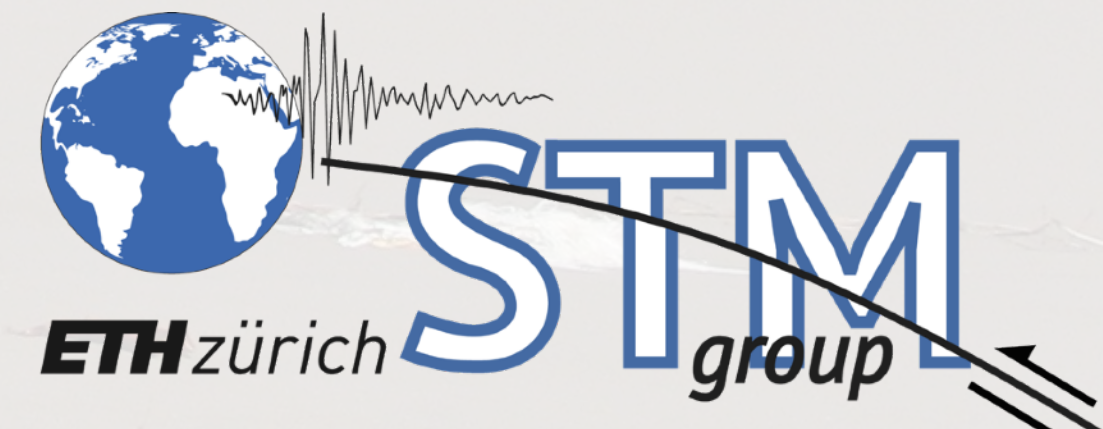

On How Lithosphere and Mantle Dynamics affect Shallow Earthquakes

Ylona van Dinther

Mario D'Acquisto, Sylvia Brizzi, Simon Preuss,
Lukas Preiswerk, Luca Dal Zilio, Iris van Zelst,
Robert Herrendörfer, Taras Gerya and collaborators



Utrecht University



» Illustrate how lithosphere- and mantle dynamics, structure and rheology can influence shallow tectonics and seismicity

• Does incoming sediment thickness increase maximum earthquake magnitude?

- Yes; larger sedimentary wedge → trench moves seaward and slab unbends → slab dip ↓ → seismogenic width ↑ → M_{max} ↑
- Modeling long-term dynamics and sediment presence increases M_{max} by an order of magnitude !

• Do lower crustal and mantle depth temperature and rheology affect seismicity and tectonics?

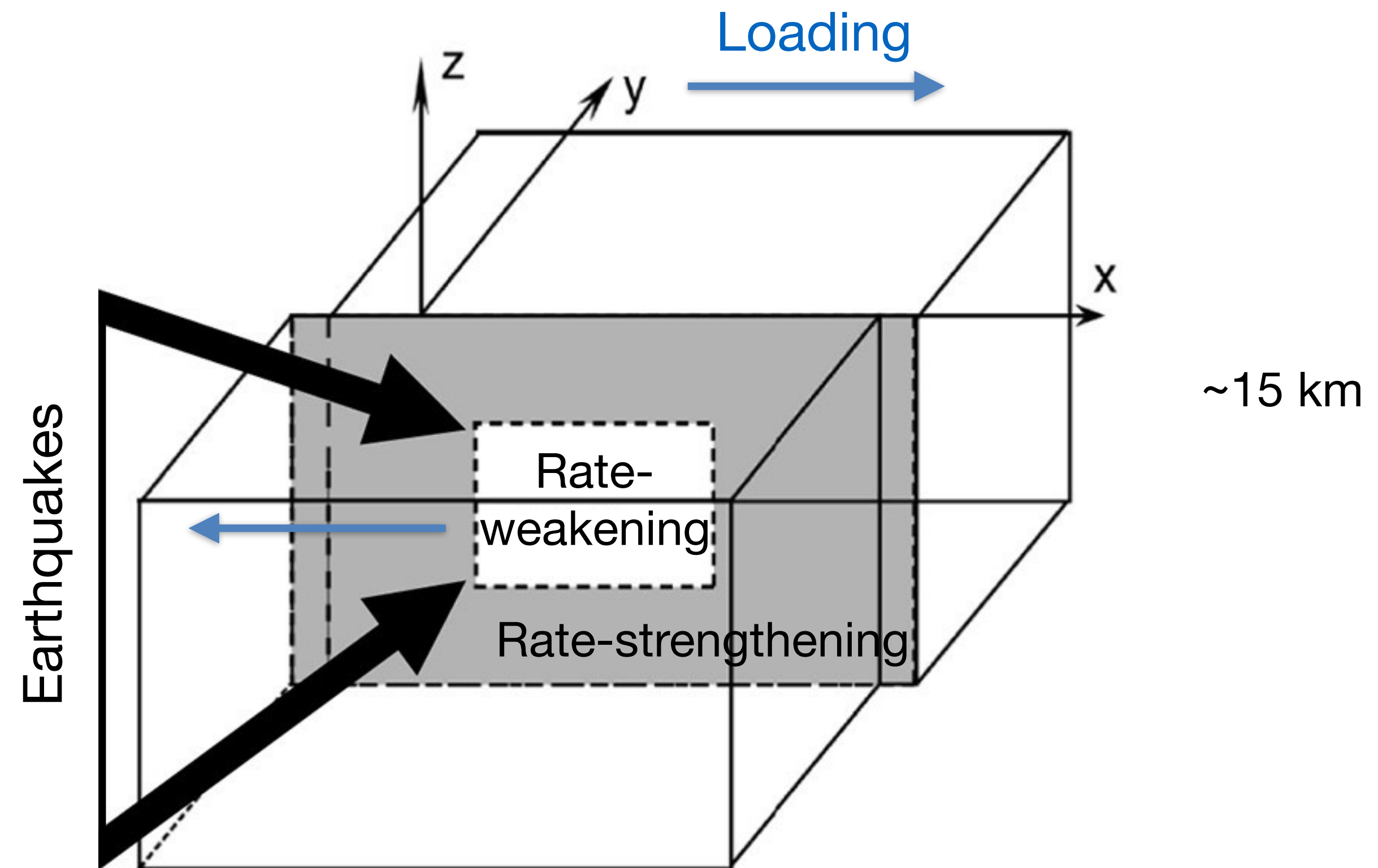
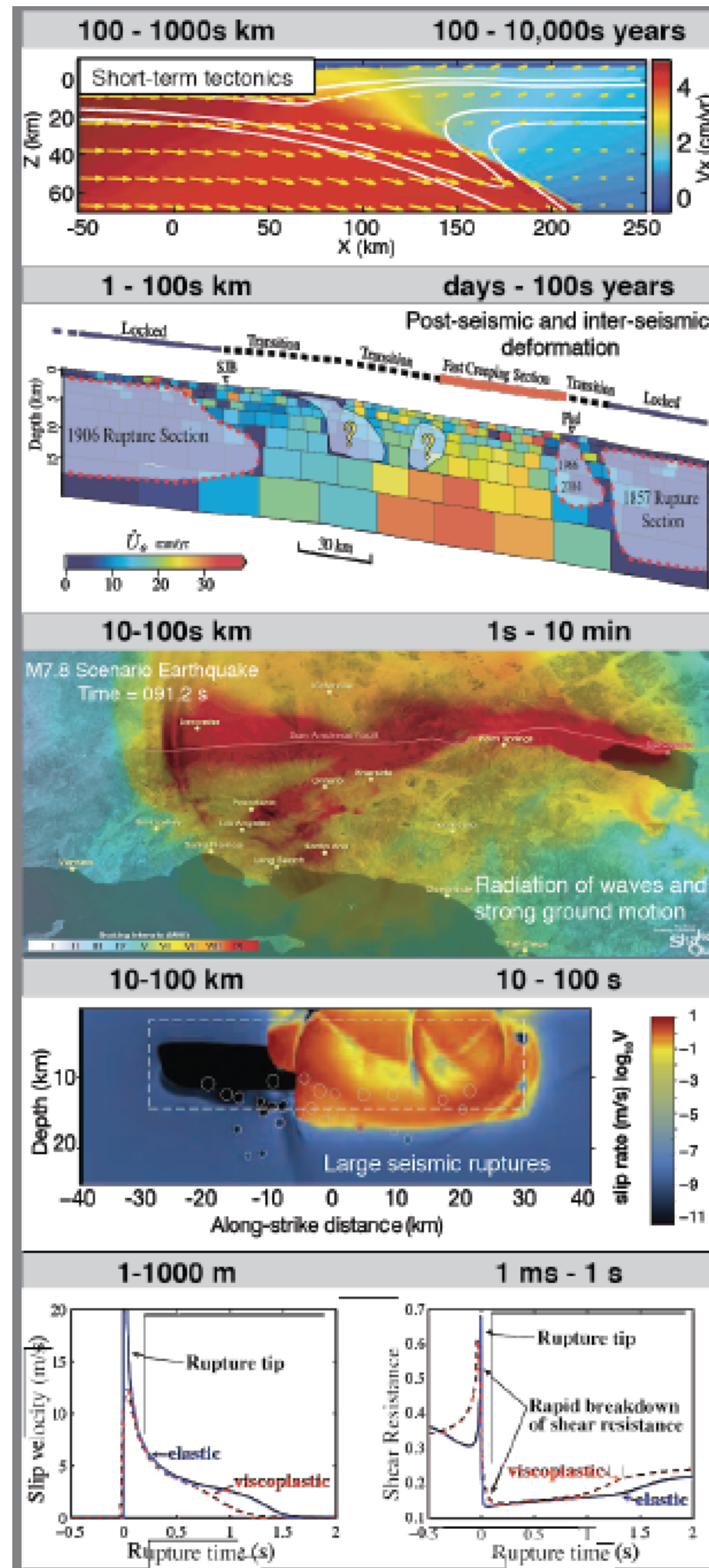
- Yes, e.g., in tectonic settings driven by complex loading, such as Northern Apennines

• Does the mantle affect surface displacements at time scales of minutes to days?

- Yes, STM models predicted a secondary zone of “coseismic” uplift, which was confirmed by observations of 4 out of 4 great megathrust earthquakes
- Accelerated slab penetration causes upward return “flow”

An upper crustal perspective on earthquakes

- Most for simplicity ignore what happens below upper crust. For what settings does that hold?

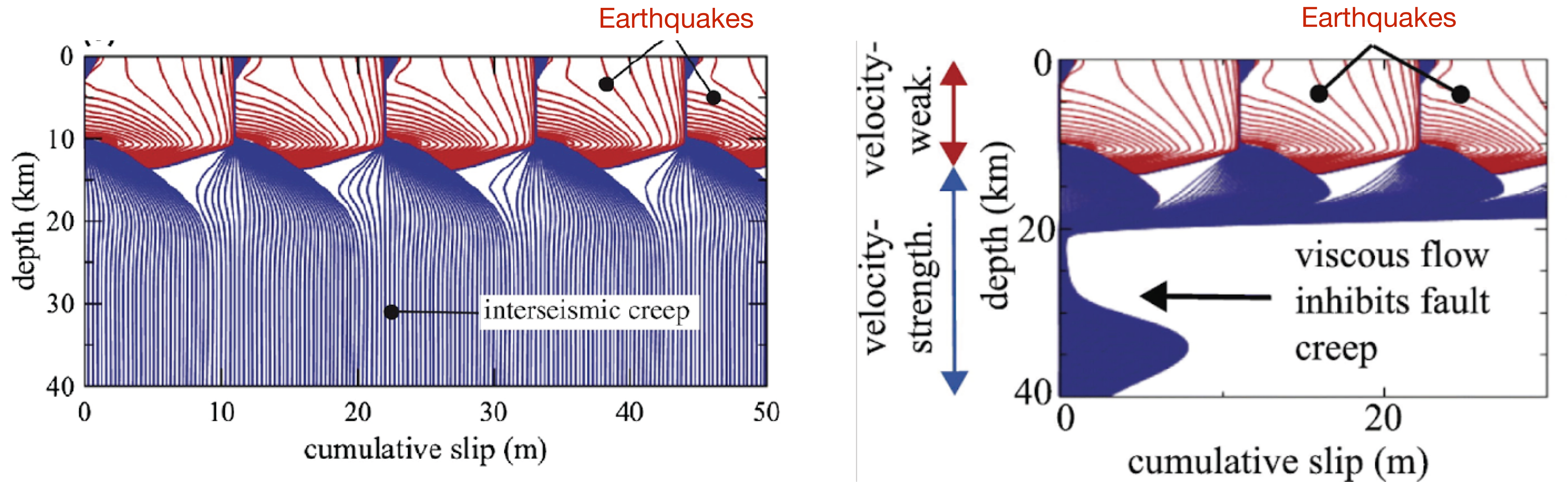


E.g., Schaal & Lapusta, JGR, 2019

- » Maybe not for orogens, subduction zones, mid ocean ridges, and likely also not really in strike-slip faults...

State-of-the-art in modeling seismic cycles

- First implementation of a powerlaw viscous rheology (Allison and Dunham, 2018):

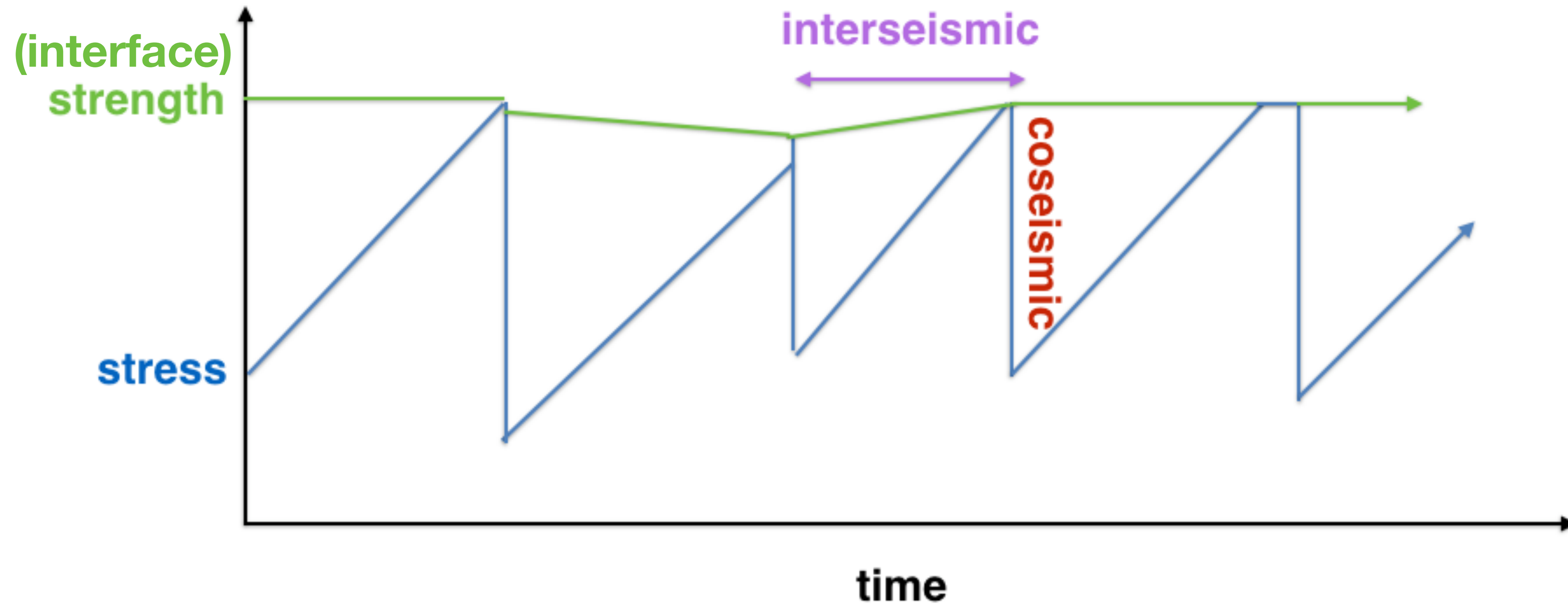


- Challenges next decade (Lapusta, Dunham, Avouac, Denolle, van Dinther, Faulkner Fialko, Katijama et al., NSF, 2019):
 - **Fluids, inelasticity, structural complexity lithosphere, shear heating, chemical reactions, thermomechanical coupling**

» We can join forces!

Seismic cycles

- Earthquakes occur when fault stress exceed its strength
 - » **Stress and strength thus regulate earthquake nucleation, propagation and arrest**



What affects (a)seismic slip?

rheology
 μ, ν, η

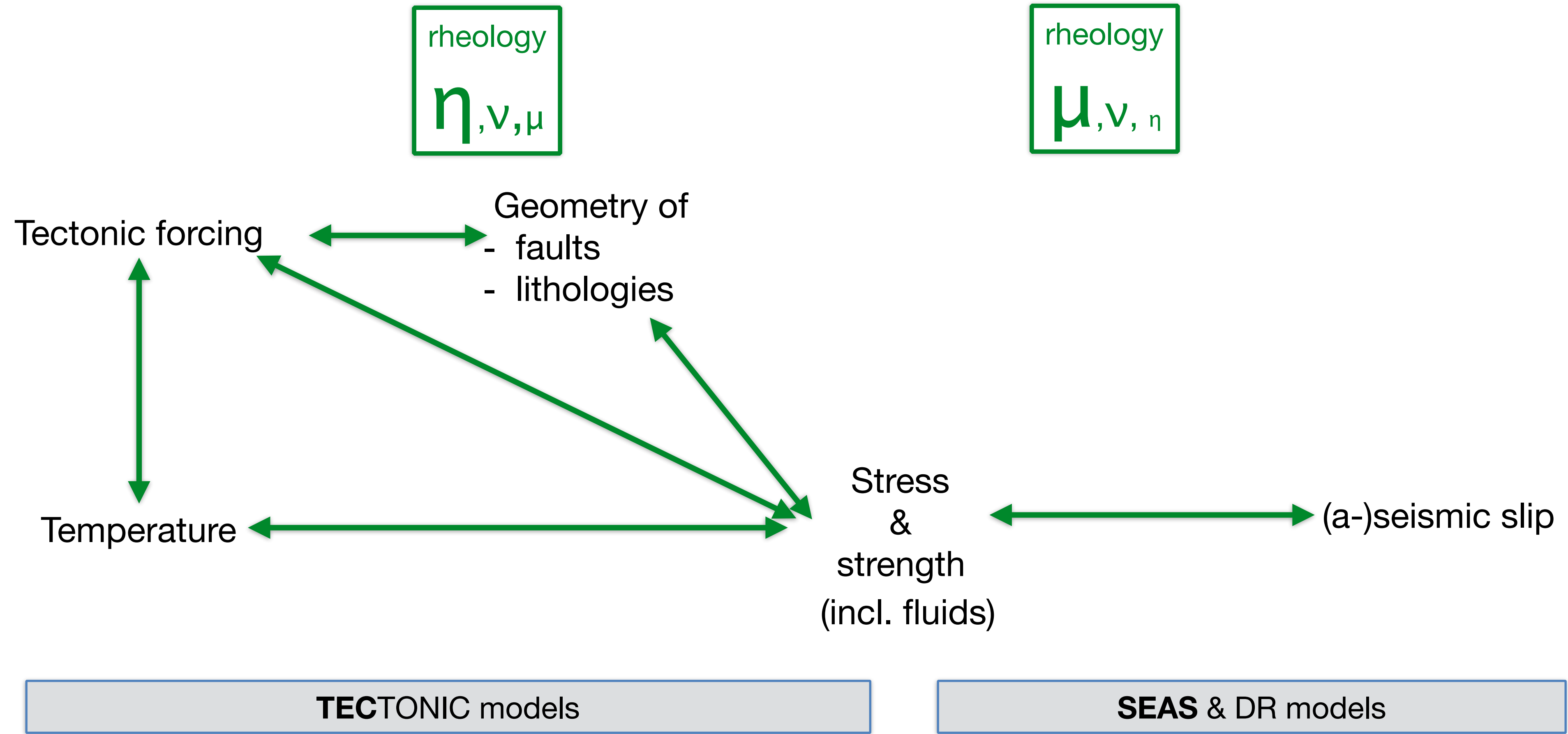
Stress & strength \longleftrightarrow (a-)seismic slip

SEAS & DR models

SEAS = sequences of SEismic and Aseismic Slip
aka "seismic cycle"

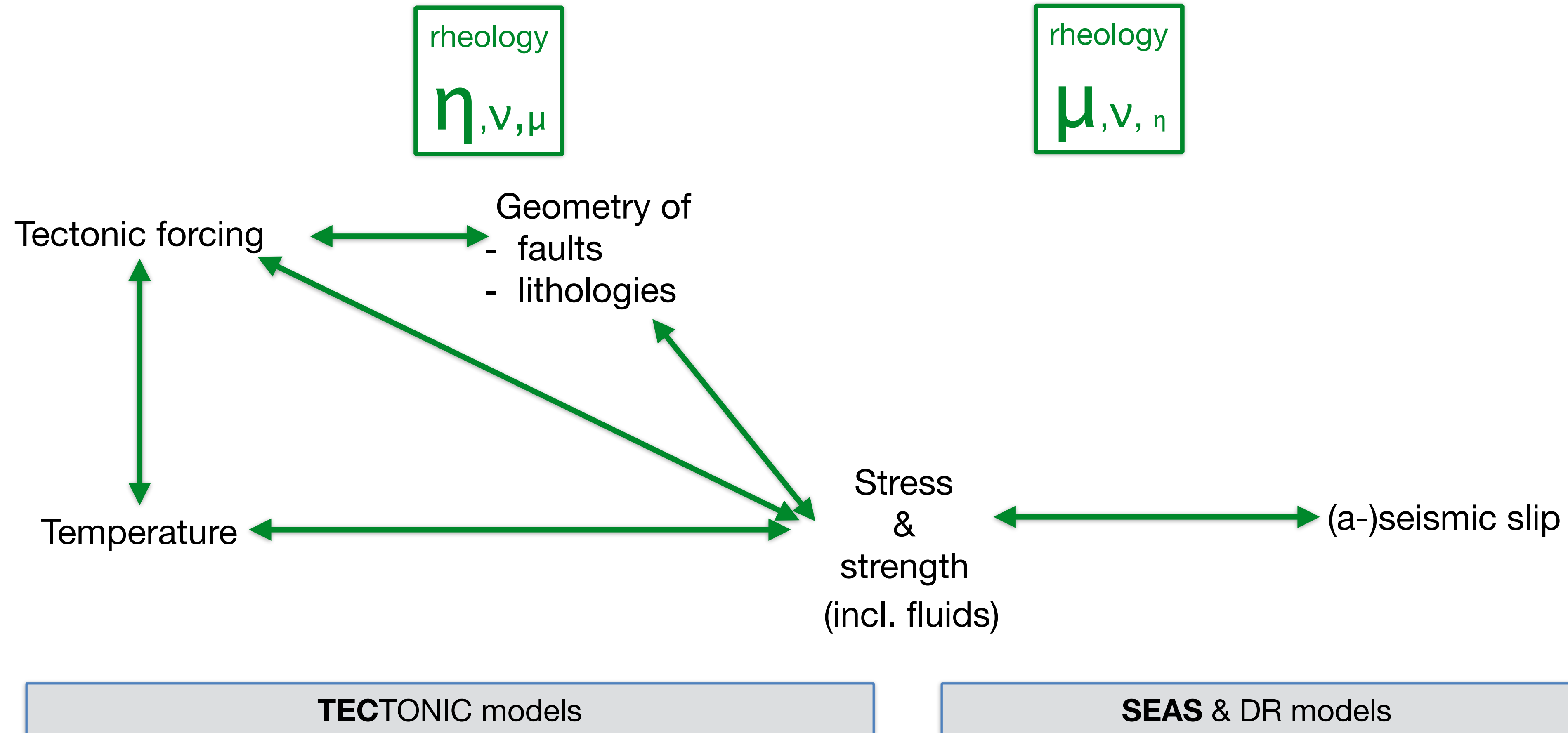
DR = Dynamic earthquake Rupture

What affects (a)seismic slip?



