



Distribution and partitioning of trace elements during crustal anatexis: a LA-ICP-MS study of metapelitic enclaves within El Hoyazo dacite, SE Spain

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The El Hoyazo volcano is located in the Neogene Volcanic Province (NVP) of south-eastern Spain, consists of a pipe and surrounding block lava, and is composed of peraluminous ($ASI \approx 1.5$) dacites. Dacites contain frequent enclaves ($\approx 1-2$ volume %) of mainly anatectic metapelites and gabbros/basalts (Zeck 1992). One of the main types of metapelitic enclaves, the Grt-Bt-Sil type (Zeck 1970), represents fragments of middle-to-lower continental crust, partially melted at $\sim 800-850^\circ\text{C}$ and 5-7 kbar, depleted in a granitic component after melt extraction, and brought to surface from a partially molten state within the host dacites ≈ 6 Ma ago (Cesare et al. 2003). Grt-Bt-Sil enclaves are composed of Pl (35-45 vol. %), Sil and glass intimately intergrown (25-35%), Bt (5-15%), Grt (5-10%), glass (former melt phase; 5-10%), Crd (1-5%), Gph (1-2%), Afd (1-2%), and accessory minerals (Apt, Ilm, Mnz, Zrn; $< 1\%$). Silicate glass is present in the rock matrix and within most of minerals as melt inclusions (MI). Melt abundance before segregation has been estimated to be $\approx 30-55\%$ (Cesare et al. 1997). EMP analyses of glasses yielded leucogranitic compositions with small but systematic differences among each textural location (Acosta-Vigil et al. 2007). Recent LA-ICP-MS analyses of glass show differences in trace element composition among each textural location as well. Trace element partitioning during anatexis of enclaves occurred as follows. **Li:** Crd>melt>Pl>Bt. **Rb:** Bt>melt. **Cs:** melt>Bt>Crd. **Be:** Crd>>melt. **B:** melt>Pl>>Bt. **Sr:** Pl>>Afd>melt. **Ba:** Afd>Bt>>melt>Pl. **Sc:** Grt>>Bt>>melt. **V:** Bt>>Grt. **Cr:** Bt>>Grt>Pl>melt. **Co:** Bt>Grt>Crd. **Ni:**

Bt>>**Crd**≈**Pl**. **Cu**: **Bt**≈**Pl**>melt. **Zn**: **Bt**>**Crd**>**Grt**≥melt. **As**: melt>>**Pl**. **Nb, Ta**: **Bt**>>melt. **Y, Zr, Hf, U, Th, REE**: accessory phases (**Mnz, Apt, Zrn**). Considering the above modal abundance of phases in the enclaves, trace element distribution during anatexis occurred as follows. **Li**: melt>**Pl**>>**Bt**. **Rb**: melt>**Bt**. **Cs**: melt>>**Bt**. **Be**: **Crd**≈melt. **B**: melt>>**Pl**. **Sr**: **Pl**>>melt. **Ba**: **Bt**>melt>**Pl**≈**Afd**. **Sc**: **Grt**>**Bt**≈melt. **V**: **Bt**>>**Grt**. **Cr**: **Bt**>>**Pl**>**Grt**≈melt. **Co**: **Bt**>>**Grt**>melt. **Ni**: **Bt**>>**Pl**>melt. **Cu**: melt≈**Pl**>**Bt**. **Zn**: **Bt**>melt>>**Grt**≈**Crd**≈**Pl**. **As**: melt>>**Pl**. **Nb, Ta**: **Bt**>melt. **Y, Zr, Hf, Th, U, REE**: accessory phases>>>melt, except for **Y** and **HREE** hosted also to some extent in **Grt** (“>” means approximately double, “>>” means approximately one order of magnitude larger). Hence, with respect to the starting source rock melt was largely enriched in **Li, Cs, B** and **As**, moderately enriched in **Rb**, likely with comparable concentrations in **Be**, moderately impoverished in **Ba, Sc, Cu** and **Nb**, and largely impoverished in **Sr, V, Cr, Co, Ni, Zn, Y, Zr, U, Th** and **REE**. Calculated **Zrn** and **Mnz** saturation temperatures are ≈100-200°C lower than those obtained from conventional geothermometry, and for many **MI** are even lower than the wet granite solidus. The disagreement may be explained by rapid melting and melt solidification right after (i.e. short-lived melting event), and/or armoring of **Zrn** and **Mnz** in restitic major mineral phases, and/or rapid (re-)crystallization of restitic phases and armoring of melt in **MI**.