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Sedimentology and paleogeography of Plio- Pleistocene Bakhtyari Formation at Ghalat and Garu- Charmakan Mountains , NW of Shiraz , Iran

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The studied region is located at Folded Zagros Zone , in which Bakhtyari Formation , as the youngest one in this Zone , has a considerable extension in the region .The region consists of some low – high anticlinal and synclinal structures which are parallel to each other and exhibit beautiful outcrops. The conglomeratic Bakhtyari and calcareous – dolomitic Asmari- Jahrum Formations overlie the synclines and anticlines respectively . Two elevated synclines (Ghalat and Garu- Charmakan Mountains) form abnormal reliefs with the height of more than 3000 meters and , contrary to the most part of the Folded Zagros , overhanging adjacent anticline , which separates them . The trend of the reliefs is NW – SE , and Cenozoic continental and marine deposits form their stratigraphic succession .In this succession , from top to bottom , a series of Formations crops out : (1) Plio- Pleistocene Bakhtyari F. (polymictic conglomerate) ; (2) Upper Miocene Aghajari F.(conglomerate , sandstone , mudstone) ; (3) Lower Miocene Gachsaran F. (marl and lenses or layers of gypsum and anhydrite in marl) ; (4) Oligocene – Miocene Asmari F.(limestone and dolomitic limestone) ; and (5) Eocene Jahrum F.(dolomite and dolomitic limestone).

To introduce conglomeratic facies of Bakhtyari F. a number of six cross sections were measured . The information resulted from these measurments showed that , from the view point of its (coarse) clasts compositon , Bakhtyari F. generally consists of two types of conglomerates : (1) Limestone – Chert Clasts Supported Conglomerate (LCCSC) , and (2) Limestone Clasts Supported Conglomerate(LCSC). The chronological and spatial relationship between these conglomerates , due to existence of some faults at their approximate boundary, are not easily distinguished, and it seems each has its own independent time and place spreading . Analytical works on sedimen-

tology shows that, in response to the uplifting of Zagros Mountaint Ranges, these coarse grained sedimentary rocks were deposited as fluvial deposits and alluvial fans by ephemeral and permanent bedload supported rivers. These deposits are usually known as alluvial facies. Since two or more lithofacies may genetically be related to each other, some researchers have described alluvial deposits as genetical facies , which may cosist of two or more lithofacies. In this research similar description method has been used and, thus, in describing the features of each conglomerate its different, but related, lithofacies have been described likewise . LCCSC consists of three different lithofacies : (1) Conglomerate(the dominant facies), in which the composition is closely related to grain size. Coarser grains are of limestone composition and finer grains are of chert composition; (2) sandstone, and (3) mudstone. Conglomeratic facies, in view point of its clasts size, is divided into two subfacies : (a) boulder- cobble conglomerate with the composition of dominantly limestone, and (b) pebble – granule (microconglomerate) facies, with the composition of dominantly chert. LCSC consists of two different lithofacies : (1) Conglomerate, which is the dominant facies of LCSC (more than 90% of the clasts are of limestone composition), and (2) sandstone, which is subordinate.

Epirogenic movements in Paleogene caused some fluctuations in the sea level of that time , which resulted in a bed of chert microconglomerate at the Jahrum- Asmari contact ; so the conglomerate as the coarse clastic sediments flew in the existing basin in Paleogene – Neogene boundary. The regression of sea in Bourdigalian caused a set of continental environments (lacustrine , fluvial and alluvial deposits) to occur. These environments, during their existence , recieved a large quantity of sediments, which was due to the erosion of newly uplifted highlands in neighbouring northern parts of Zagros Zone. Accumulation of these sediments caused a pile of them , which its vertical succession thickness reached to more than 5000 meters. The upper part of this succession forms Bakhtyari Formation , which its textural characterestics and sedimentary structures specify a non- marine depositional system and considered as " standard" lithofacies of alluvial fans and braided rivers. Researches on alluvial structures have been done in a large number of tectonic settings , one of which is forland basin. Tradationally , the deposits of forland basins are considered as continental molasse.