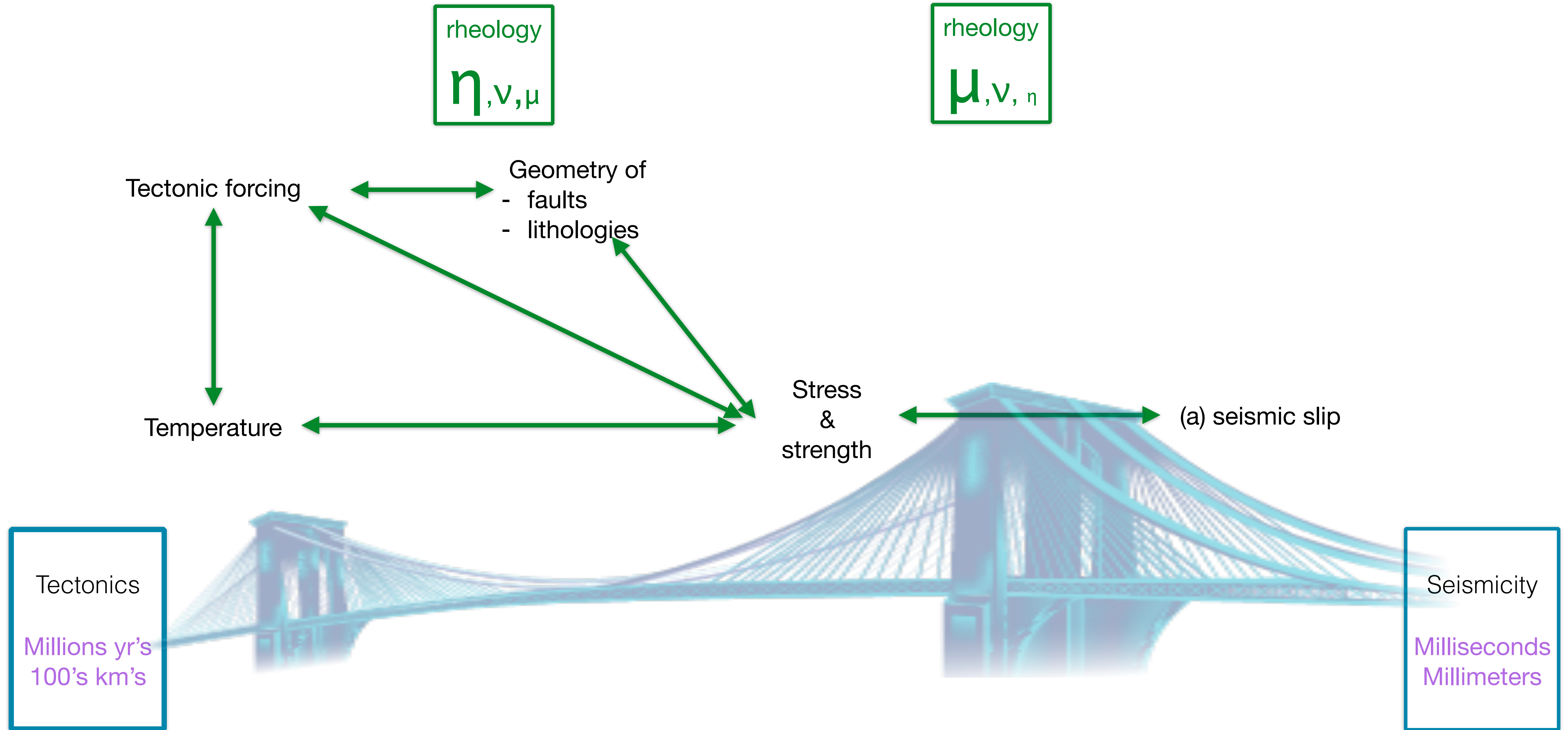
SEAS = sequences of SEismic and Aseismic Slip
aka "seismic cycle"
DR = Dynamic earthquake Rupture

Complex interaction of processes controlling (a)seismic slip

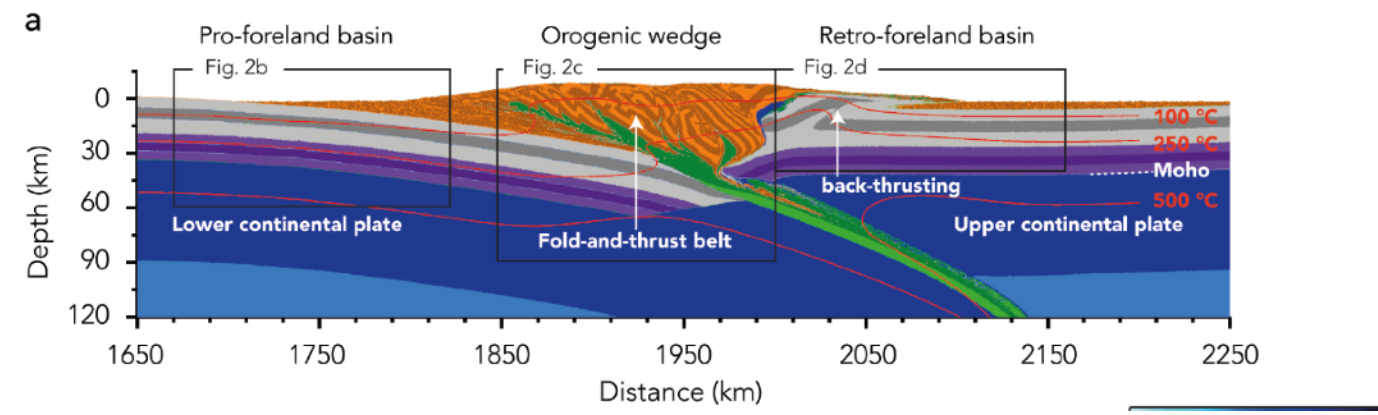


- Both feedback networks have important role for rheology, where material properties are a function of stress, temperature, fluids,...
- Complex, non-linear interactions require spontaneous simulation on both processes on both ranges of time-scales

Bridging time scales from tectonics to dynamic rupture



Bridging time scales from tectonics to dynamic earthquake rupture



Geodynamic evolution

$\Delta t = 1000$ years

Tectonics

Millions yr's
100's km's



Seismicity

Milliseconds
Millimeters

Thermo-Mechanical models (TM)

- Based on 2D finite-difference with marker-in-cell code

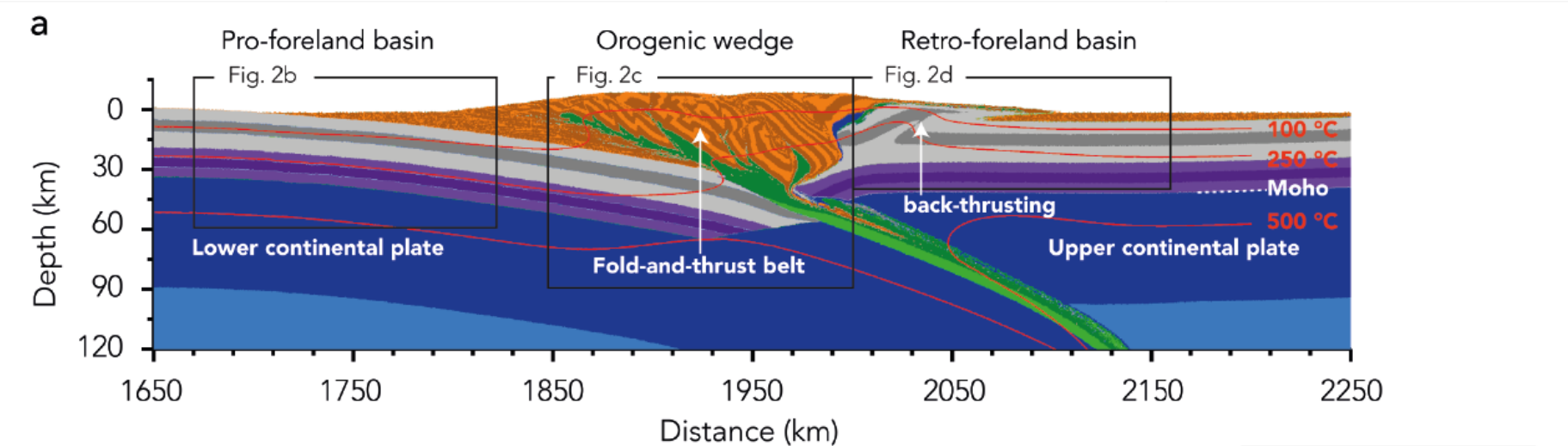
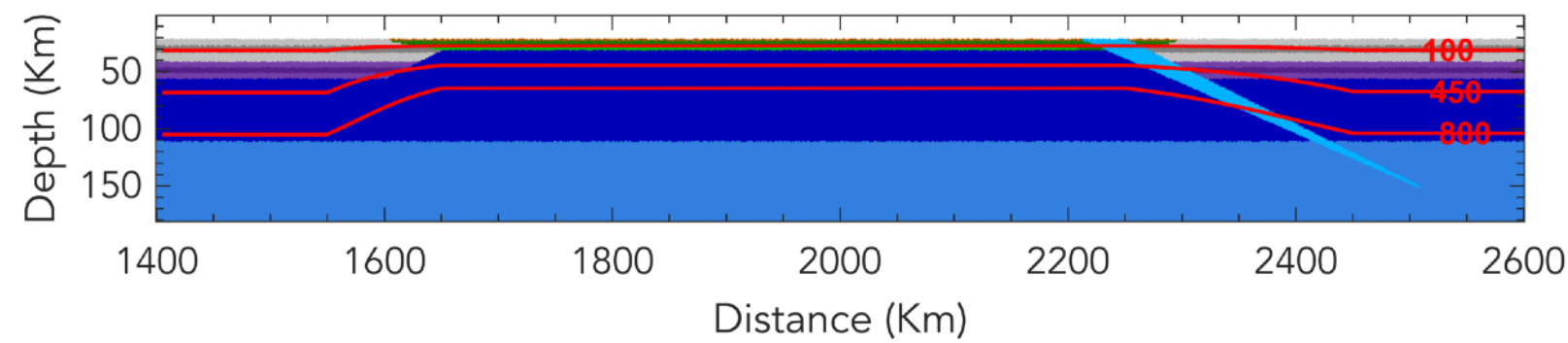
- **Input**

- Initial geometry and temperature
- Tectonic parameters
- Material parameters rock types

- **Tectonic output**

- Geometry
- Distribution physical parameters
 - Viscosity, temperature, stress, fluid pressure

Conservation of mass, momentum and energy
Visco-elasto-plastic rheology



Seismo-Thermo-Mechanical models (STM)

- Based on 2D finite-difference with marker-in-cell code

- **Input**

- Initial geometry and temperature
- Tectonic parameters
- Material parameters rock types

- **Tectonic output**

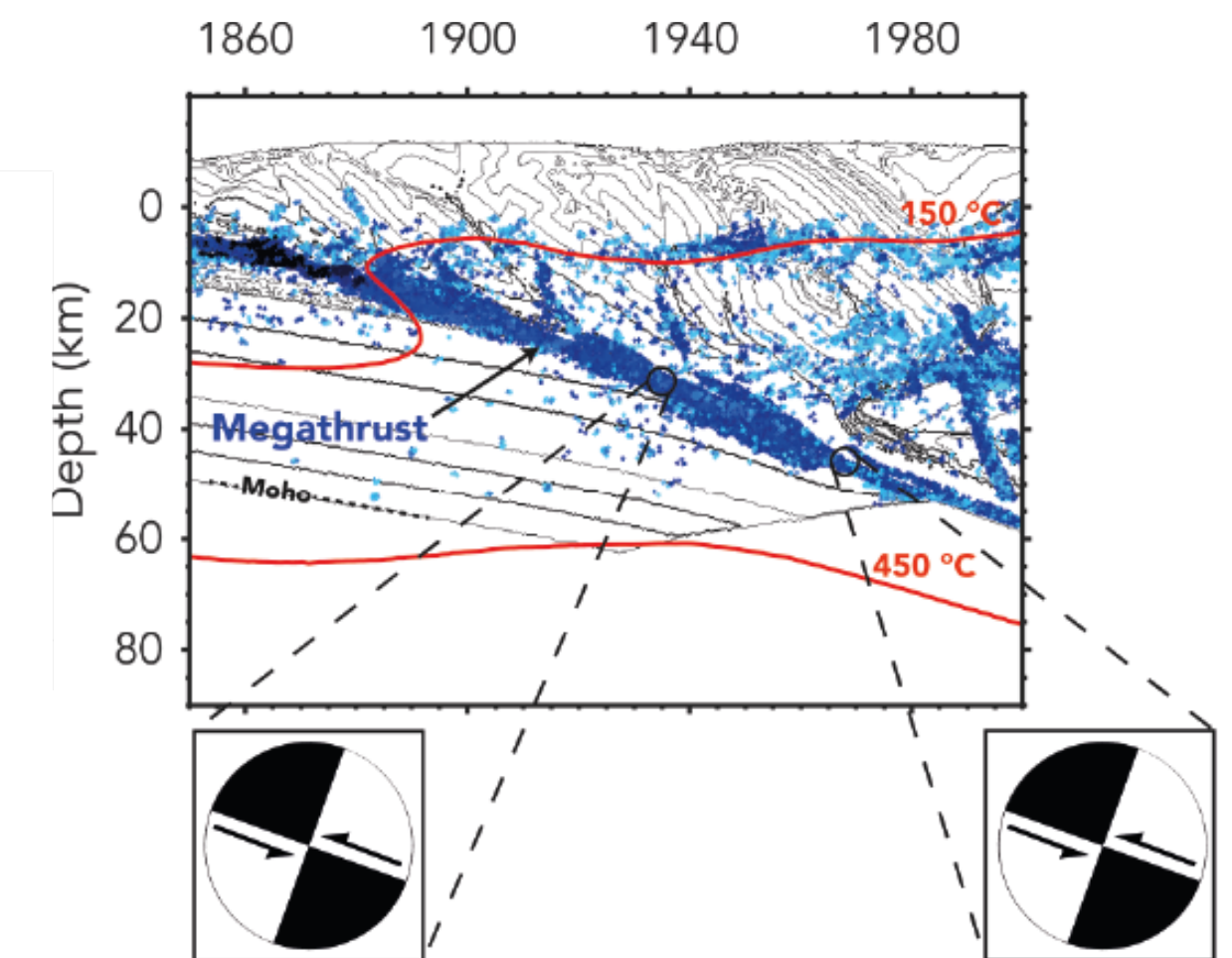
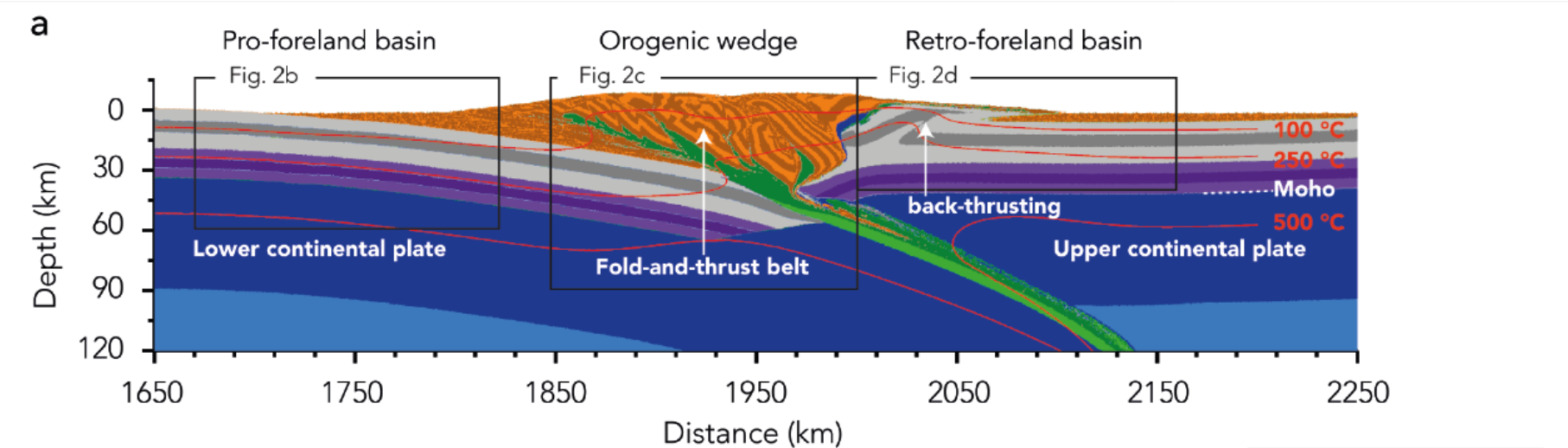
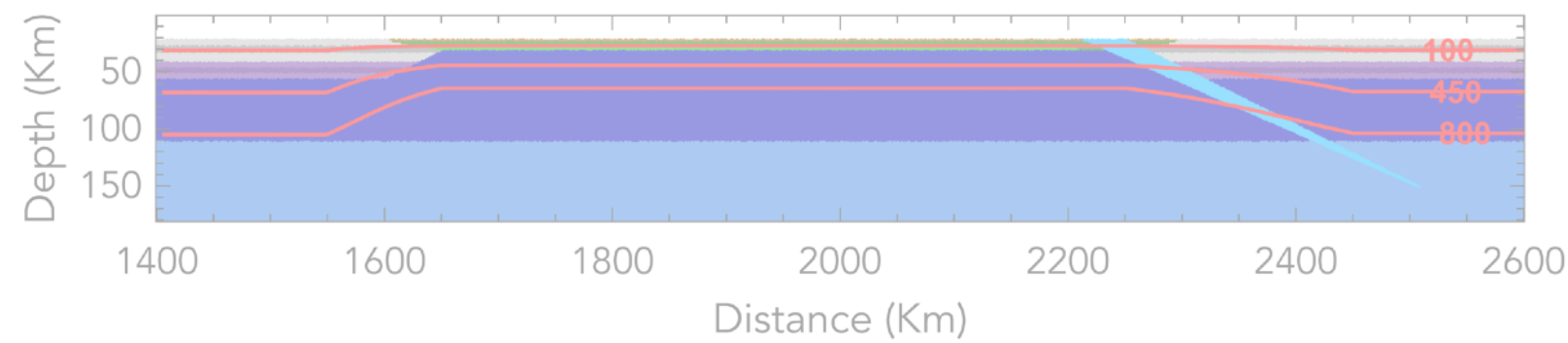
- Geometry
- Distribution physical parameters
 - Viscosity, temperature, stress, fluid pressure

- **Seismicity output**

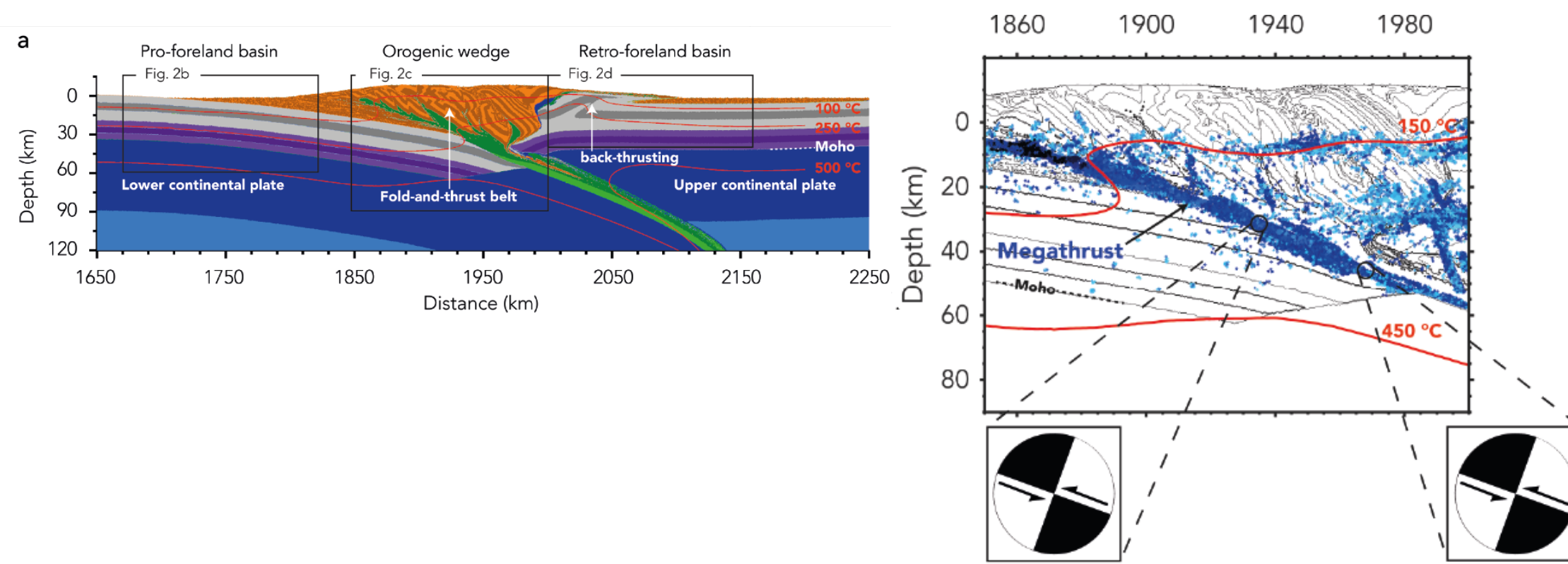
- Earthquake nucleation, propagation, arrest

Conservation of mass, momentum and energy
Visco-elasto-plastic rheology

+ inertia
+ rate(-and state) dependent friction



Bridging time scales from tectonics to dynamic earthquake rupture



Slip rate-dependent friction

$\Delta t = 1-5$ years

Tectonics

Millions yr's
100's km's

Seismicity

Milliseconds
Millimeters

- Brittle response mimicked by Drucker-Prager plasticity

- Localizes deformation when σ'_{II} reaches strength $\sigma_{yield} = C + \mu \cdot \left(1 - \frac{P_{fluid}}{P_{solid}}\right) \cdot P$

Short-term rheology: strongly slip rate dependent friction

- Brittle response mimicked by Drucker-Prager plasticity

- Localizes deformation when σ'_{II} reaches strength $\sigma_{yield} = C + \mu \cdot \left(1 - \frac{P_{fluid}}{P_{solid}}\right) \cdot P$

Slip rate $V = 2\dot{\epsilon}'_{II(p)} \Delta x$

Friction $\mu_{eff} = \mu_s(1 - \gamma) + \mu_s \frac{\gamma}{1 + \frac{V_{vp}}{V_c}}$

Short-term rheology: regularized rate-and-state dependent friction

- Brittle response mimicked by Drucker-Prager plasticity

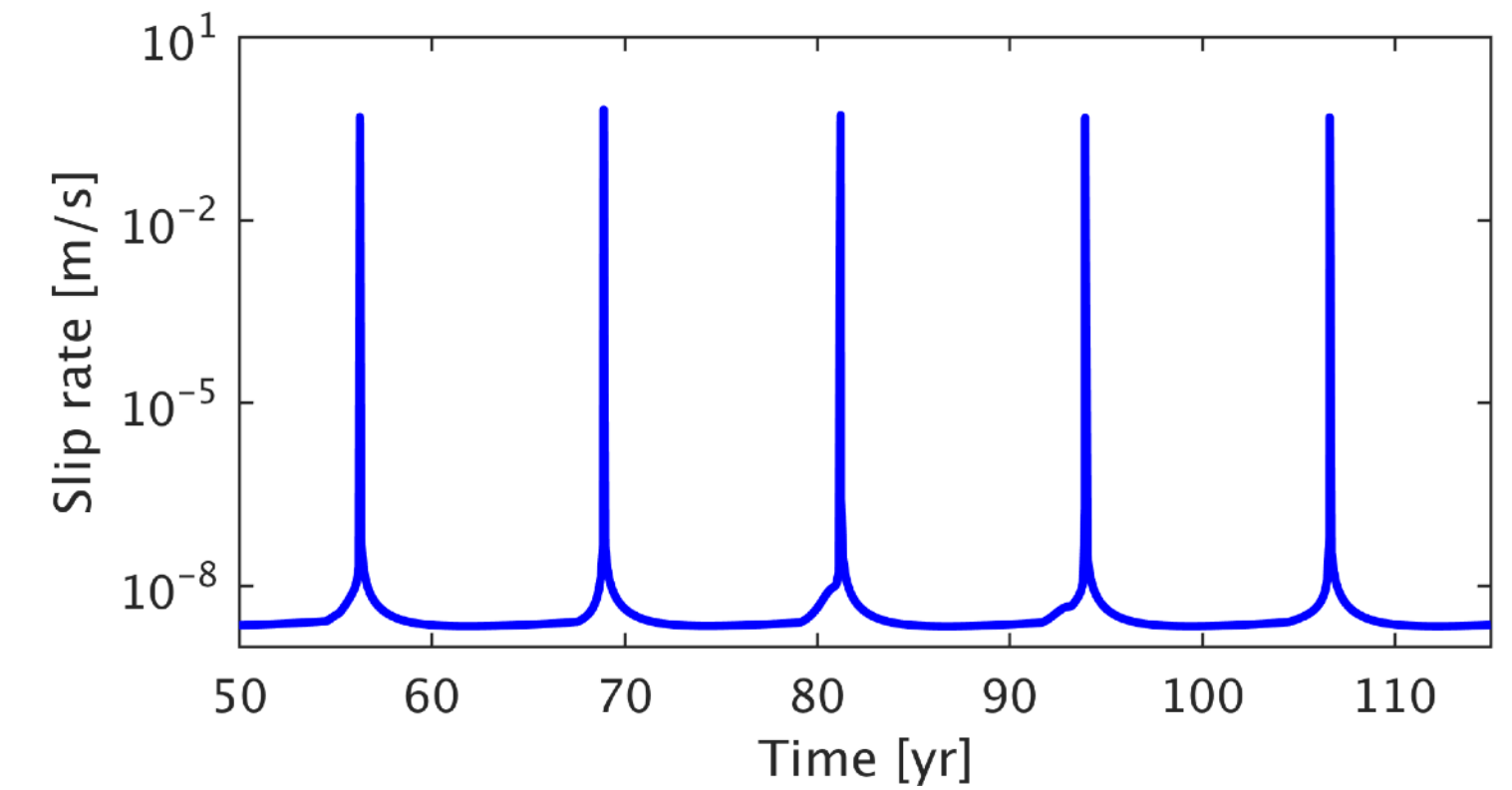
- Localizes deformation when σ'_{II} reaches strength $\sigma_{yield} = C + \mu \cdot \left(1 - \frac{P_{fluid}}{P_{solid}}\right) \cdot P$

Slip rate $V = 2\dot{\epsilon}'_{II(p)} \Delta x$

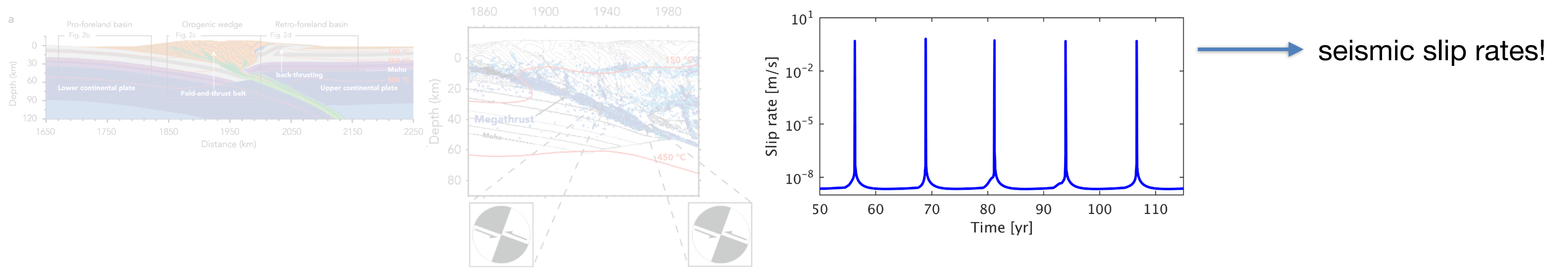
Friction $\tau_{II} = \sigma_{yield} = a P \operatorname{arcsinh} \left[\frac{V_p}{2V_0} \exp \left(\frac{\mu_0 + b \ln \left(\frac{\theta V_0}{L} \right)}{a} \right) \right]$

State evolution $\frac{d\theta}{dt} = 1 - \frac{V_p \theta}{L}$

with adaptive time stepping and Global picard iterations



Bridging time scales from tectonics to dynamic earthquake rupture



Rate-and-state dependent friction - invariant reformulation

Resolve interseismic, coseismic and postseismic phase

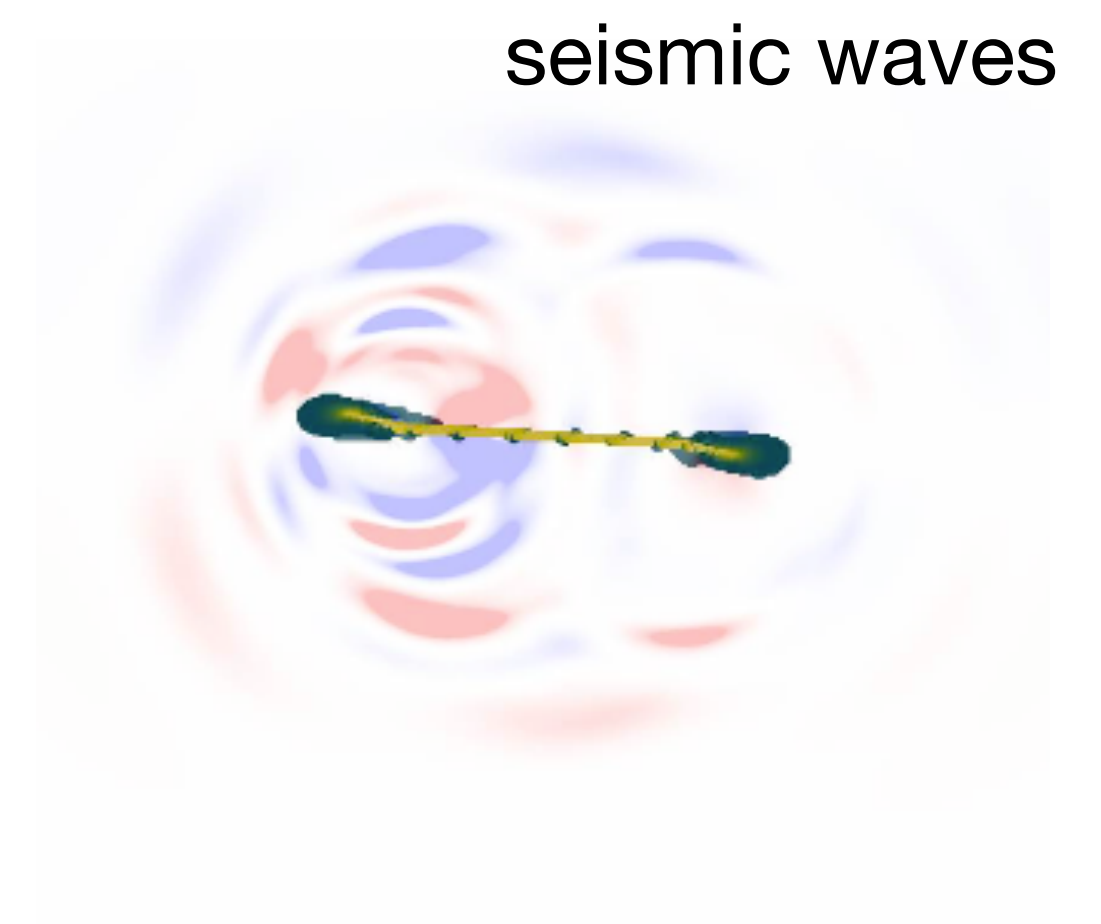
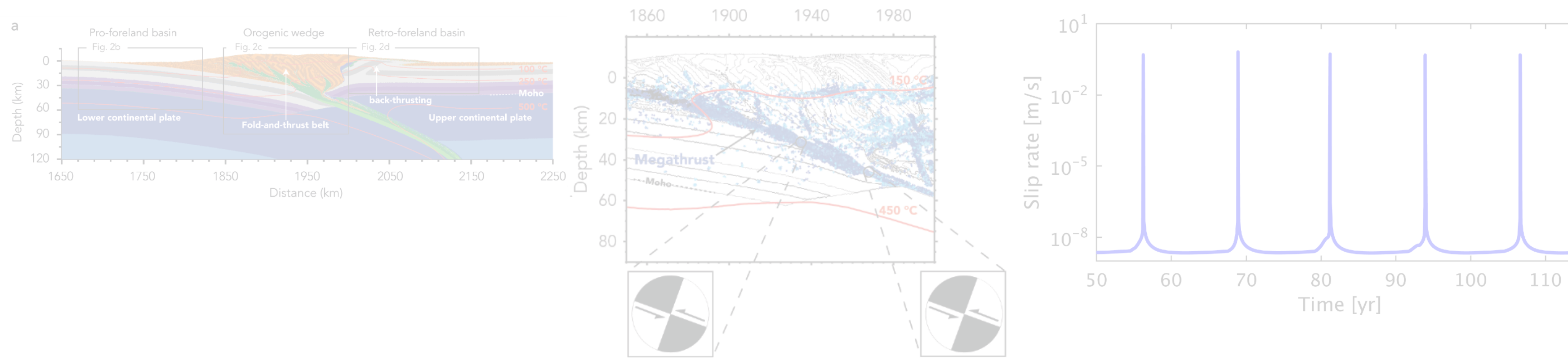
Simulate whole slip spectrum: a-, slow-, seismic slip

