



The temperature dependent strontium isotope fractionation ($\delta^{88/86}\text{Sr}$) during calcium carbonate precipitation

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In order to study the influence of stable isotope fractionation during inorganic and biologically controlled CaCO_3 precipitation we've developed the analytical principles for the measurement of Strontium (Sr) isotope fractionation. We've established a measurement protocol for the application on a MC-ICP-MS (*AXIOM*) using the common bracketing standard technique. The Sr CRM NBS987 was used as reference material for all measurements and to calculate the Sr fractionation. Latter value is expressed by the δ -notation defined as:

$$\delta^{88/86}\text{Sr} = [({}^{88}\text{Sr}/{}^{86}\text{Sr})_{\text{sample}} / ({}^{88}\text{Sr}/{}^{86}\text{Sr})_{\text{standard}}] * 1000 - 1.$$

A first set of experiments focused on the temperature dependency of Sr-isotope fractionation. For this purpose we've studied inorganically precipitated aragonite and calcite prepared under controlled conditions in a temperature range from 10 to 50 °C. In addition, cultured and naturally grown corals were analyzed for their $\delta^{88/86}\text{Sr}$ values.

Repeated measurements of IAPSO seawater standard showed a mean $\delta^{88/86}\text{Sr}$ value of 0.383 ± 0.008 (2SEM) being the isotopically heaviest material measured so far. The first results of the inorganically precipitated aragonite and the natural corals revealed a clear temperature dependency of the $\delta^{88/86}\text{Sr}$ values. For inorganic aragonite the slope of this correlation is about 0.0055 *permille*°C. However, for naturally grown corals a 6 fold steeper slope of 0.033 *permille*°C was determined. This difference between inorganic and organic aragonite is still under debate and will be studied by further laboratory experiments.