



## **Carbon transfers in tropical stream water in a volcanic context: examples of a hot spot island (Reunion Island) and two subduction arc islands (Guadeloupe and Martinique, French Lesser Antilles)**

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At long-time scales, the global cycle of CO<sub>2</sub> is controlled by chemical denudation of rocks. One way to evaluate the CO<sub>2</sub> consumption rate by chemical weathering is to consider bicarbonate concentrations. However, transfers of carbon in rivers are poorly constrained both for the sources of carbon and processes that may affect dissolved inorganic carbon (DIC), for example CO<sub>2</sub> degassing, carbonate precipitation or dissolution, photosynthesis, organic matter decay.

In order to better investigate the carbon behaviour in rivers, we have used  $\delta^{13}\text{C}_{DIC}$  as a source and processes tracer. Because basaltic areas have high chemical weathering rates comparatively to other silicate places, we have focus our study on basaltic/andesitic-like single-lithology areas with no sedimentary carbonates (Guadeloupe and Martinique Islands, and Reunion Island). These tropical islands have important rainfall, thick soils and sharp relief, which induced high weathering rates (Rad et al, 2005). Main streams, vegetation and soil solutions were sampled.

The  $\delta^{13}\text{C}_{DIC}$  of the sampled streams ranges from  $-19.6\text{‰}$  to  $-5.2\text{‰}$ , with DIC ranging from 110 to 2540  $\mu\text{M}$ . In absence of carbonate rocks, potential sources of carbon in these streams are atmospheric CO<sub>2</sub> ( $\delta^{13}\text{C}_{CO_2} = -7.7\text{‰}$ ), organic matter ( $\delta^{13}\text{C}_{COM} = -29\text{‰}$ ), and magmatic CO<sub>2</sub> ( $\delta^{13}\text{C}_{CO_2} = -3.5\text{‰}$ ). The Reunion Island data show a quite different  $\delta^{13}\text{C}_{DIC}$  vs. DIC trend from that of the Antilles Islands. The Reunion trend can be mainly explained by a mixing between soil solutions (carbon from

organic matter) and magmatic CO<sub>2</sub>. This influence of magmatic contribution is confirmed by major elements like Mg/Na vs. Ca/Na (Louvat, 1997 and Rad et al., 2005), and the  $\delta^{13}\text{C}_{DIC}$  vs DIC allows us to estimate the different contributions of magmatic/organic sources in absence of other processes. Moreover, a variable degassing of dissolved CO<sub>2</sub> from waters is likely to occur in addition of mixing, more in Lesser Antilles than in Reunion Island (profile of river shows lower  $\delta^{13}\text{C}_{DIC}$  at the source of the river). This behaviour of DIC should be considered for recalculation of CO<sub>2</sub> consumption rate by chemical weathering.

**References:**

S. Rad, et al. 2005. Journal of Geochemical Exploration;

P. Louvat, and C. J. Allègre, 1997. Geochimica et Cosmochimica Acta, Vol. 61, 17, pp 3645-3669