Δt = milliseconds - years

Tectonics
Millions yr's
100's km's

Seismicity
Milliseconds
Millimeters

Bridging time scales from tectonics to dynamic earthquake rupture



Dynamic earthquake rupture

Fault evolution

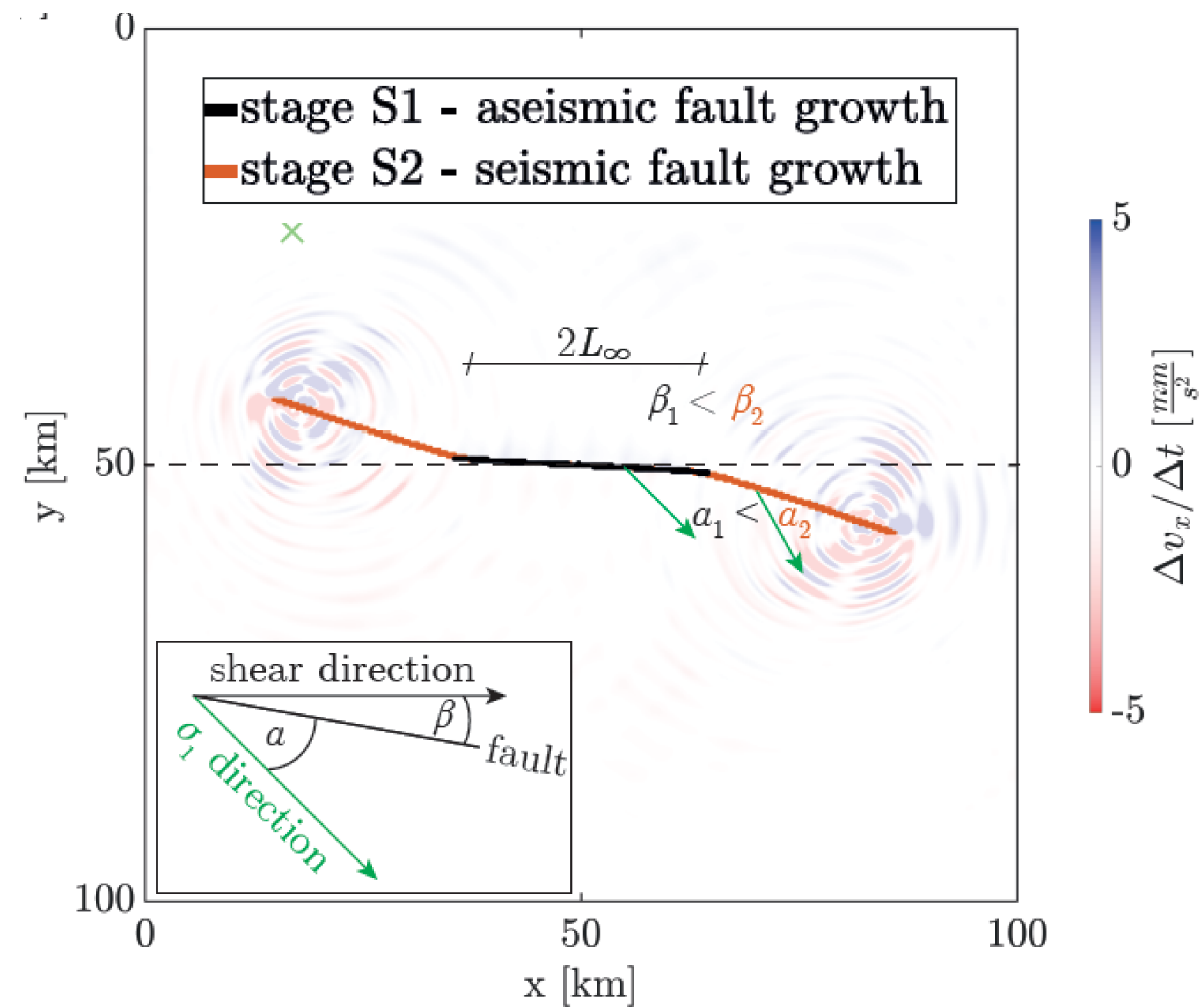
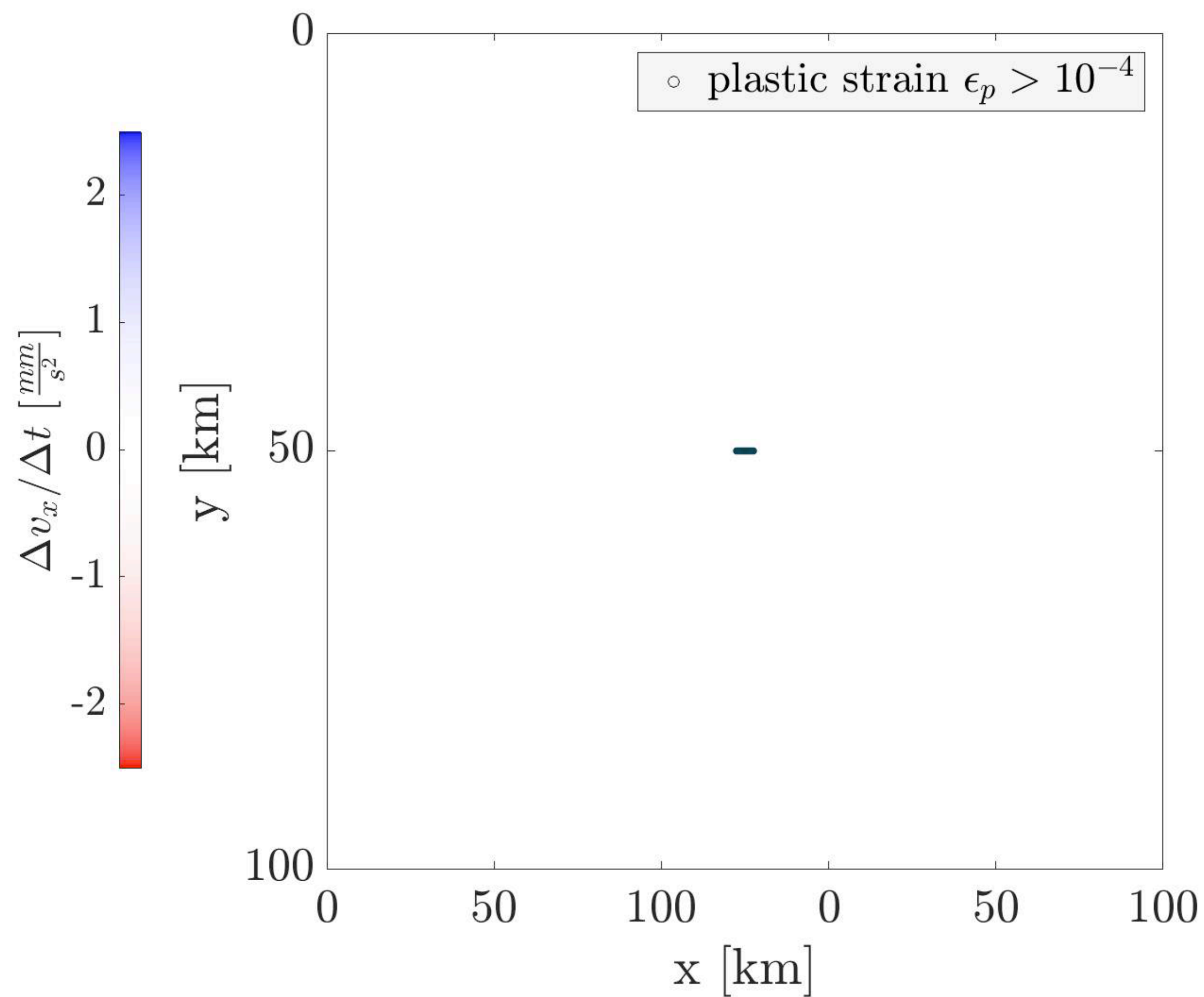
$\Delta t = \text{milliseconds}$

Tectonics
Millions yr's
100's km's

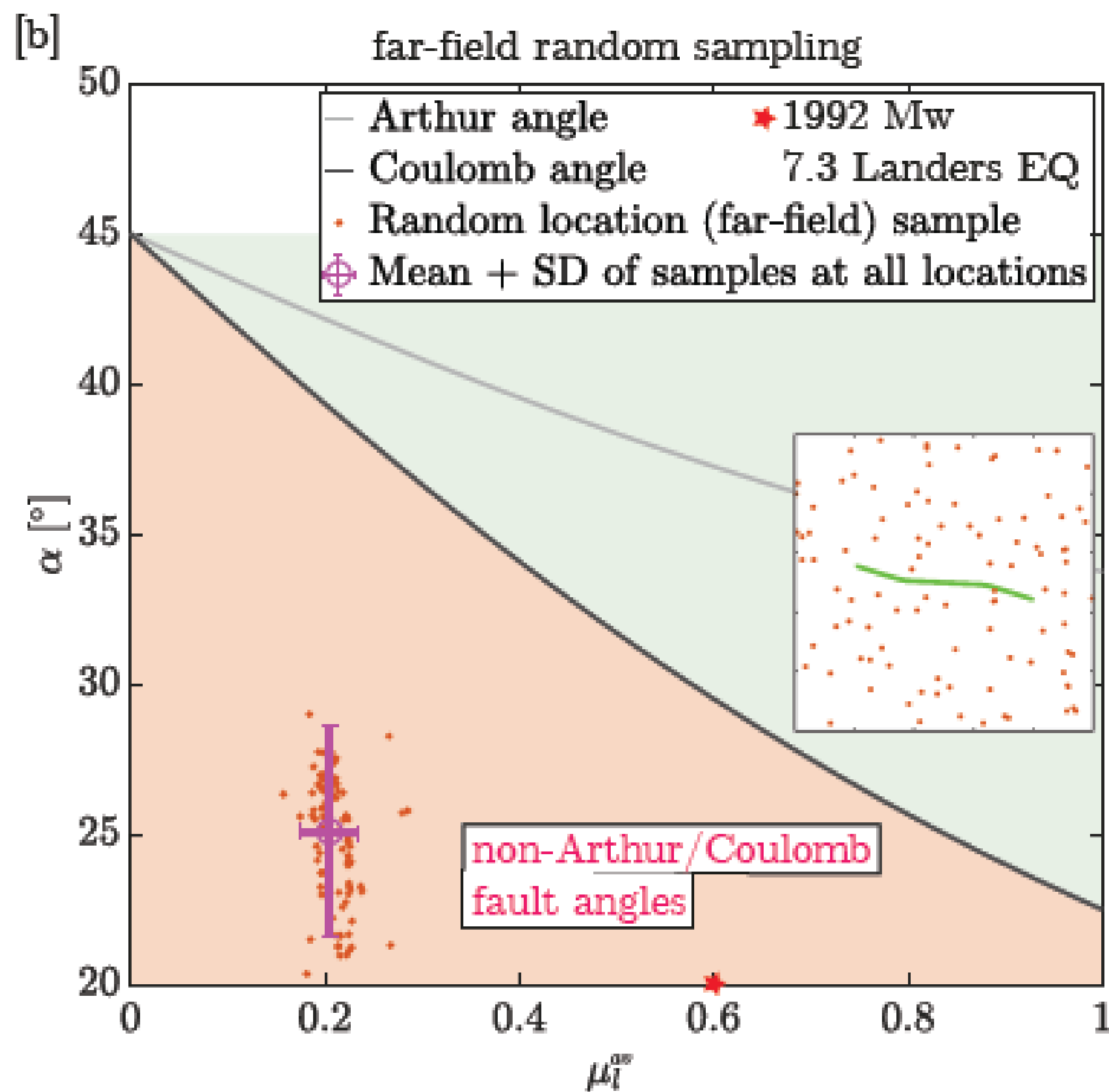
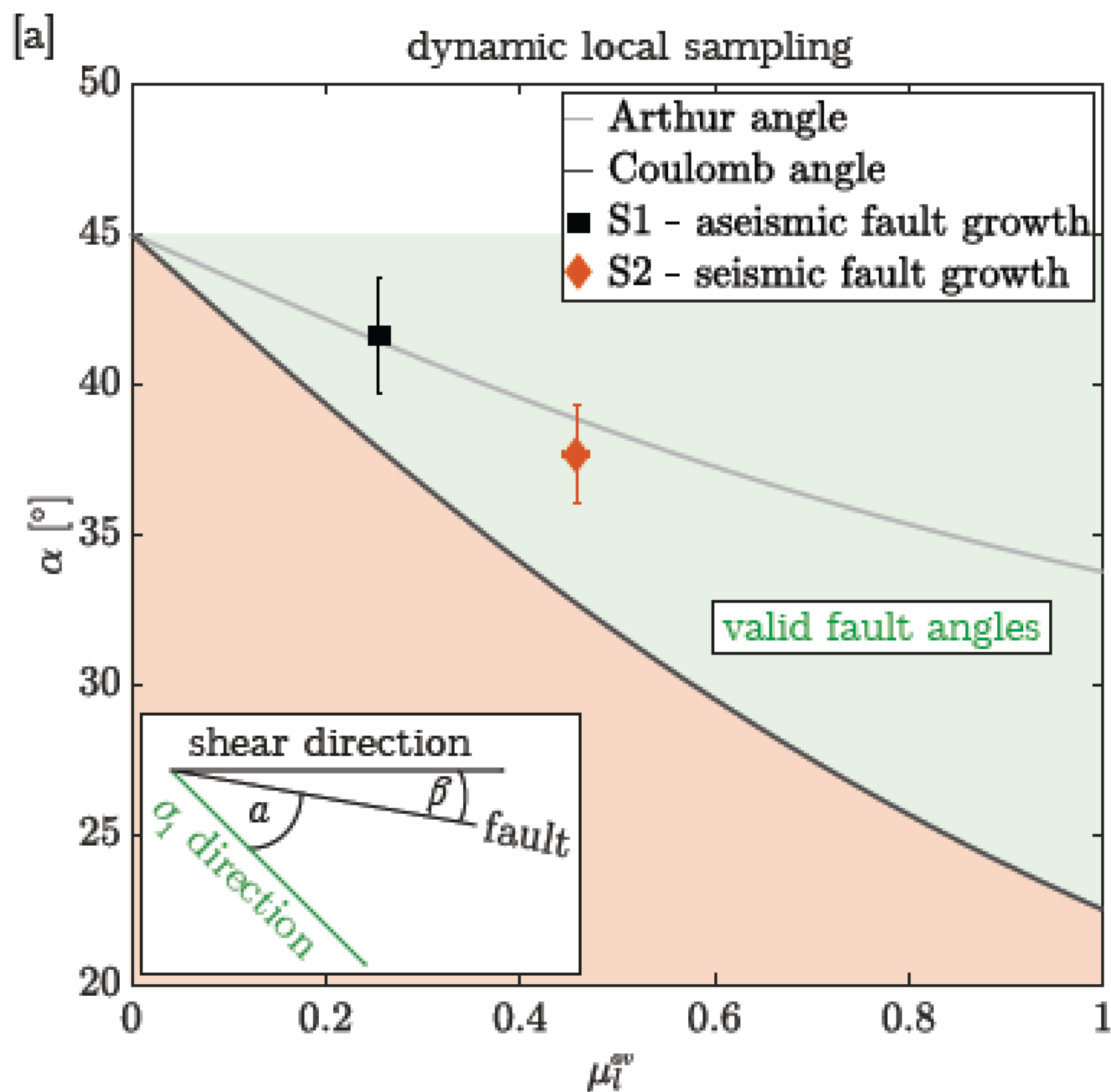
Seismicity
Milliseconds
Millimeters

Two fault growth modes exist: seismic and aseismic

Time = 234.851 a



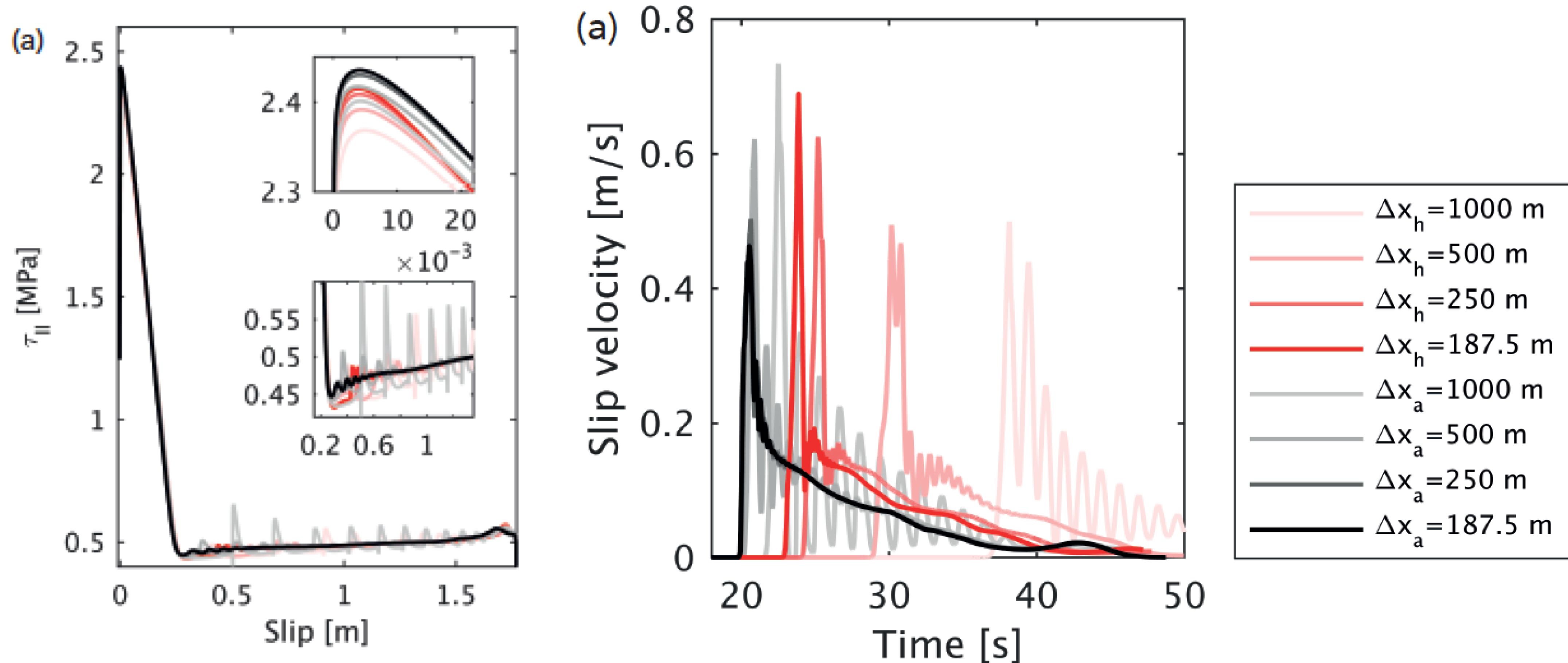
» Mis-orientation may indicate seismic fault growth



Do these angles and characteristics depend on grid size?

- Plasticity is grid size dependent (e.g., Vermeer and de Borst, 1984)
- Length-scale in slip rate formulation helps with grid convergence (e.g., Needleman, 1988)

$$V = \underbrace{2\dot{\epsilon}'_{II(p)}}_{\text{plastic strain rate}} \underbrace{\Delta x}_{\text{fault width}}$$

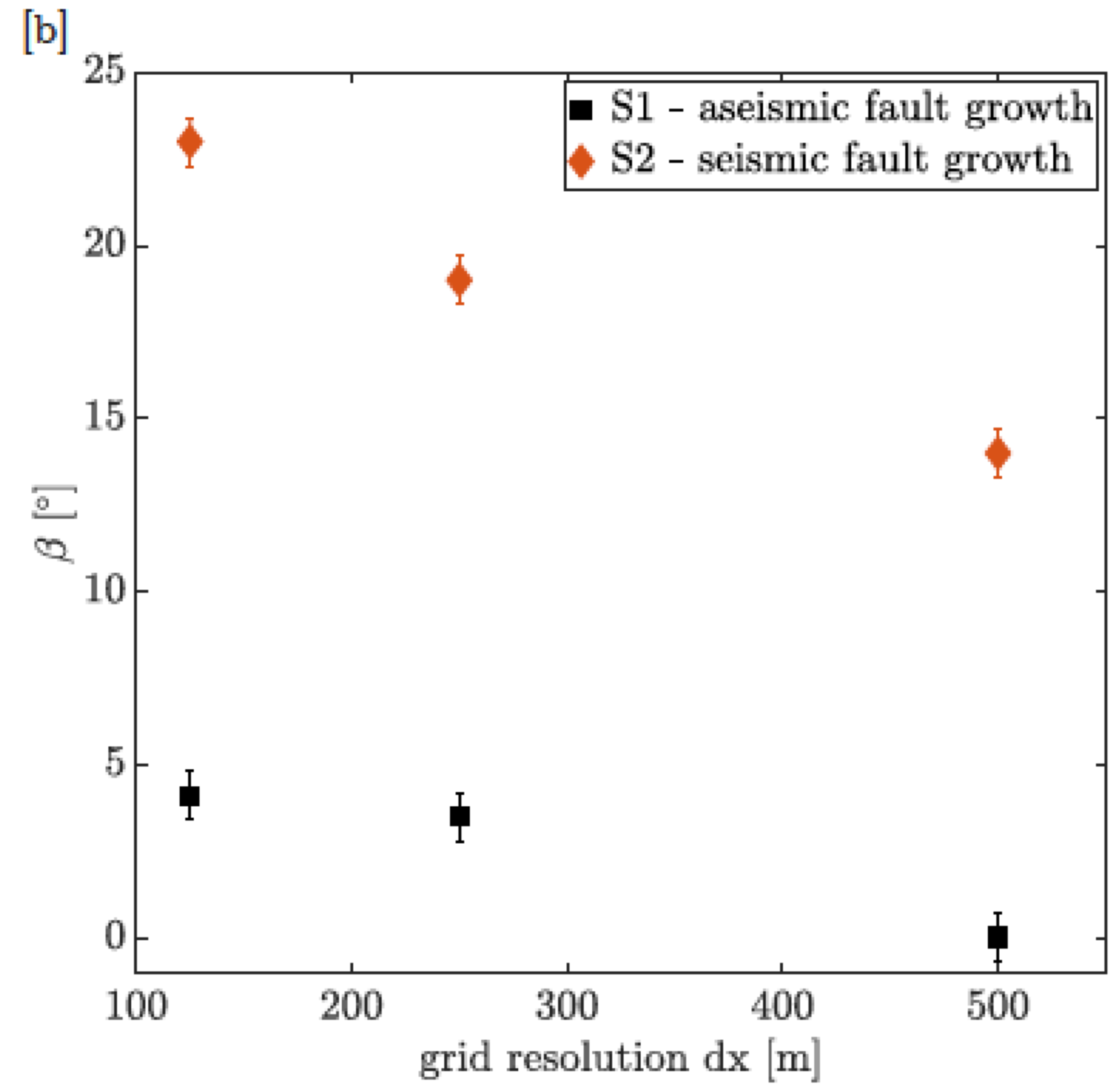
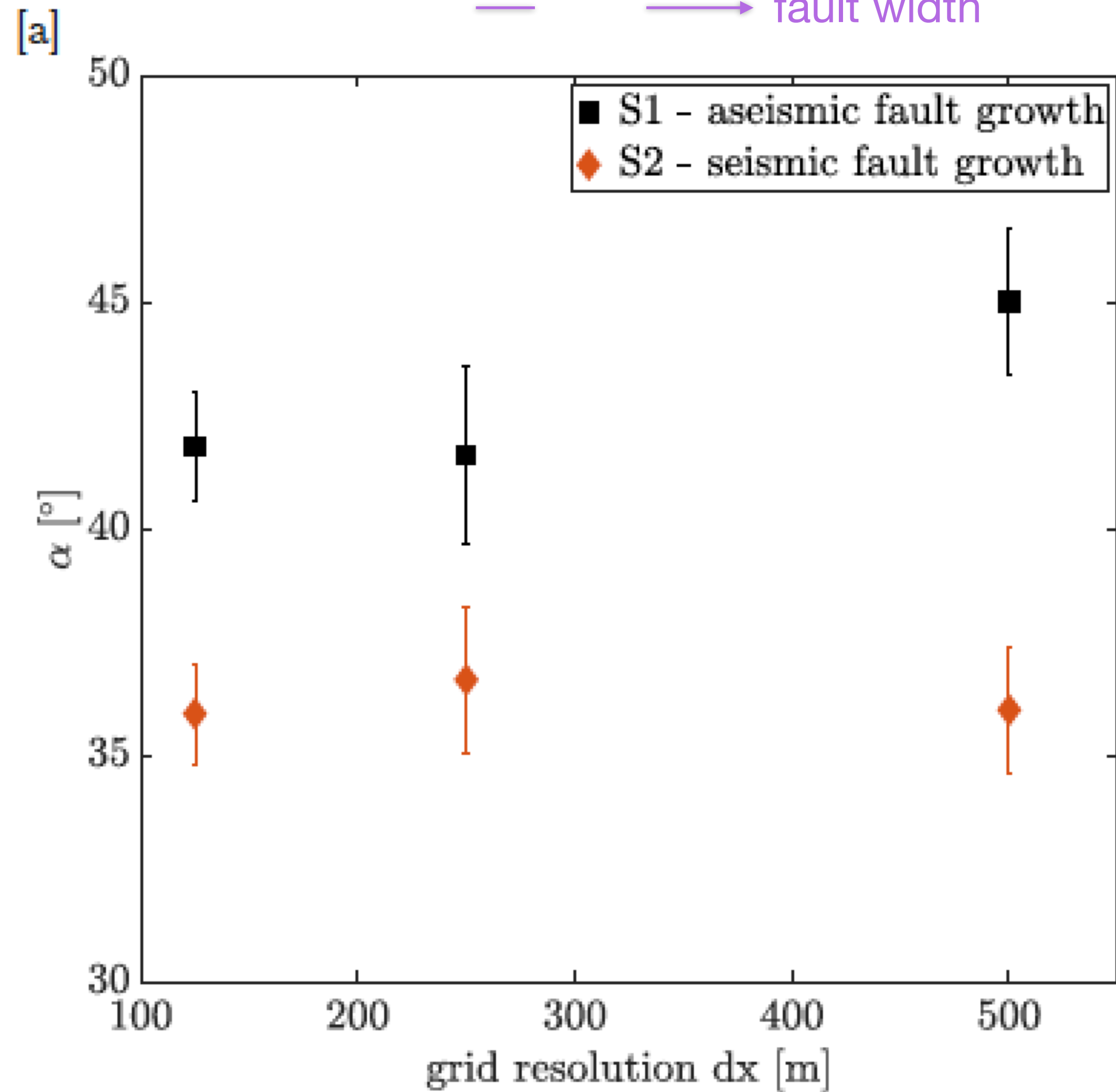


Do these angles depend on grid size?

