



$^{40}\text{Ar}/^{39}\text{Ar}$ ages and geochemistry of Maranhão CAMP tholeiites (Brazil): implications for low and high-Ti basalts sources

R. Merle (1), A. Marzoli (1), C. Verati (2), H. Bertrand(3), M. Chiaradia (4), L. Reisberg (5), G. Bellieni (1), M. Ernesto (6)

(1) Università di Padova, Italia (renaud.merle@unipd.it, andrea.marzoli@unipd.it, giuliano.bellieni@unipd.it), (2) Géosciences Azur, France (verati@unice.fr), (3) ENS-Lyon and UCBL, France (herve.bertrand@ens-lyon.fr), (4) Université de Genève, Suisse (chiaradia@terre.unige.ch), (5) CRPG-CNRS (reisberg@crpg.cnrs-nancy.fr), (6) Universidade de São Paulo, Brazil (marcia@iag.usp.br)

The Central Atlantic magmatic province (CAMP) emplaced at the Triassic-Jurassic boundary, is composed by prevailing low-Ti ($\text{TiO}_2 < 2.0$ wt%) and rare high-Ti basalts ($\text{TiO}_2 > 2.0$ wt%). The latter were so far confined to the circum-Atlantic regions of northernmost South America (French Guyana, Surinam and Cassiporé area of Brazil) and of western Africa (Liberia, Sierra Leone). Here we report the first data for high-Ti basalts sampled up to 800 km inland in the Maranhão basin of Brazil. The Western Maranhão Basin tholeiites (WMBT) are generally evolved basalts to basaltic andesites ($\text{MgO} = 2.6\text{-}7.9$ wt%). Three distinct groups are clearly distinguished: low-Ti group ($\text{TiO}_2 < 1.3$ wt%; $5.6 < \text{MgO} < 7.9$ wt%), intermediate-Ti tholeiites ($\text{TiO}_2 \approx 2.1$ wt%; $6.6 < \text{MgO} < 7.2$ wt%) and high-Ti group ($\text{TiO}_2 = 3.4\text{-}3.7$ wt%; $2.6 < \text{MgO} < 2.7$ wt%).

The new $^{40}\text{Ar}/^{39}\text{Ar}$ plateau ages obtained on plagioclase separates for intermediate-Ti (199.7 ± 2.4 Ma) and high-Ti tholeiites (197.2 ± 0.5 Ma and 198.2 ± 0.6 Ma) are indistinguishable and are also identical to previously analyzed low-Ti tholeiites (198.5 ± 0.8 Ma) and to the mean $^{40}\text{Ar}/^{39}\text{Ar}$ age of the Brazilian CAMP (mean 199 ± 2.4 Ma).

The three chemical groups display almost flat to very slightly enriched REE patterns precluding a garnet-bearing source. Positive Pb and negative Nb anomalies are observed in the three groups. There are strong in the low-Ti WMBT patterns, less pronounced in the high-Ti WMBT and very weak in the intermediate-Ti WMBT, suggesting a variable continental contribution to the geochemical characteristics of these lavas.

The intermediate-Ti and high-Ti groups display MORB-like Sr-Pb isotopic signatures (initial $^{87}\text{Sr}/^{86}\text{Sr} = 0.7030\text{-}0.70306$, $^{206}\text{Pb}/^{204}\text{Pb} = 17.85\text{-}17.94$, $^{207}\text{Pb}/^{204}\text{Pb} = 15.50\text{-}15.52$, $^{208}\text{Pb}/^{204}\text{Pb} = 37.75\text{-}37.81$; $^{87}\text{Sr}/^{86}\text{Sr} = 0.70341\text{-}0.70356$, $^{206}\text{Pb}/^{204}\text{Pb} = 18.00\text{-}18.03$; $^{207}\text{Pb}/^{204}\text{Pb} = 15.48\text{-}15.49$; $^{208}\text{Pb}/^{204}\text{Pb} = 37.73\text{-}37.86$ respectively). Nevertheless, the intermediate-Ti samples have Nd compositions less radiogenic than MORB ($\epsilon\text{Nd} = +6.19$ and $+6.43$) but more than the high-Ti ones ($\epsilon\text{Nd} = +5.63$ and $+6.02$) and slightly radiogenic $^{207}\text{Pb}/^{204}\text{Pb}$ for low $^{206}\text{Pb}/^{204}\text{Pb}$ ratios. The low-Ti group like other low-Ti analogues of CAMP, displays more enriched signatures ($^{87}\text{Sr}/^{86}\text{Sr} = 0.706155\text{-}0.707129$, $\epsilon\text{Nd} = -1.92$ and -1.93), in particular, radiogenic $^{207}\text{Pb}/^{204}\text{Pb}$ values for relatively low $^{206}\text{Pb}/^{204}\text{Pb}$ ($^{206}\text{Pb}/^{204}\text{Pb} = 18.168$; $^{207}\text{Pb}/^{204}\text{Pb} = 15.621$; $^{208}\text{Pb}/^{204}\text{Pb} = 38.300$) suggestive of a crustal component. However, $^{187}\text{Os}/^{188}\text{Os}$ initial ratios of WMBT, ranging between 0.12830 and 0.13161 (the low-Ti group being slightly more radiogenic than the two other groups) preclude any large assimilation of continental crust. Thus, the enriched signature of the low-Ti WMBT could be due to melting of a subduction-related metasomatised subcontinental lithospheric mantle. The high-Ti and intermediate WMBT could be derived from an asthenospheric source possibly slightly contaminated by the continental crust and by the metasomatised subcontinental mantle, respectively.