- Length-scale in slip rate formulation helps with grid convergence (e.g., Needleman, 1988), **but not enough for evolving fault**

$$V = 2\dot{\varepsilon}'_{II(p)} \Delta x$$

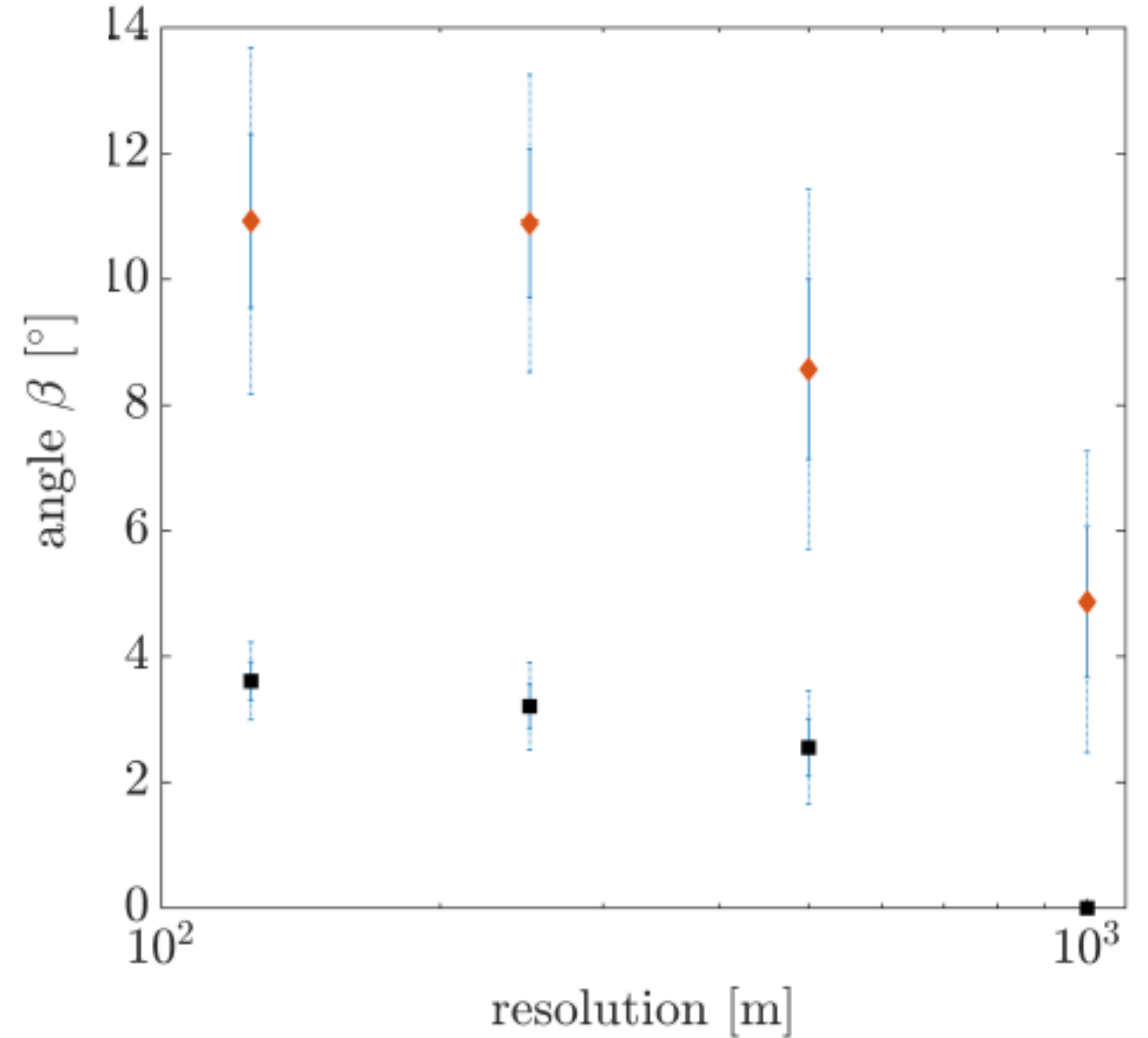
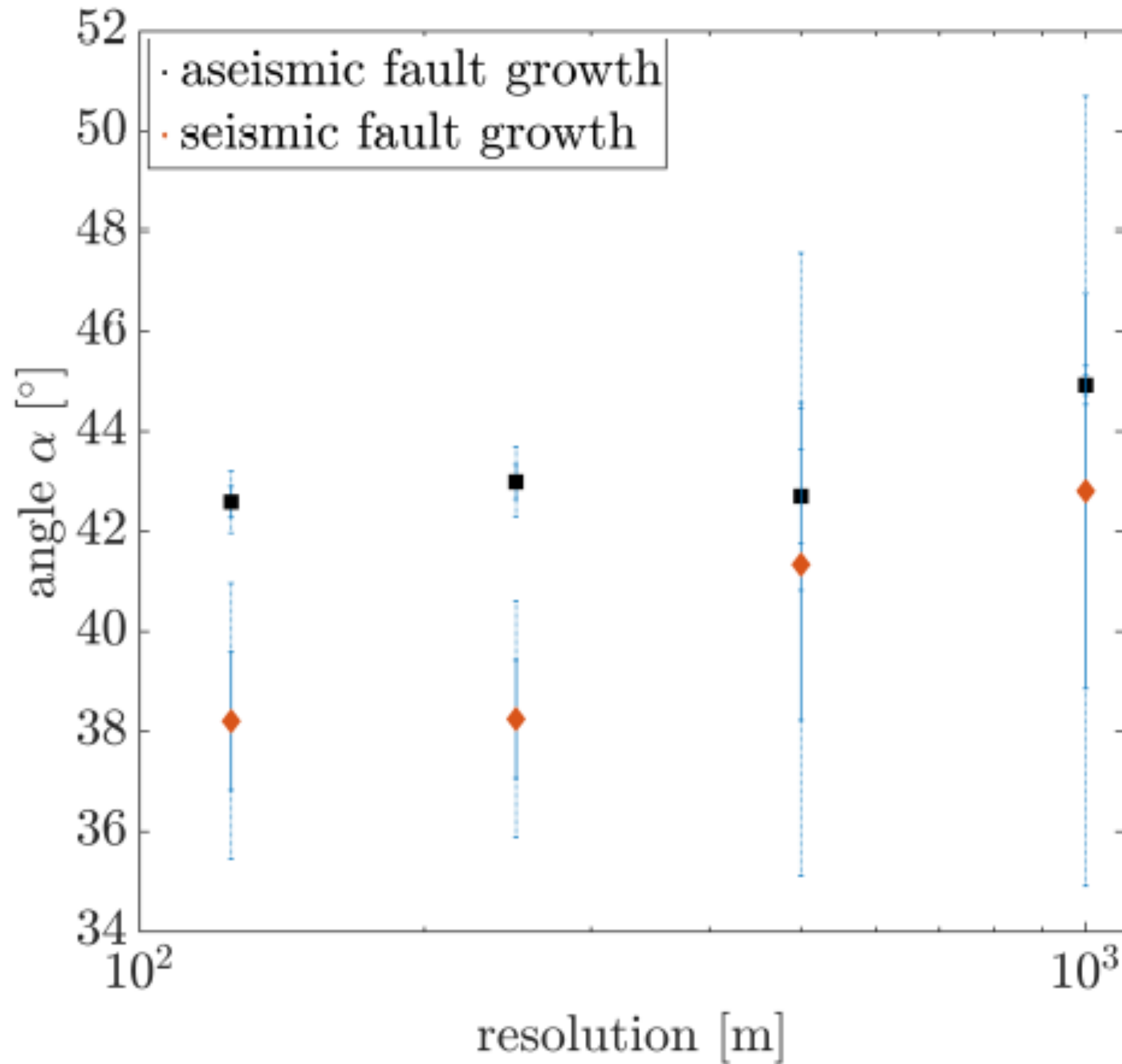
— Δx — \rightarrow fault width



Do these angles depend on grid size?

- What is fault width W if a fault has not yet localized?

$$W_V^2 = W_{max} \log\left(1 + K \frac{V_0}{V_p}\right)$$



- » Illustrate how lithosphere- and mantle dynamics, structure and rheology can influence shallow tectonics and seismicity
- **Does incoming sediment thickness increase maximum earthquake magnitude?**

How sediment thickness influences subduction dynamics and seismicity

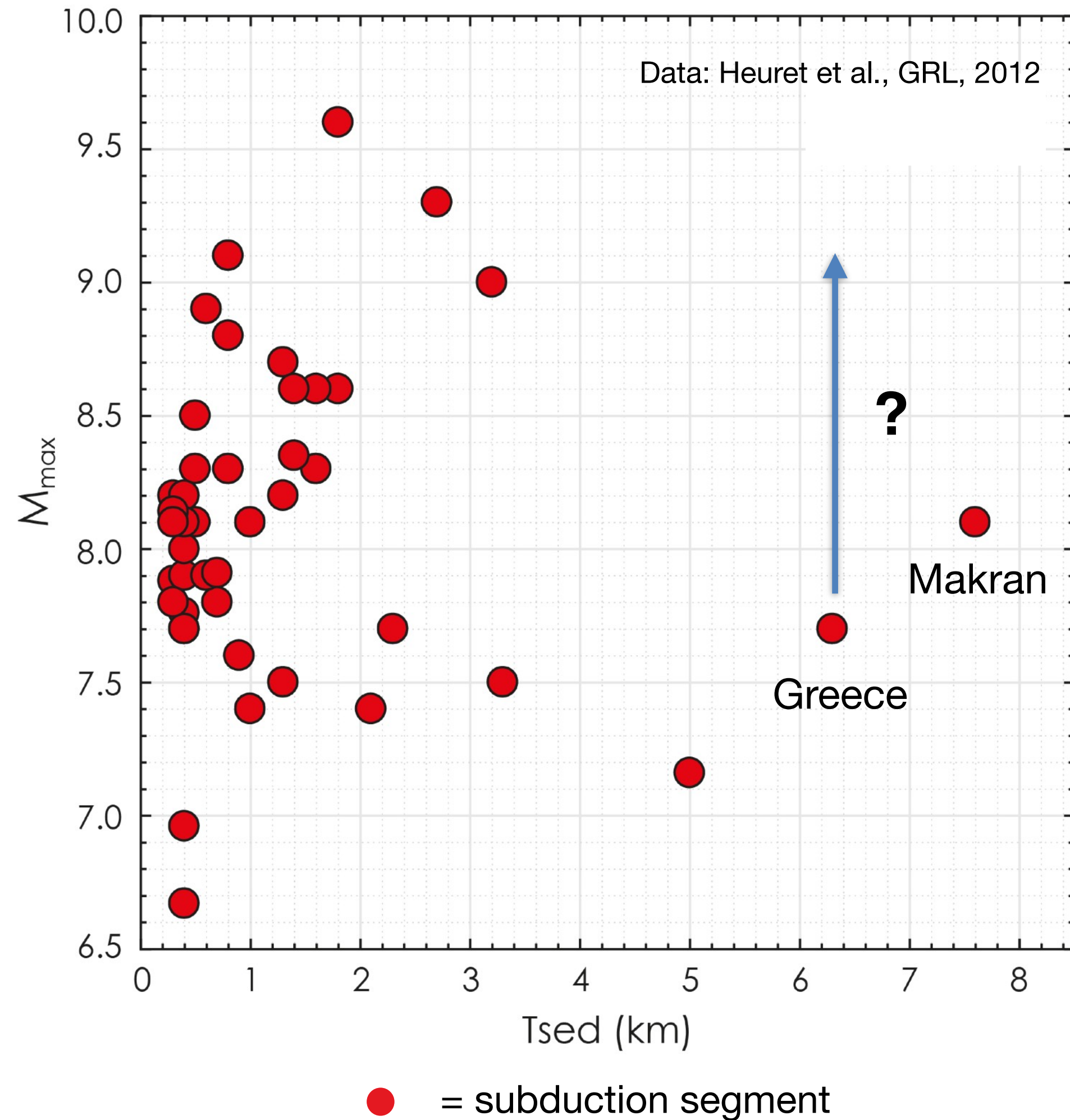
S. Brizzi^{1,2*}, I. van Zelst³, F. Funiciello¹, F. Corbi¹, Y. van Dinther^{3,4}

¹Laboratory of Experimental Tectonics, University of Roma Tre, Italy; ²Natural and Experimental Tectonics research group, University of Parma, Italy; ³Seismology and Wave Physics, ETH Zürich, Switzerland; ⁴Department of Earth Sciences, Utrecht University, Utrecht, the Netherlands

Under review at Earth Planetary Science Letters

Does incoming sediment thickness increase maximum earthquake magnitude?

- Various observations suggest T_{sed} increases M_{max} (e.g., Ruff, 1989; Heuret et al., 2012; Scholl et al., 2015; Seno, 2017; Brizzi et al., 2018)
- But does it? And how?

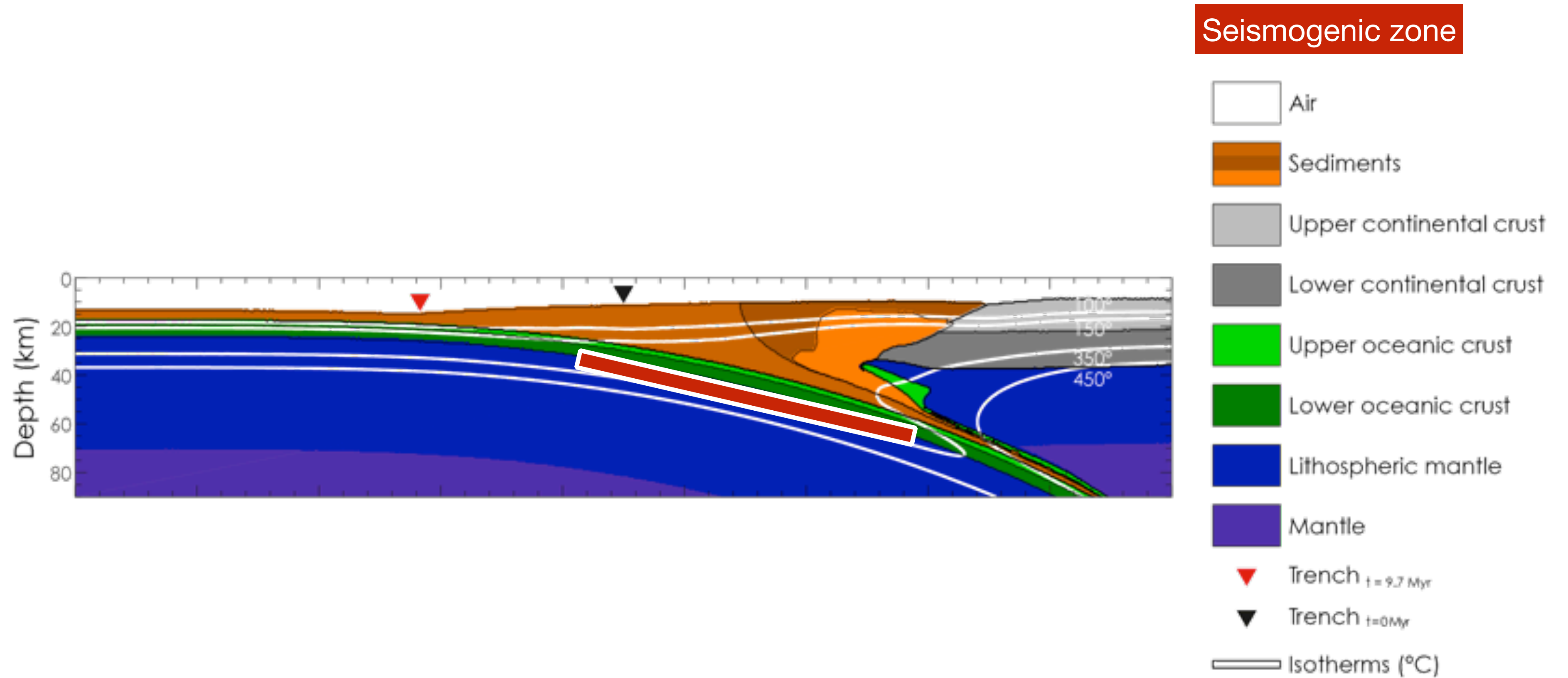


- Cross-scale modeling is needed because
 - Observation window \ll recurrence interval
 - Concurrent influence of multi-parameters

Sediment thickness controls geometry of convergent margin

Sediment thickness:

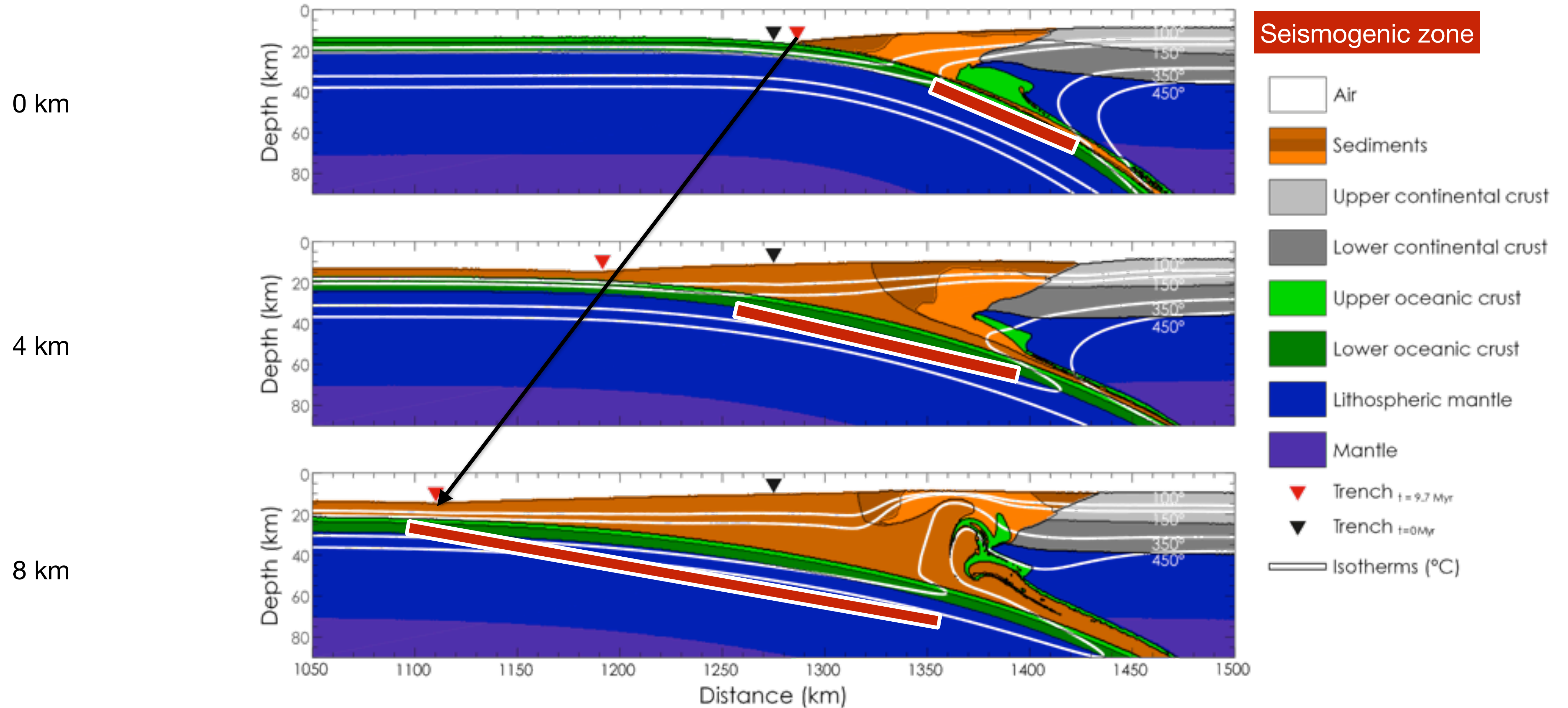
4 km



Sediment thickness controls geometry of convergent margin

- More sediments → seaward growth wedge → trench retreat and unbending → shallower dip → wider seismogenic zone

Sediment thickness:

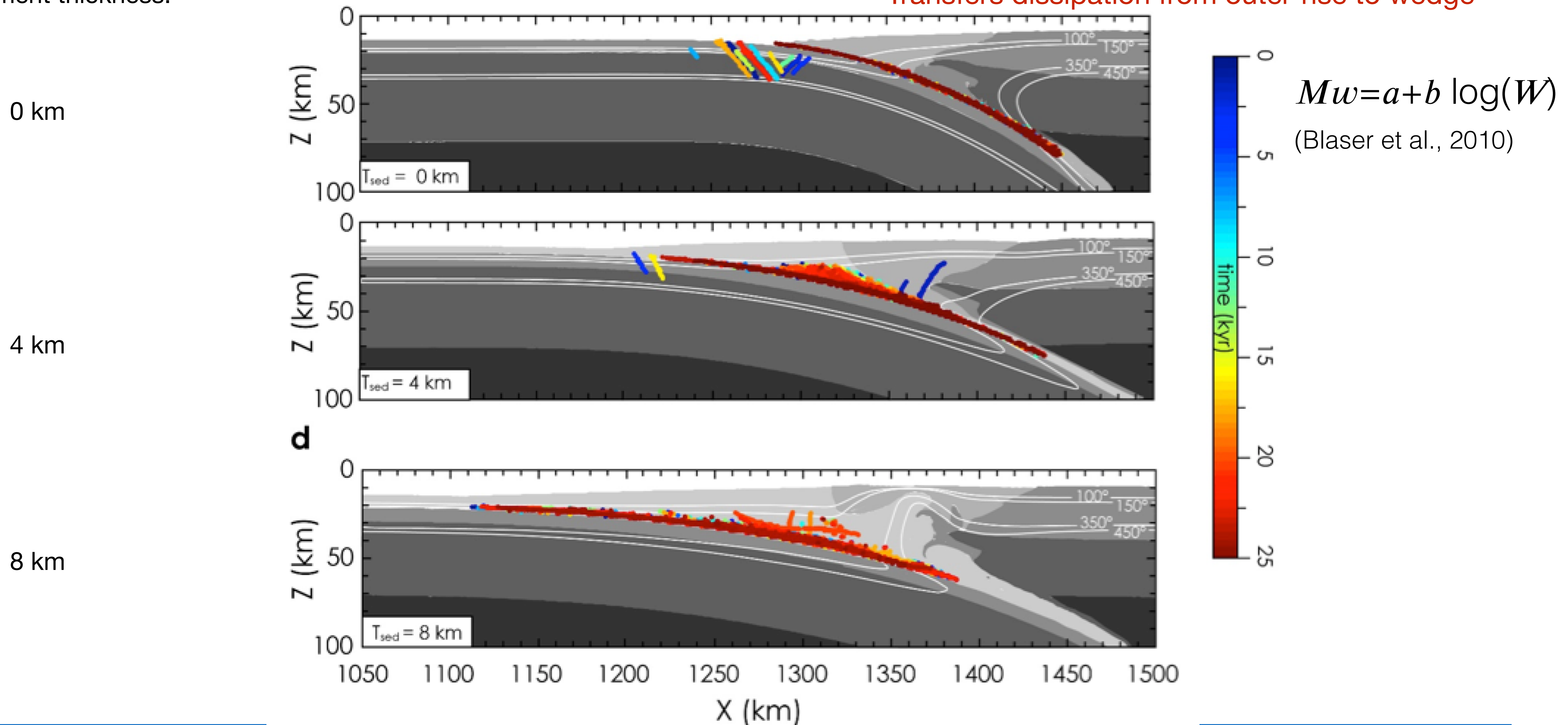


Sediment thickness controls Mmax and type of seismicity

- More sediments → less mechanical coupling → more slab retreat → shallower dip → wider seismogenic zone
 → Mmax up

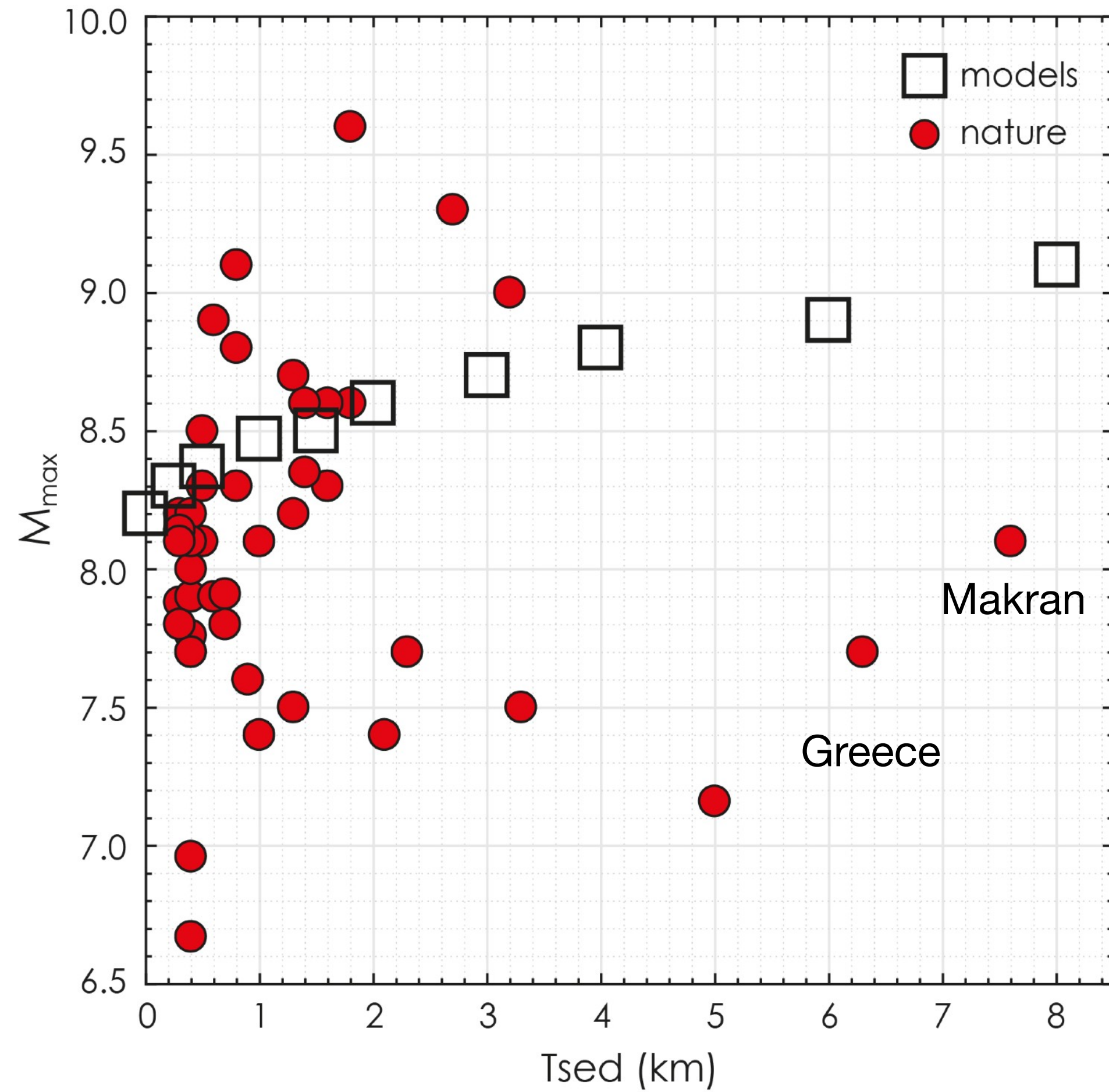
Sediment thickness:

Transfers dissipation from outer-rise to wedge



Sediment thickness controls maximum magnitude

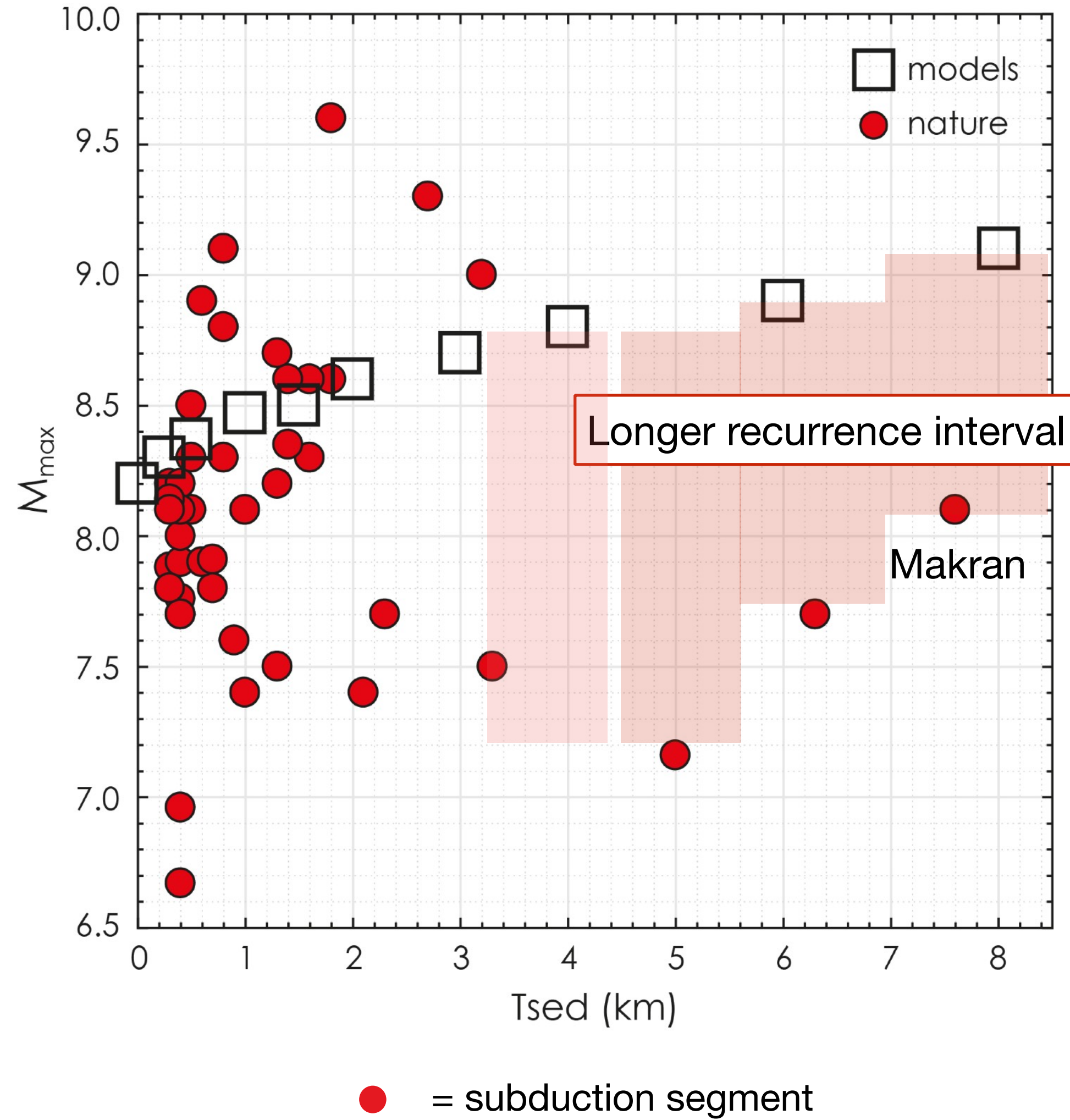
- Models - clarify and quantify suspected trend in nature



● = subduction segment

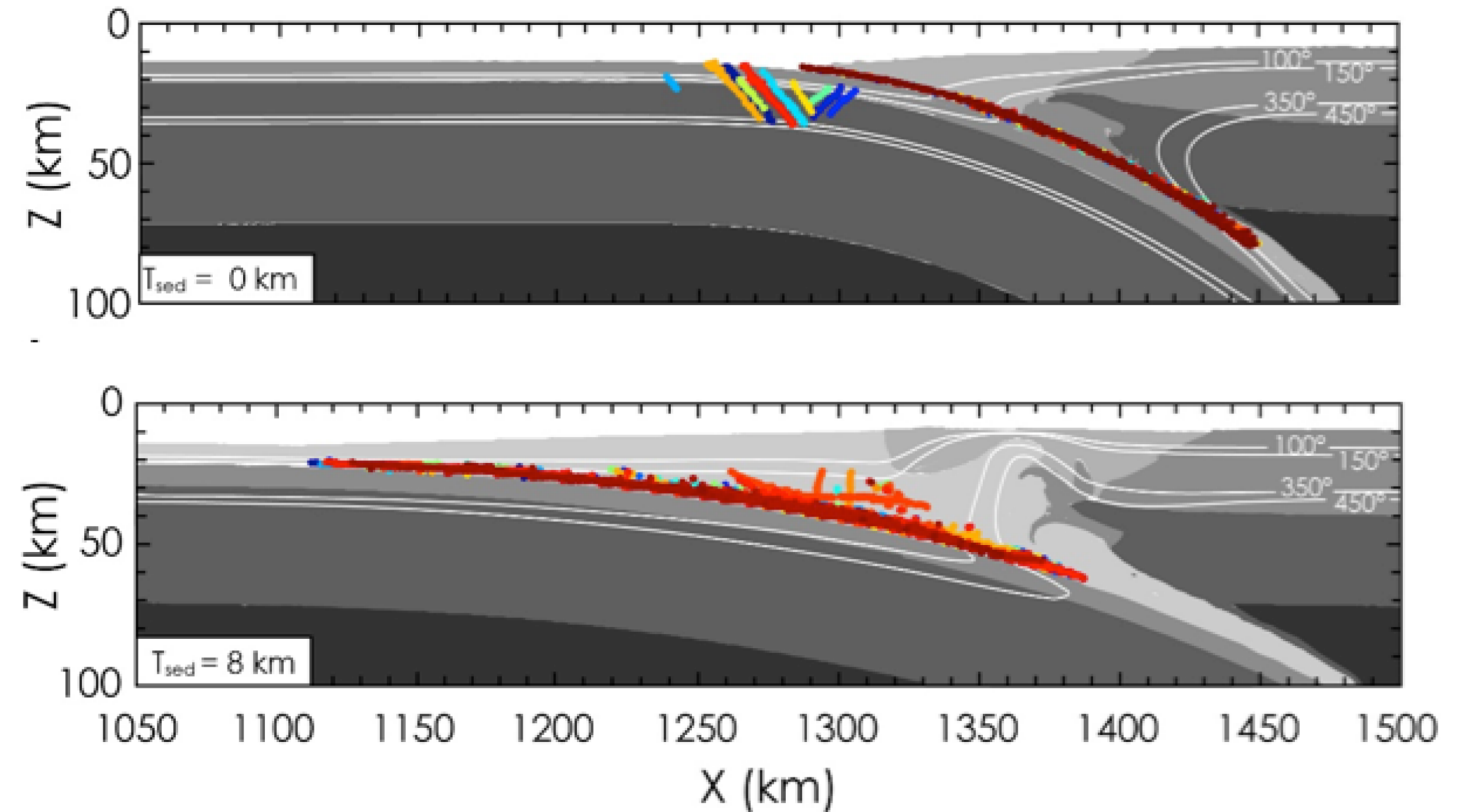
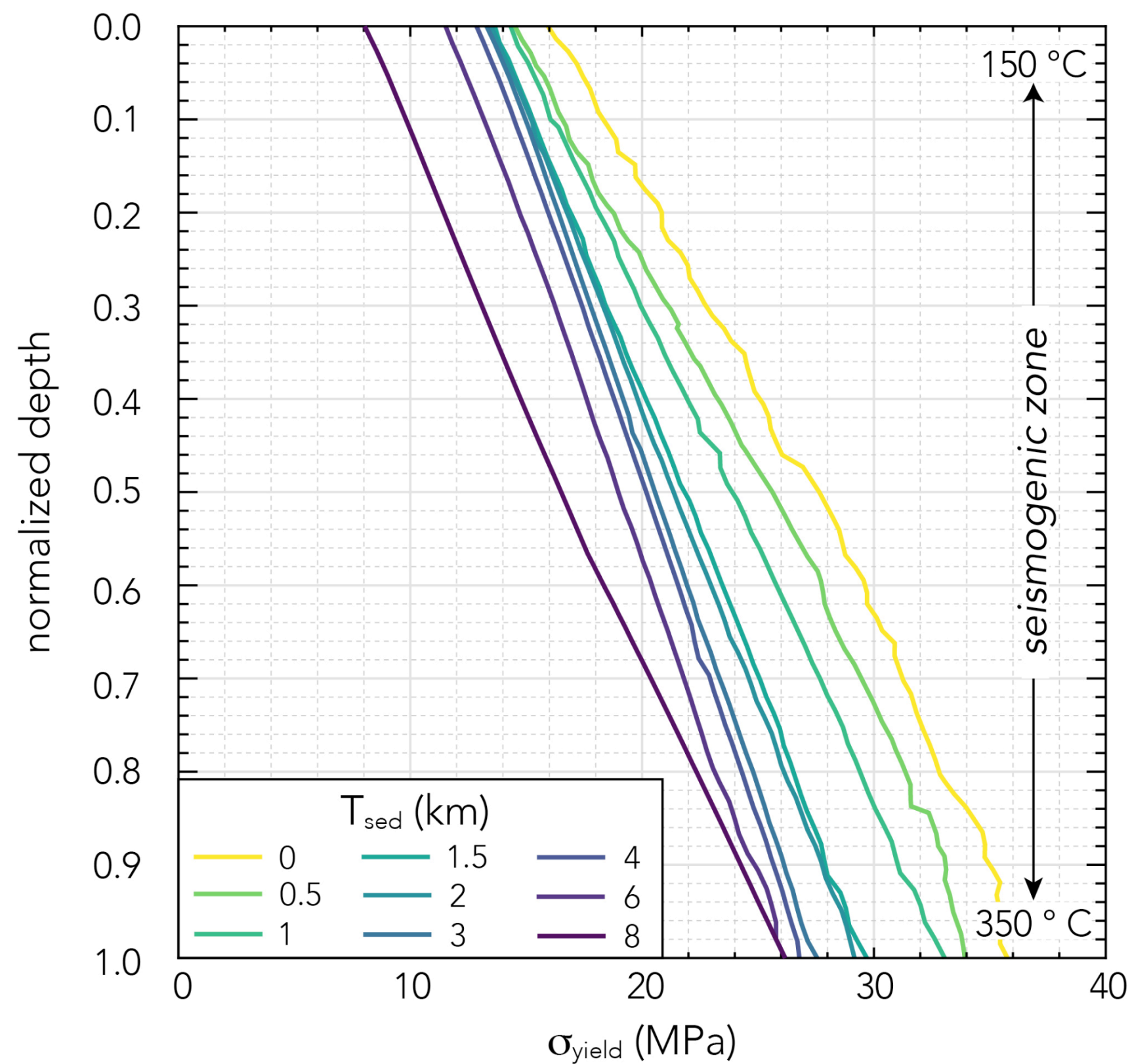
Sediment thickness controls maximum magnitude

- Models
 - clarify and quantify suspected trend in nature
 - provide an explanation for why we might not have yet seen such large magnitudes



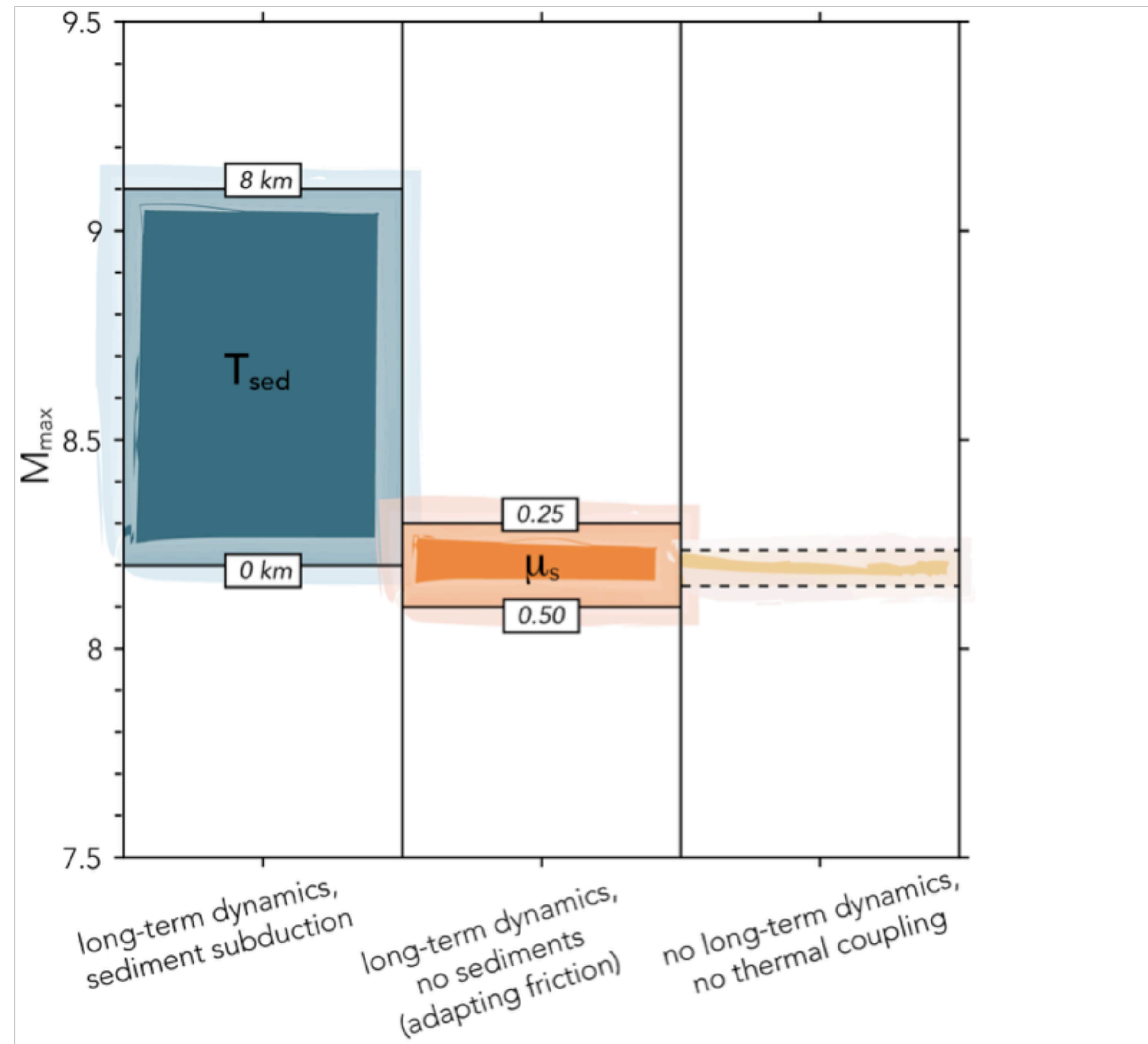
Why does megathrust interface strength decrease with T_{sed} ?

- Intuitive explanation for slab dip decrease through slab retreat as interface is increasingly weaker is not cause
- What makes interface weaker, if “all” models have weak sediments?
 - Warmer incoming sediments → seismogenic zone depth ↓ & a lighter forearc structure → pressure in seismogenic zone ↓



Why model long-term dynamics?

- Simulating long-term dynamics and sediment presence significantly changes quantification



- » Illustrate how lithosphere- and mantle dynamics, structure and rheology can influence shallow tectonics and seismicity
- Does incoming sediment thickness increase maximum earthquake magnitude?
- **Do lower crustal and mantle depth temperature and rheology affect seismicity and tectonics?**

Tectonics and seismicity in the Northern Apennines driven by slab retreat and crustal delamination

Mario D'Acquisto^{1,2}, Luca Dal Zilio^{1,3}, Irene Molinari^{1,4}, Edi Kissling¹, Taras Gerya¹,
Ylona van Dinther^{1,2}

¹*Department of Earth Sciences, ETH Zürich, Sonneggstrasse 5, Zürich, Switzerland.*

²*Department of Earth Sciences, Utrecht University, Princetonlaan 8A/4, Utrecht, Netherlands*

³*Division of Geological and Planetary Science, California Institute of Technology, 1200 E California*

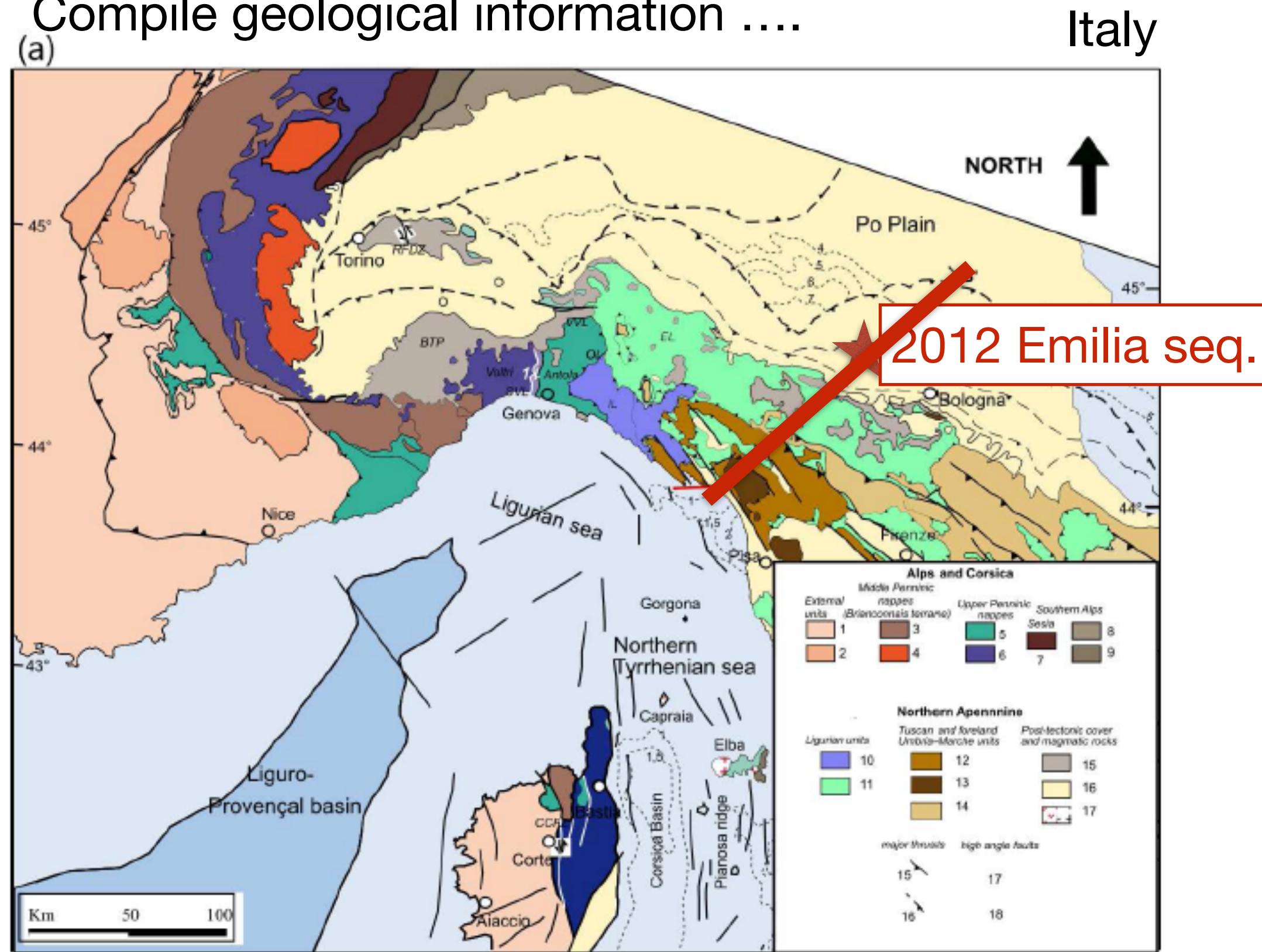
⁴*Sezione di Bologna, Istituto Nazionale di Geofisica e Vulcanologia, via Donato Creti 12, Bologna, Italy*

Submitted to Journal of Geophysical Research - Solid Earth

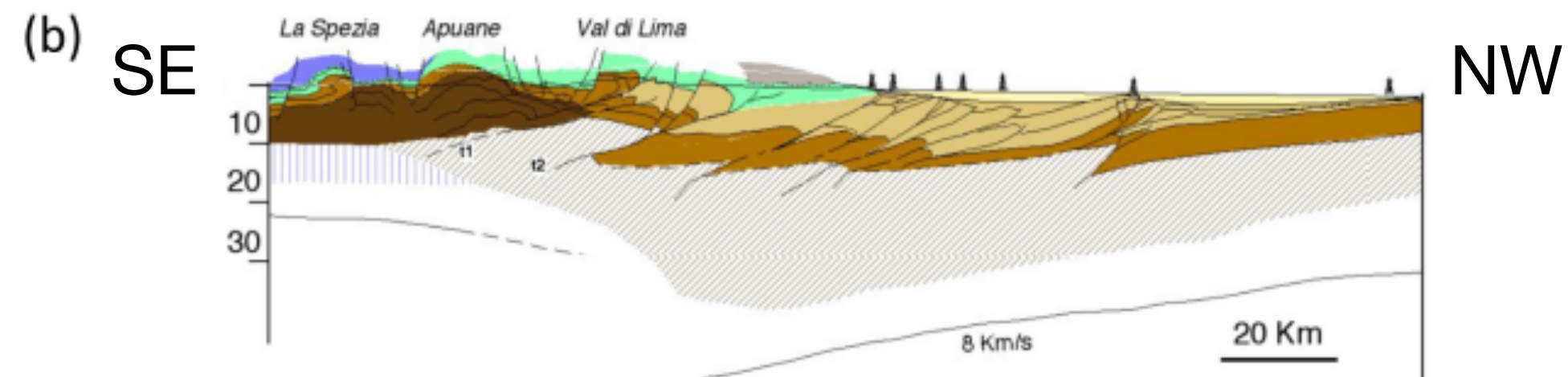
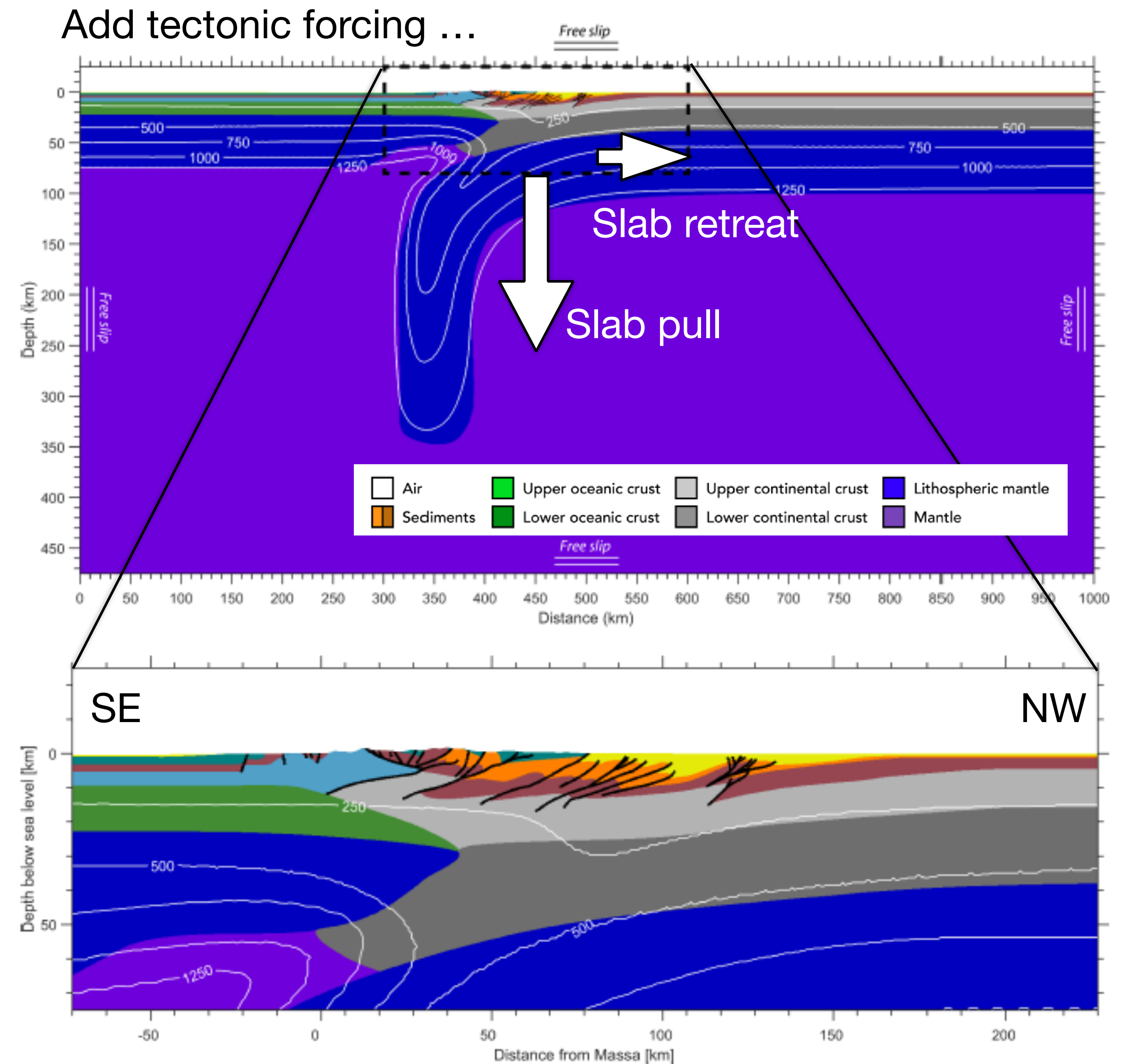
Building an instantaneous model

- Combine available information to build regional model of Northern Apennines

Compile geological information

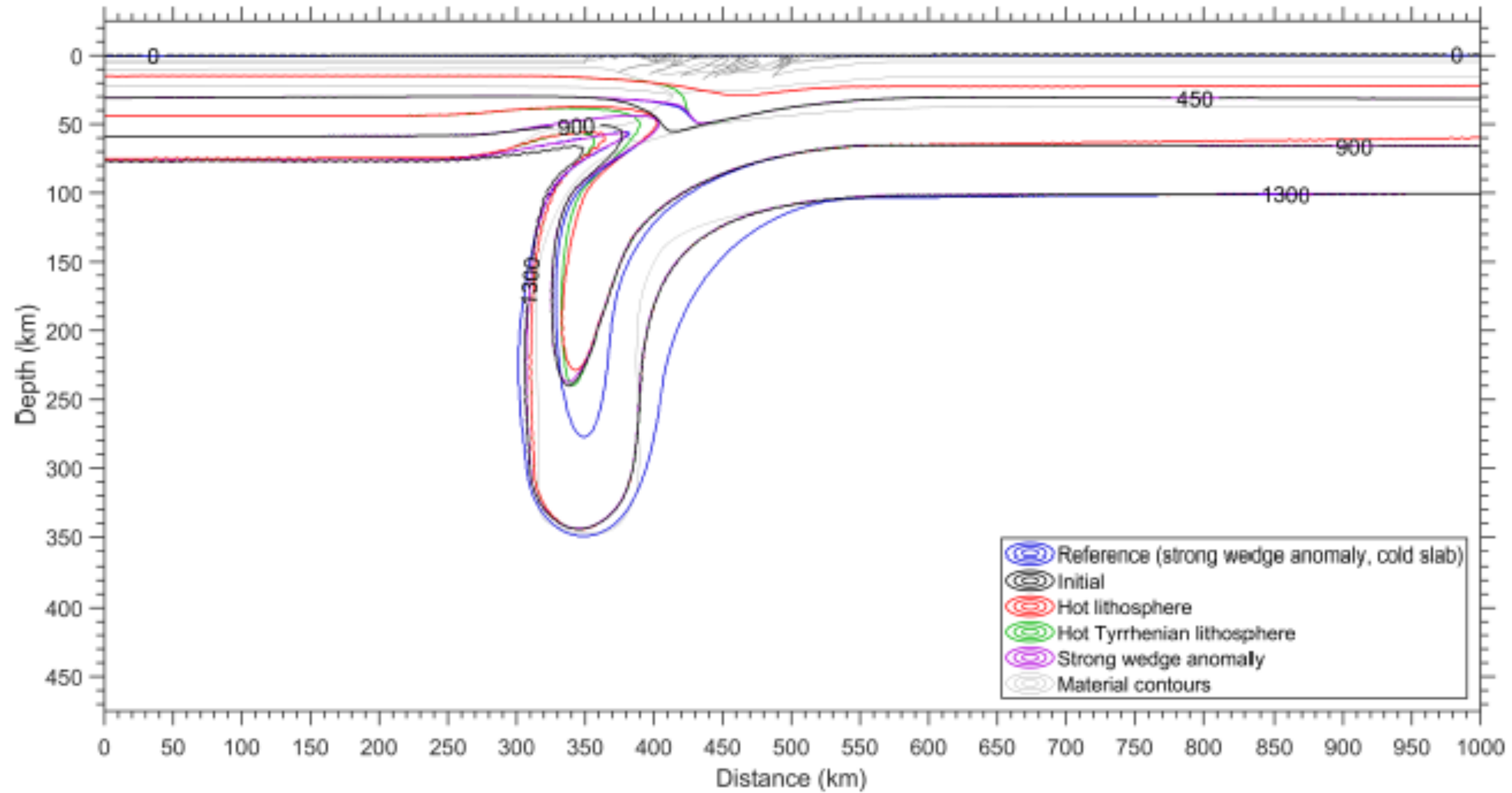


Add tectonic forcing ...



Add thermal model

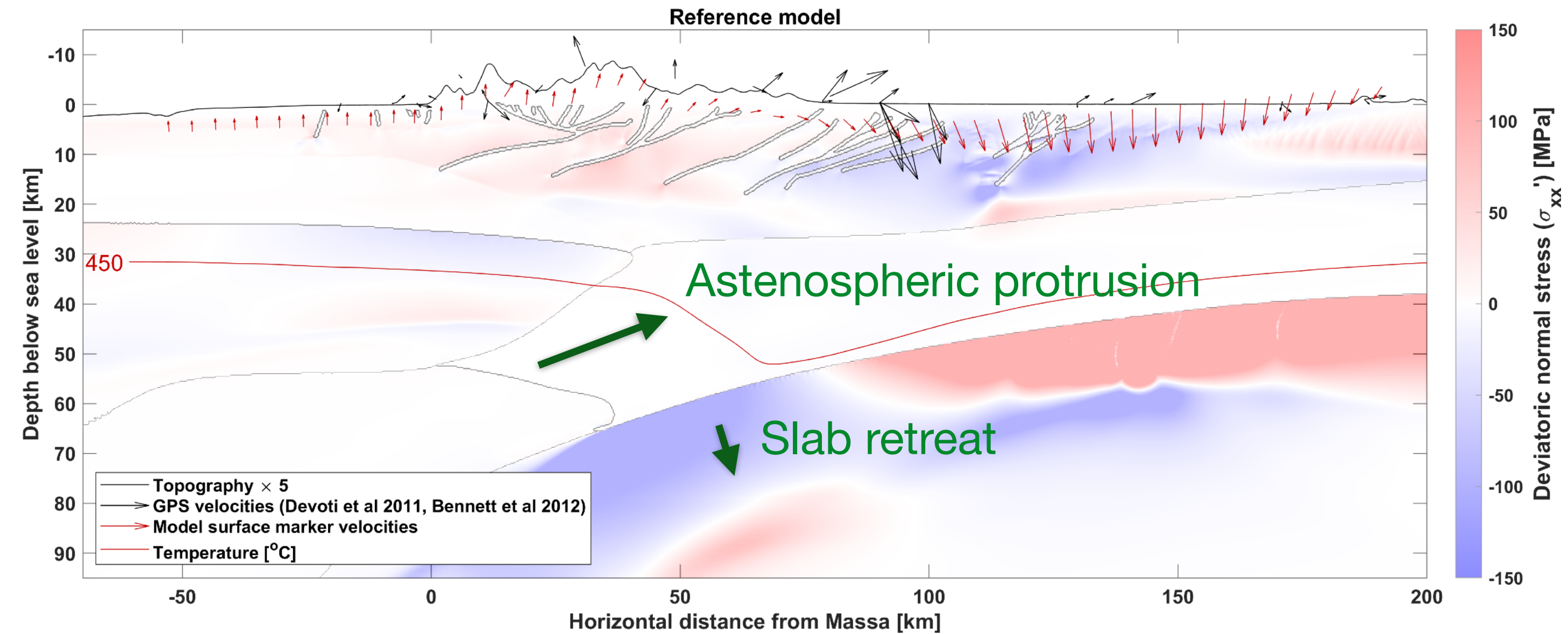
- Using generic slab retreating simulations



Long-term dynamics needs buoyant and highly ductile material beneath suture zone

- Run to sensible long-term deformation regimes
- To agree with stress regime and surface displacements need → → →
 - weak lower crust
 - high mantle wedge temperatures

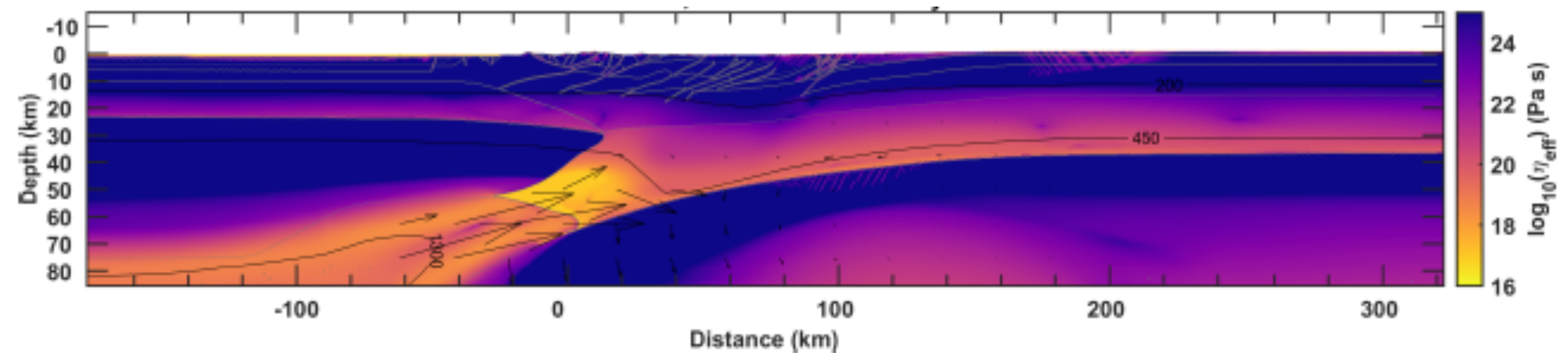
World Stress Map: ← extension ← compression →



Extension and uplift
in Apennine range

Compression and subsidence
in Po basin

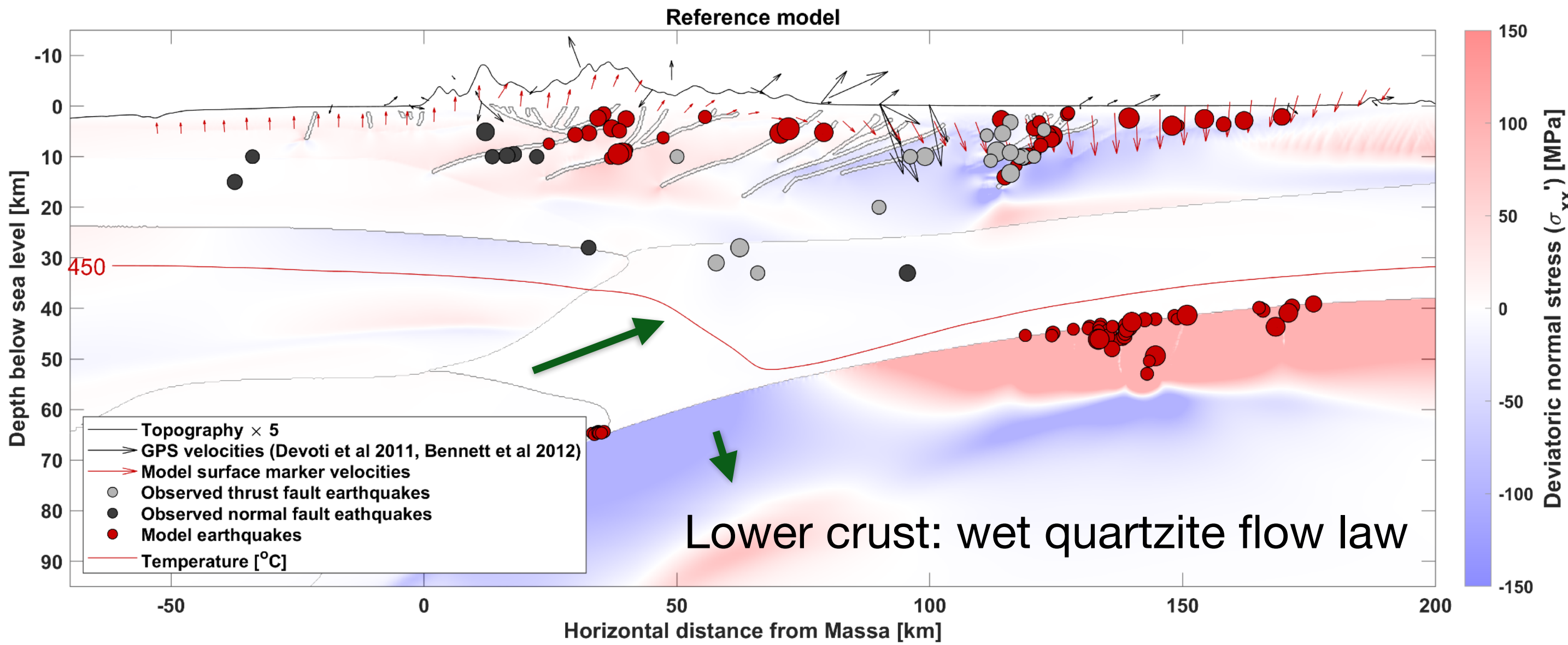
Viscosity distribution



Seismicity broadly agrees with data

- Switch to 1 year time steps

World Stress Map:

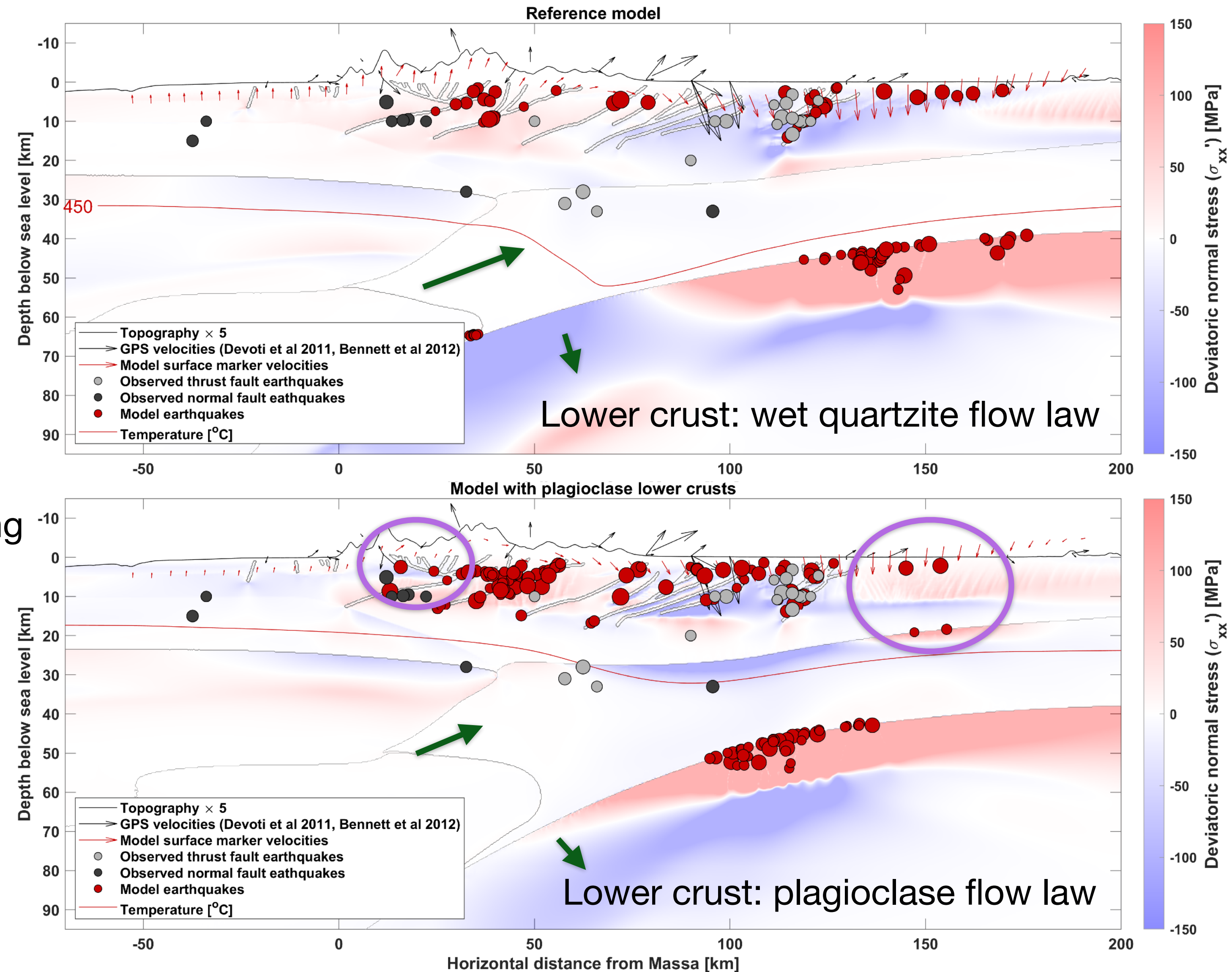


Normal faulting events
in Apennine range

Thrust faulting events
in Po basin

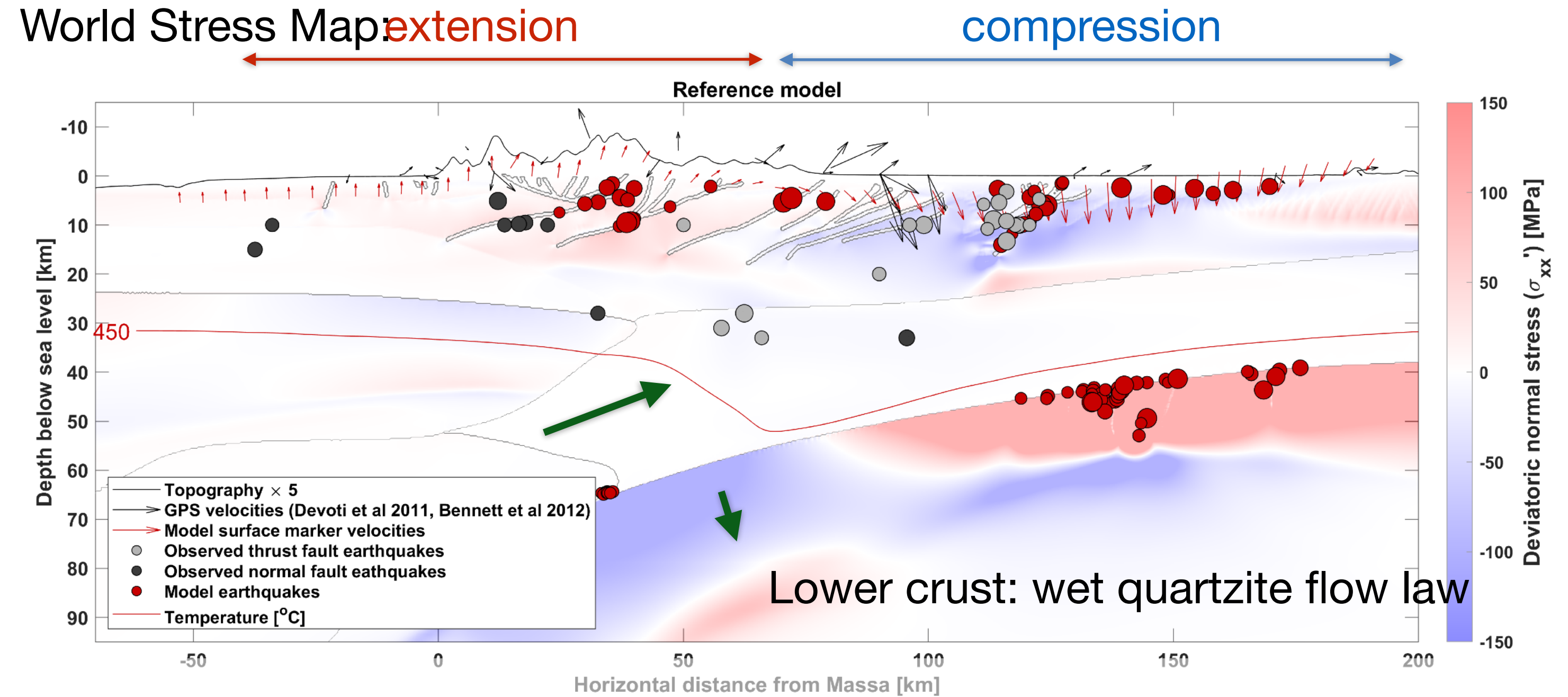
Lower crustal rheology affects stresses and seismicity

World Stress Map: ← extension → compression



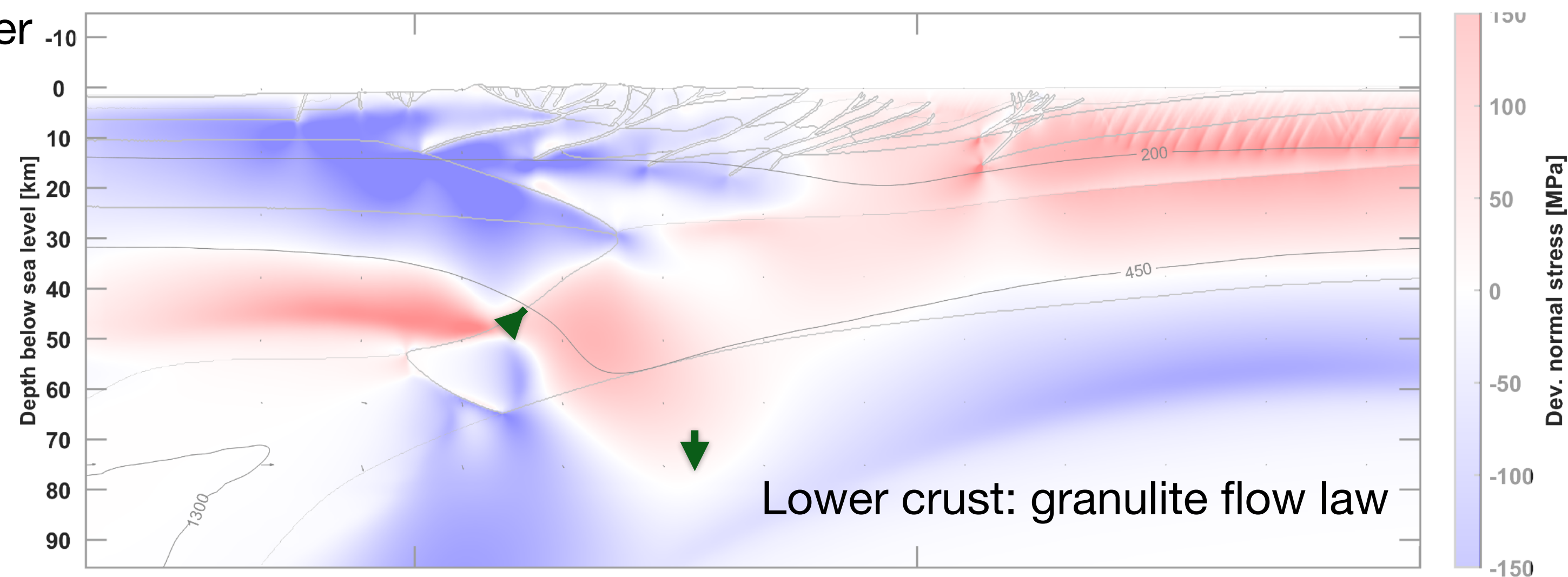
- Plagioclase flow law for lower crust is too strong
 - » Mismatch stress and earthquake type

Lower crustal rheology affects stresses and seismicity



- Granulite flow law for lower crust is even stronger
 - » Complete mismatch stress regime

» Need protrusion, delamination, retreat

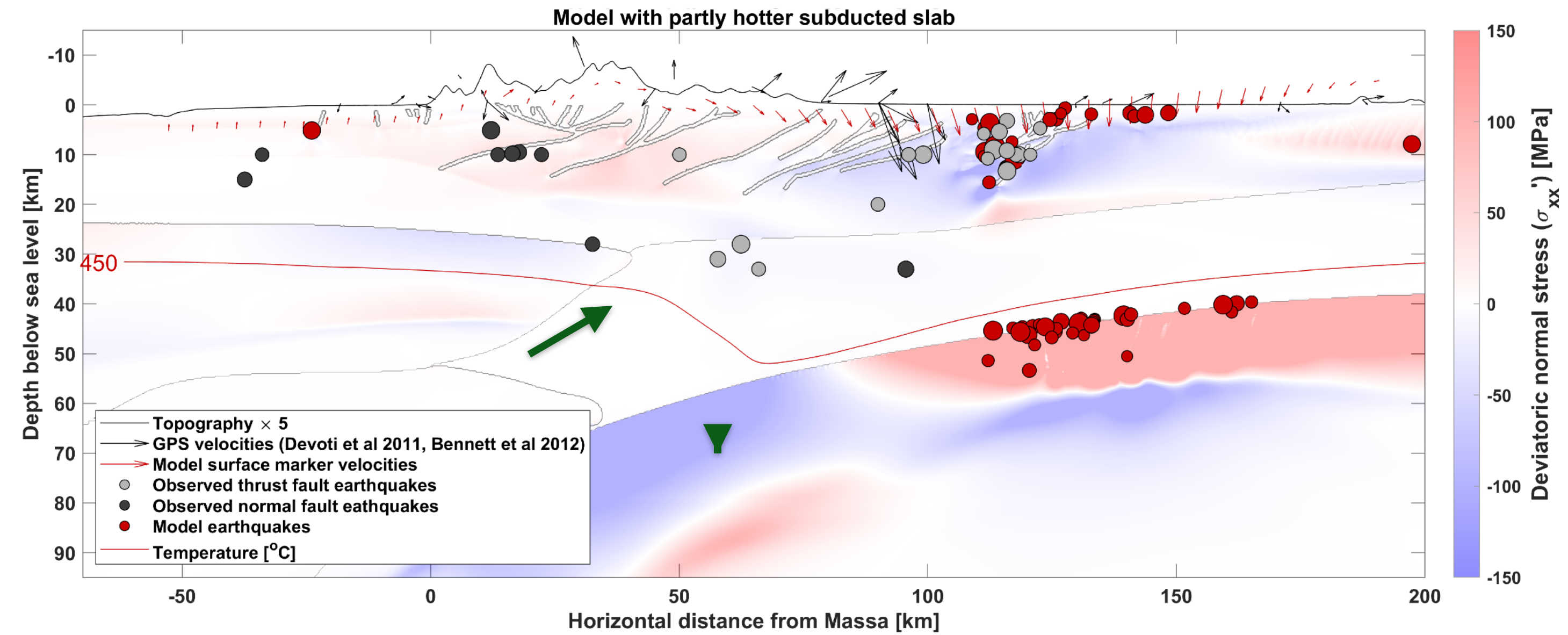
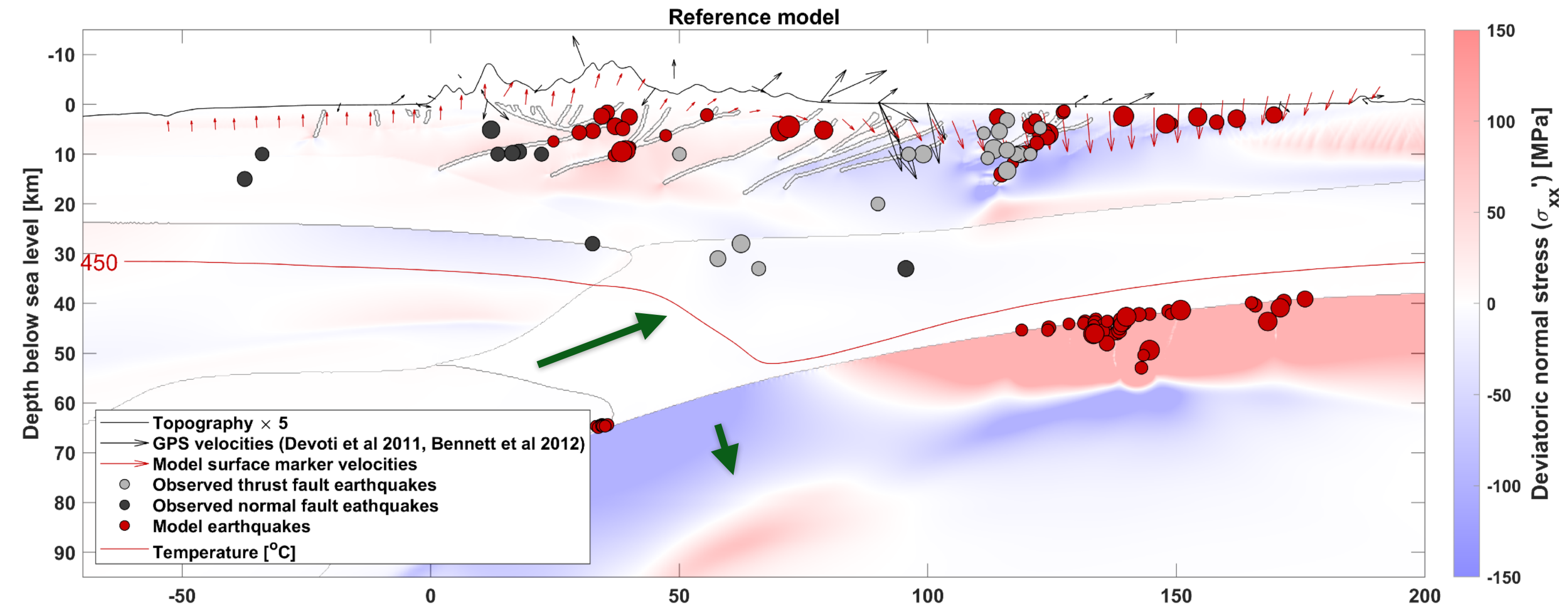
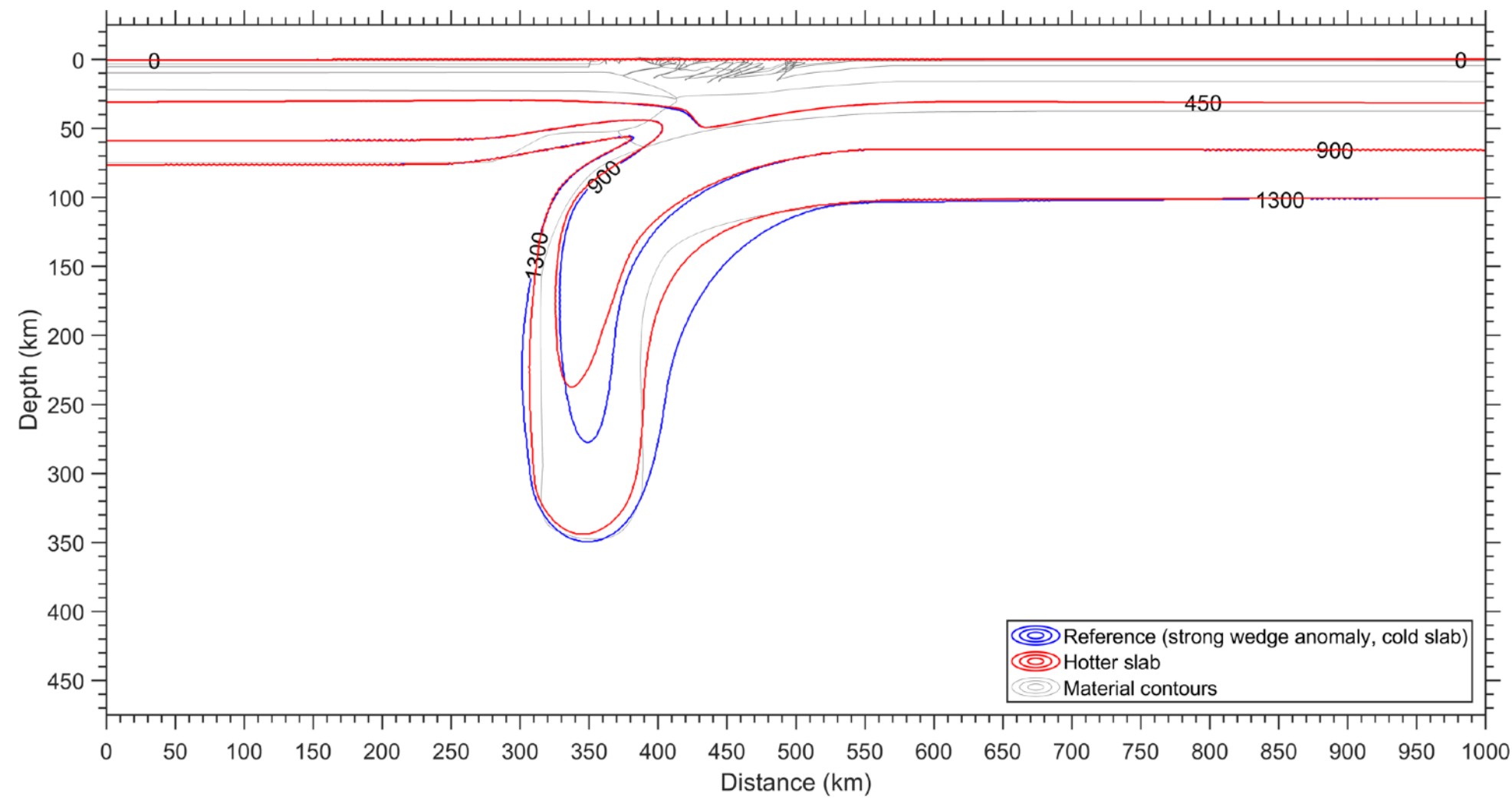


Large-scale dynamics affects seismicity

World Stress Map:

extension

compression



- Reduce slab pull
 - » Shuts down extensional seismicity in range

» Should care about depths larger than ~15km

- Seismicity can be another observable to constrain rheologies and structure

Presentation objectives

- » Illustrate how lithosphere- and mantle dynamics, structure and rheology can influence shallow tectonics and seismicity
- Does incoming sediment thickness increase maximum earthquake magnitude?
- Do lower crustal and mantle depth temperature and rheology affect seismicity and tectonics?
- **Does the mantle affect surface displacements at time scales of minutes to days?**

Pure Appl. Geophys.


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<https://doi.org/10.1007/s00024-019-02250-z>

Pure and Applied Geophysics

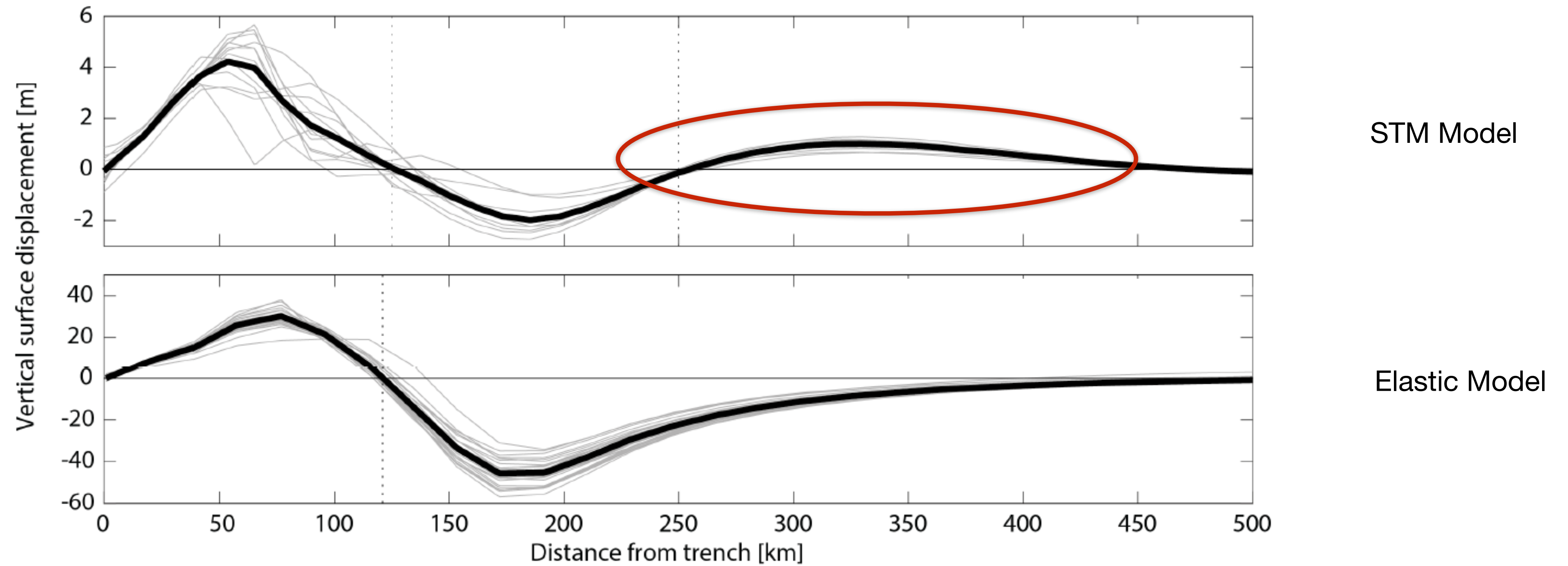


A Secondary Zone of Uplift Due to Megathrust Earthquakes

YLONA VAN DINTHER,^{1,2}  LUKAS E. PREISWERK,^{1,3,4} and TARAS V. GERYA⁴

A secondary zone of coseismic uplift

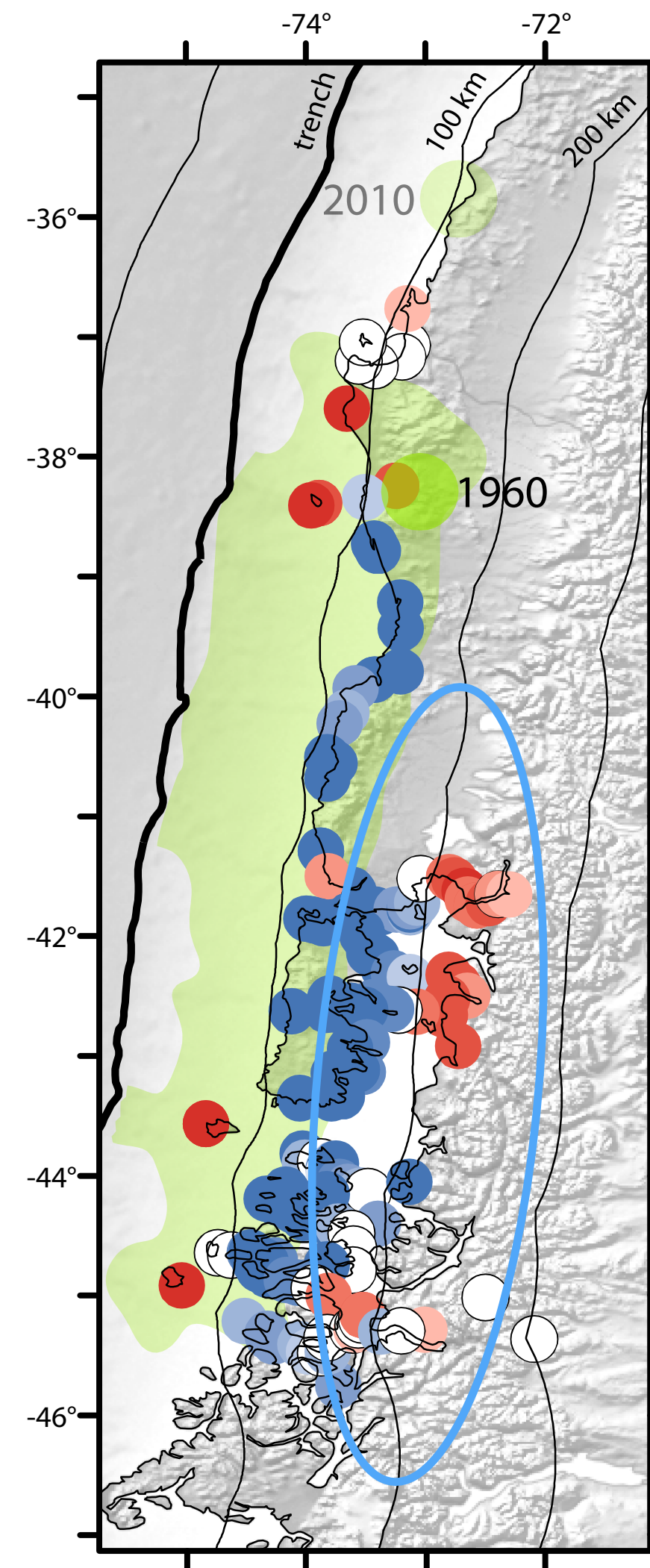
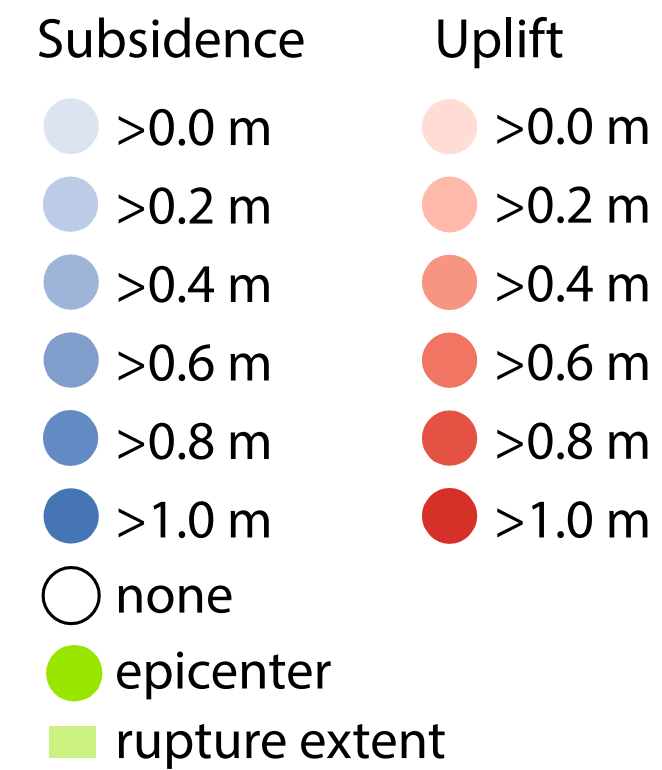
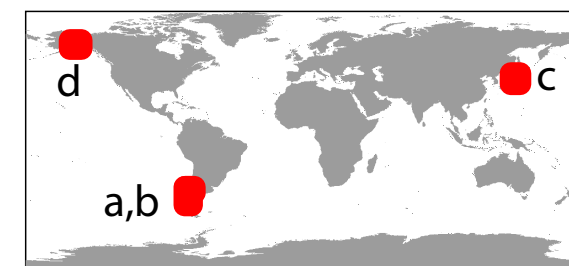
- STM models identified a new physical phenomena: a secondary zone of coseismic uplift



A secondary zone of coseismic uplift exists

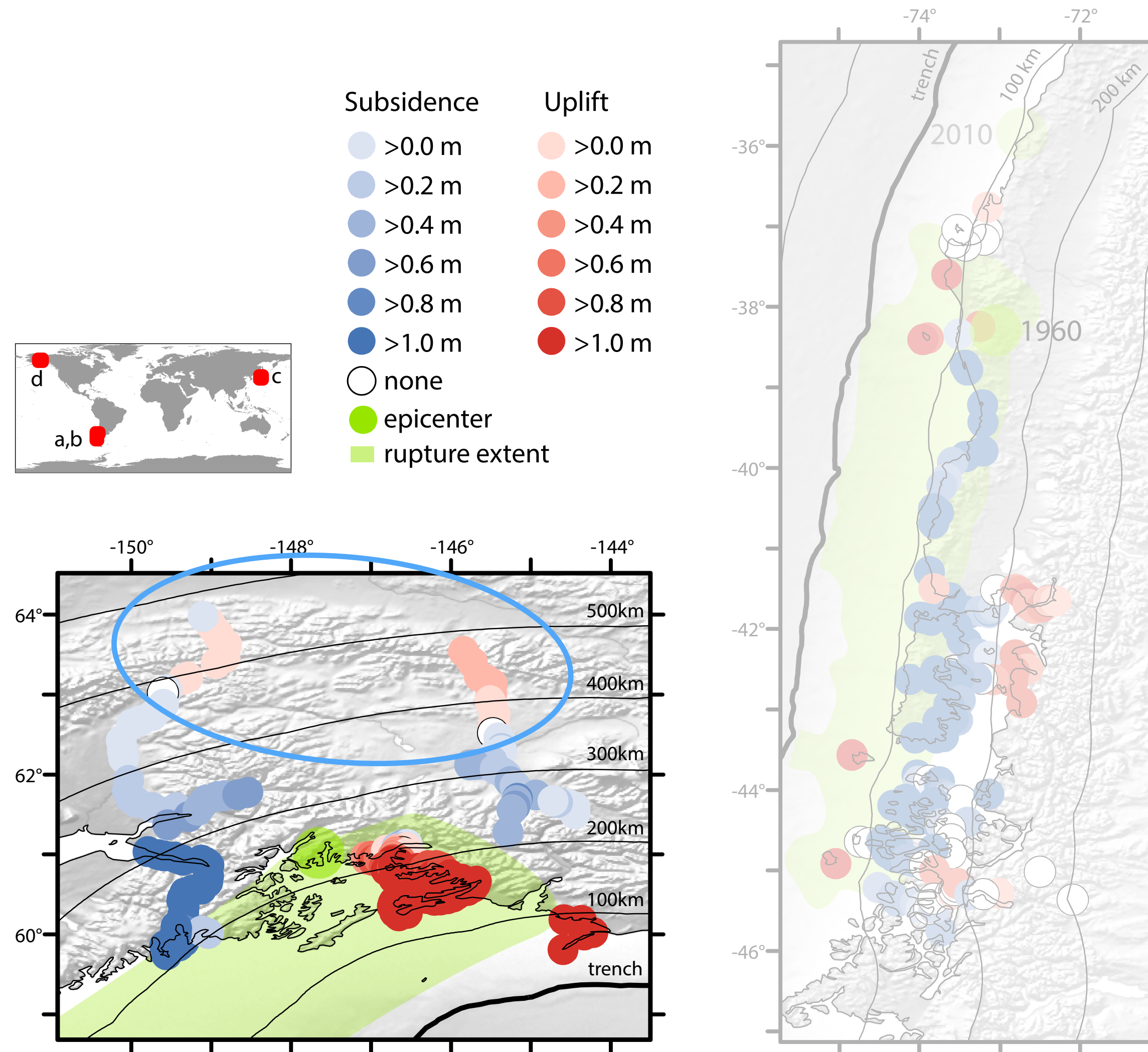
- STM models identified a new physical phenomena: a secondary zone of coseismic uplift
- Its existence is confirmed for 4 out of 4 megathrust earthquakes analyzed, albeit at different distances and extents

1960 M9.5 Valdivia earthquake



A secondary zone of coseismic uplift exists

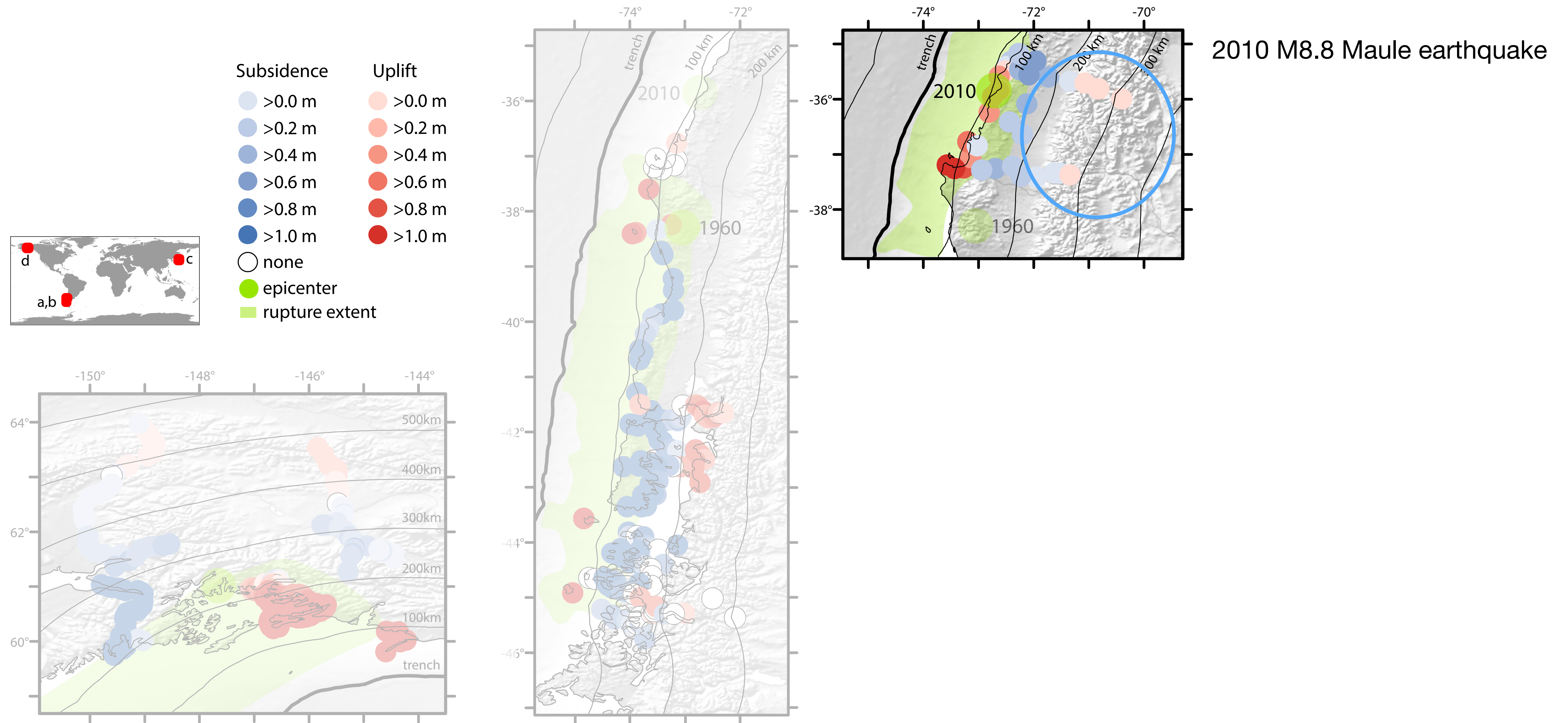
- STM models identified a new physical phenomena
- Its existence is confirmed for 4 out of 4 megathrust earthquakes analyzed, albeit at different distances and extents



1964 M9.2 Alaska earthquake

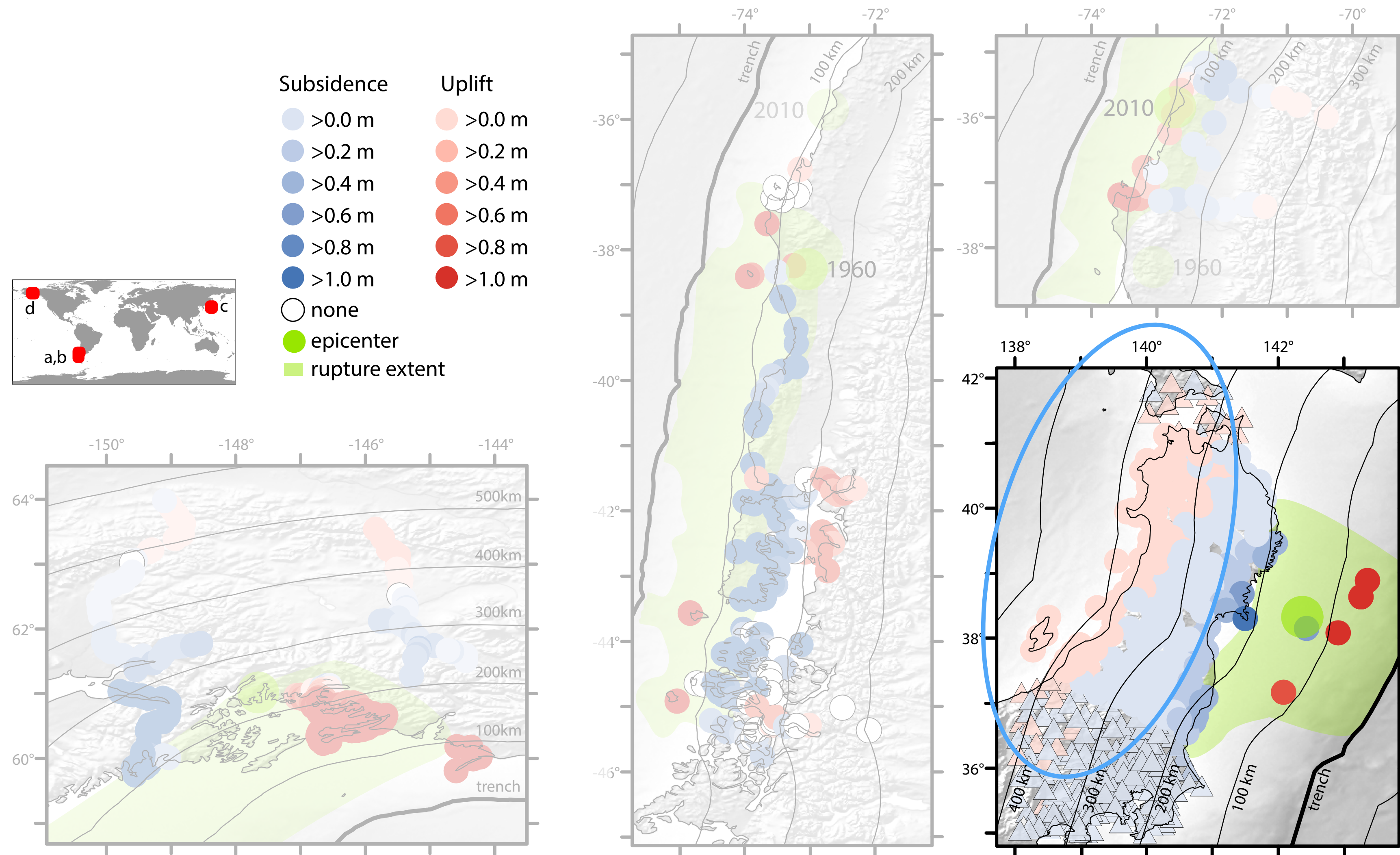
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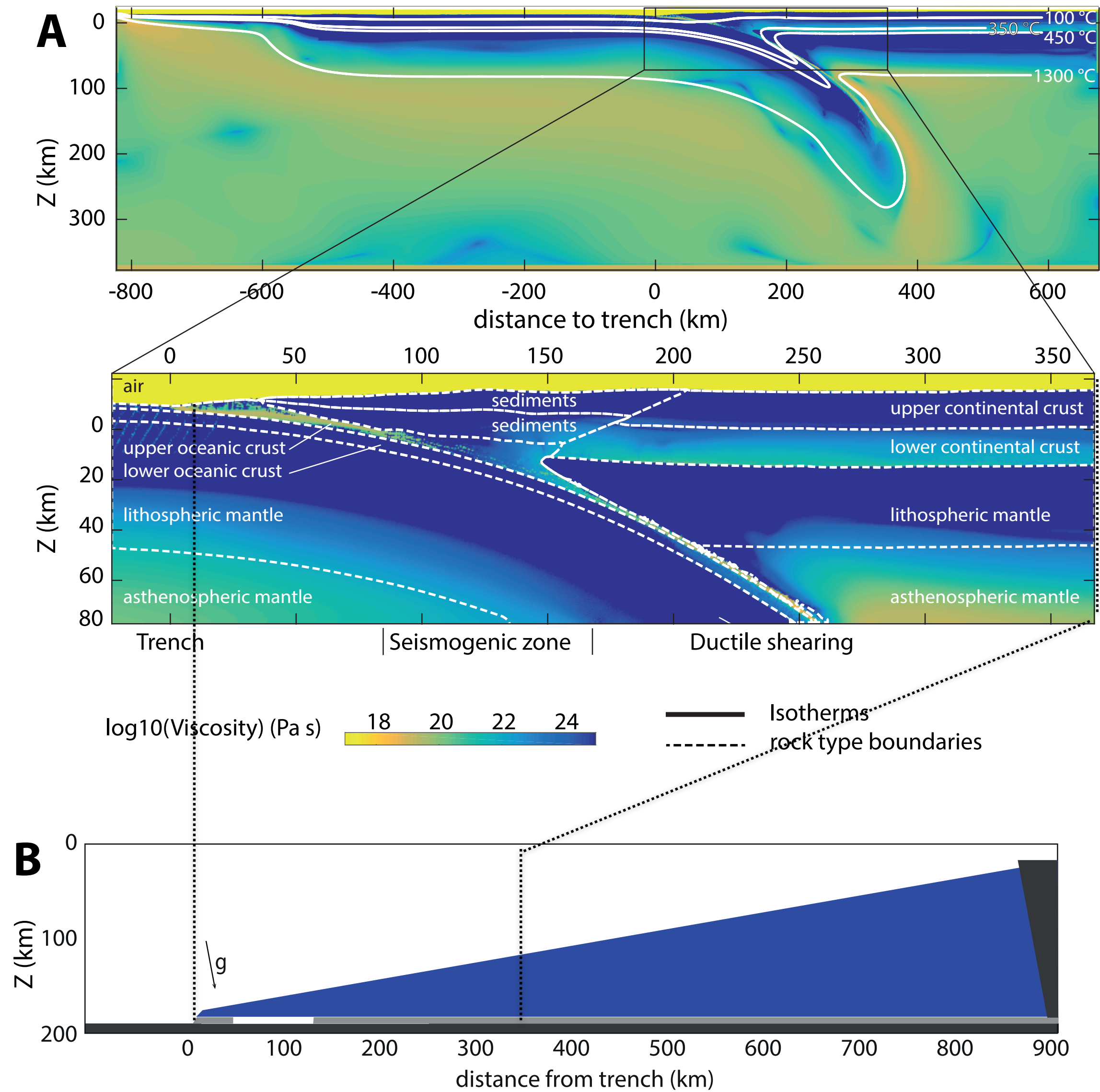
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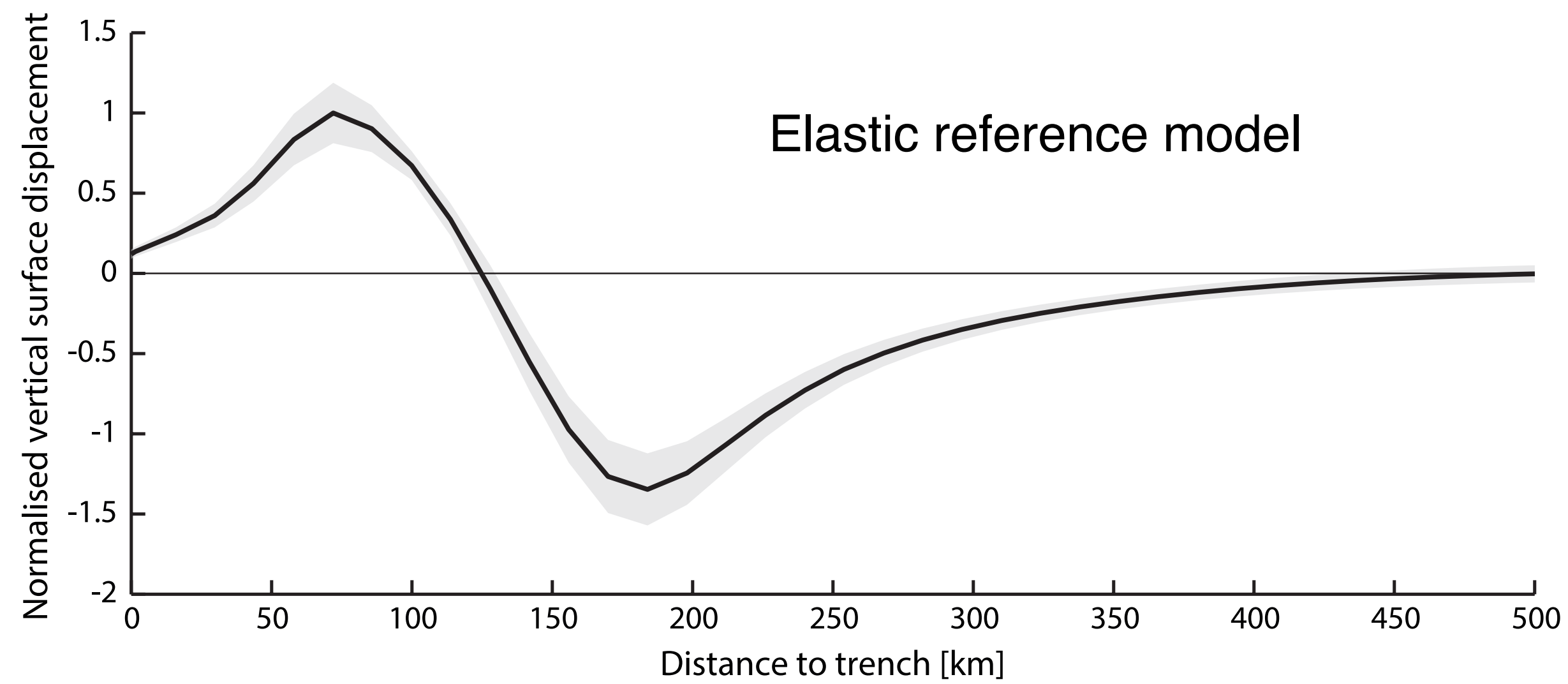
2011 M9.0 Tohoku earthquake

Systematic exploration through simple and complex models

- First attempt to understand based on simple models

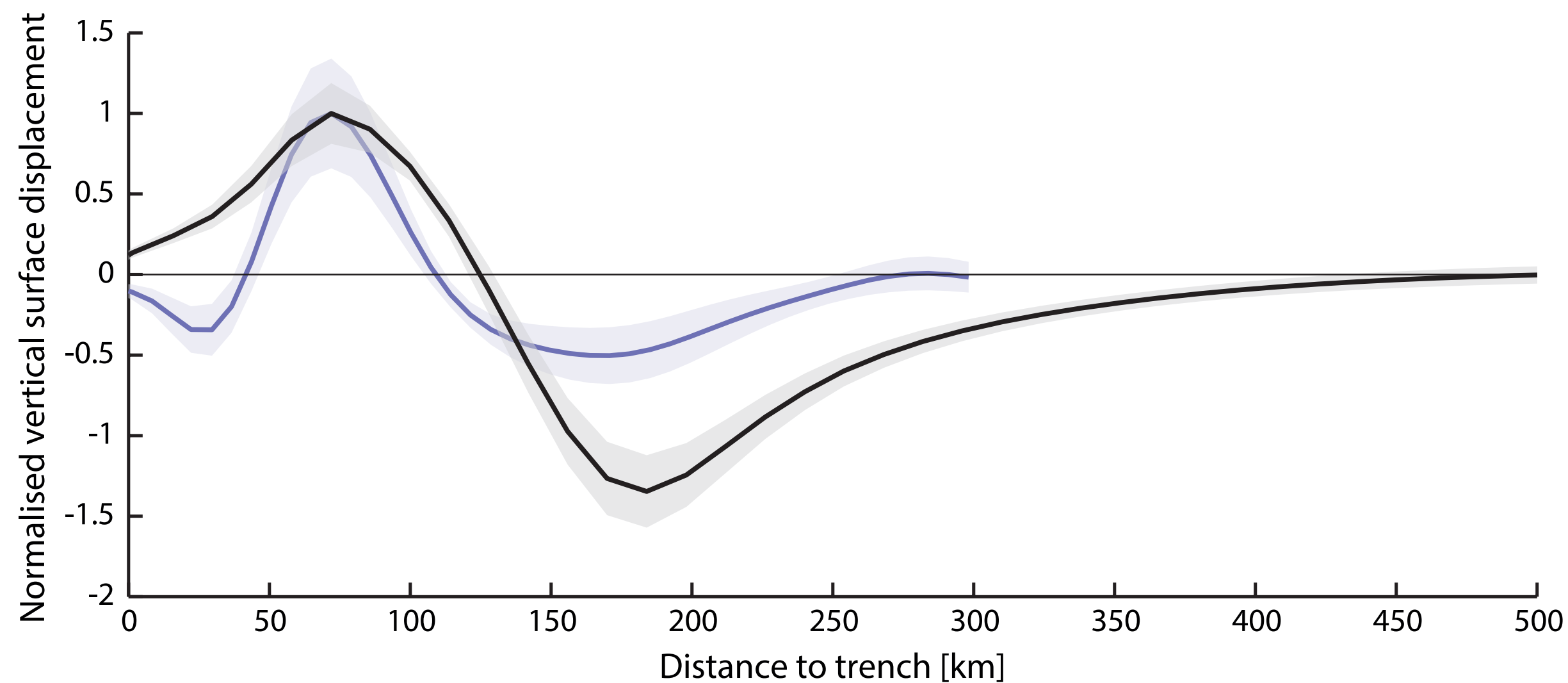
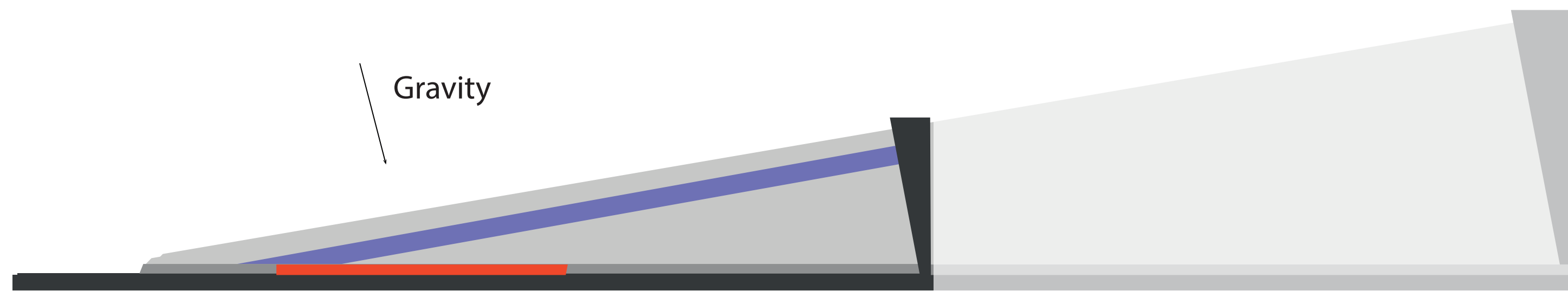


What is the governing physical mechanism?



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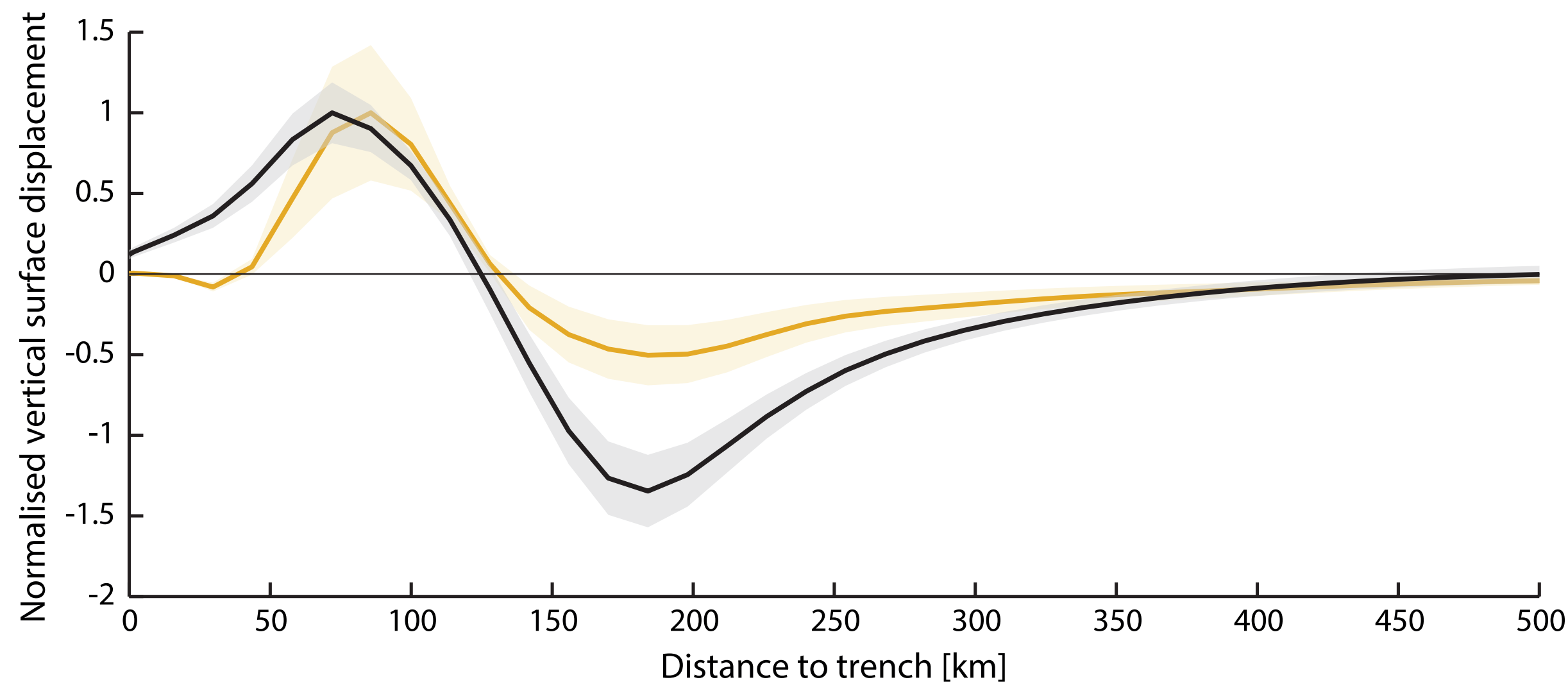
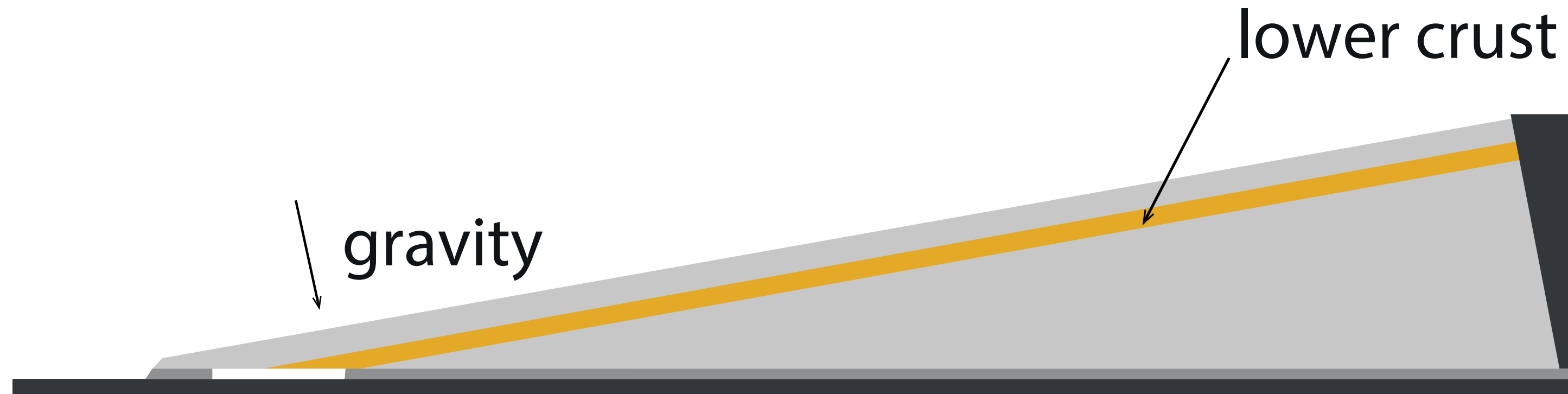
- (1) Elastic rebound after interseismic buckling of **visco**-elastically layered lithosphere



- » Visco-elastic structure is important for surface displacements
- » A back stop could play a role, but it is arguable if that is persistently present

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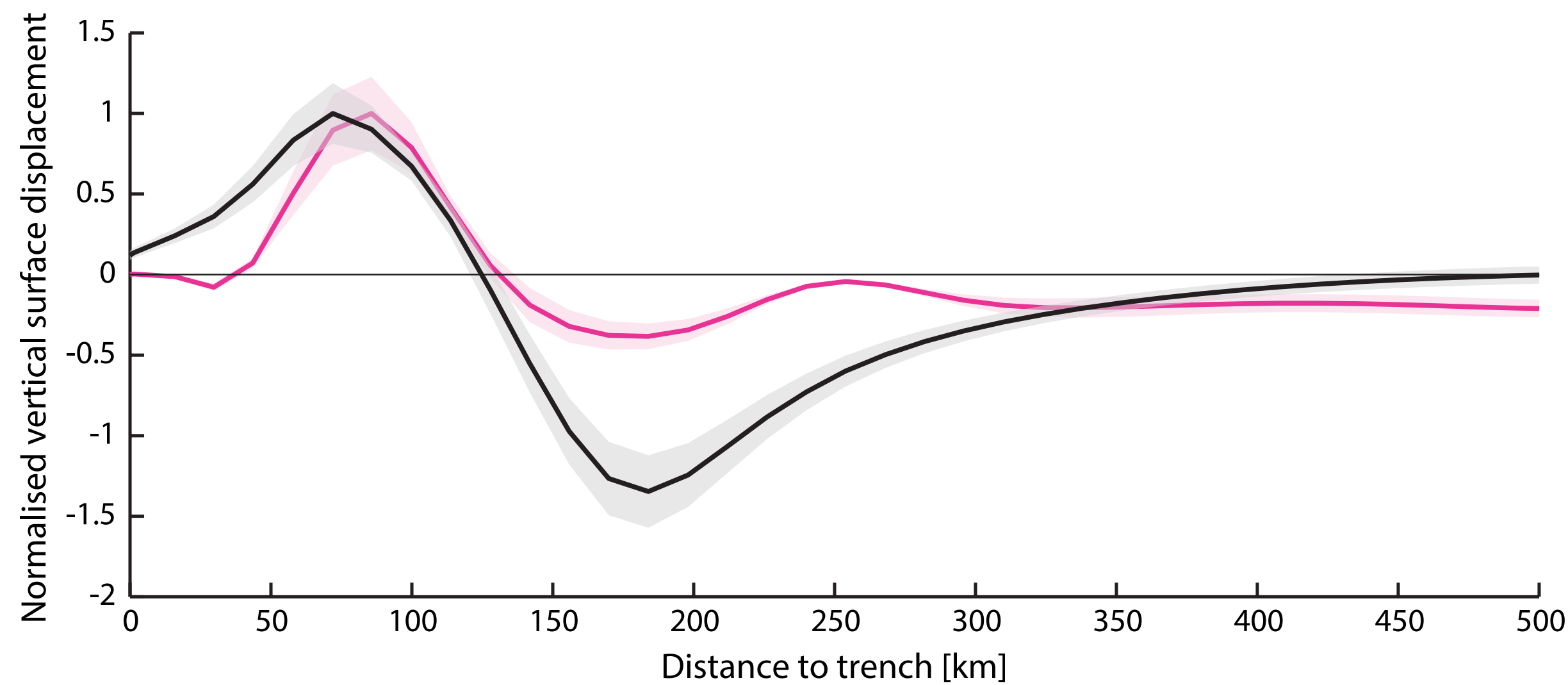
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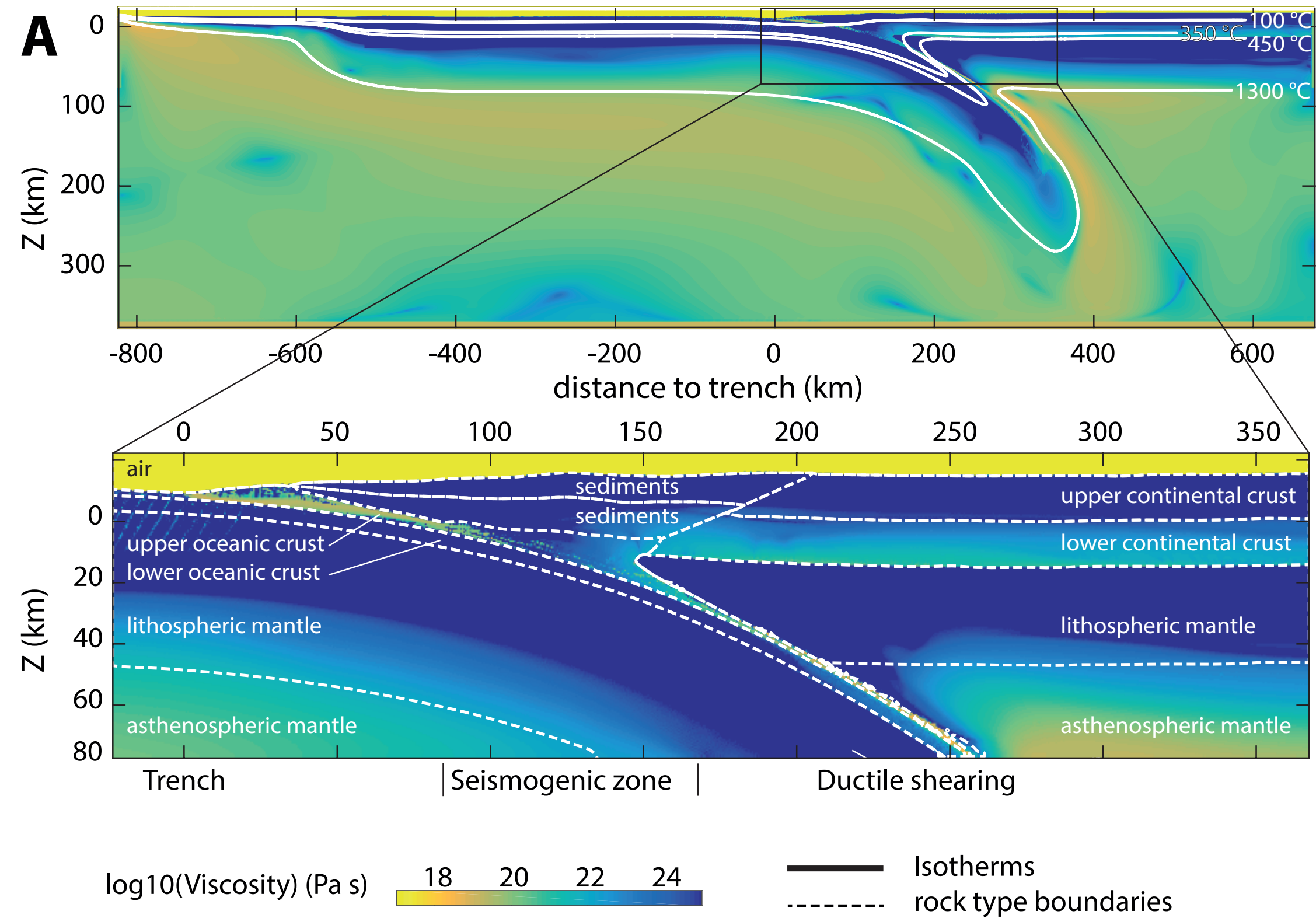
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» Visco-elastic structure is important for surface displacements

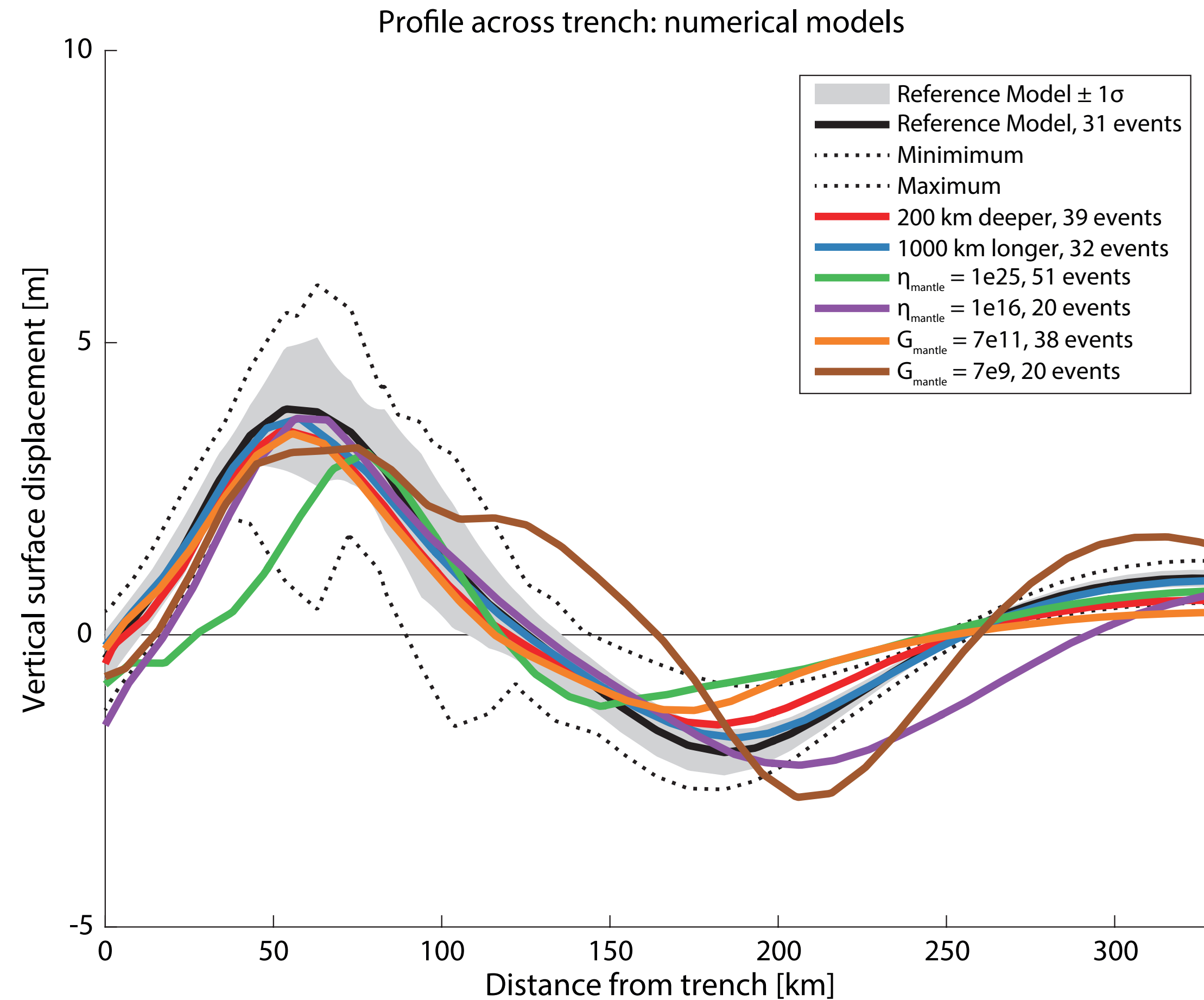
Systematic exploration through simple and complex models

- First mechanism is rebound of elastically buckled lithosphere → assess more realistic models

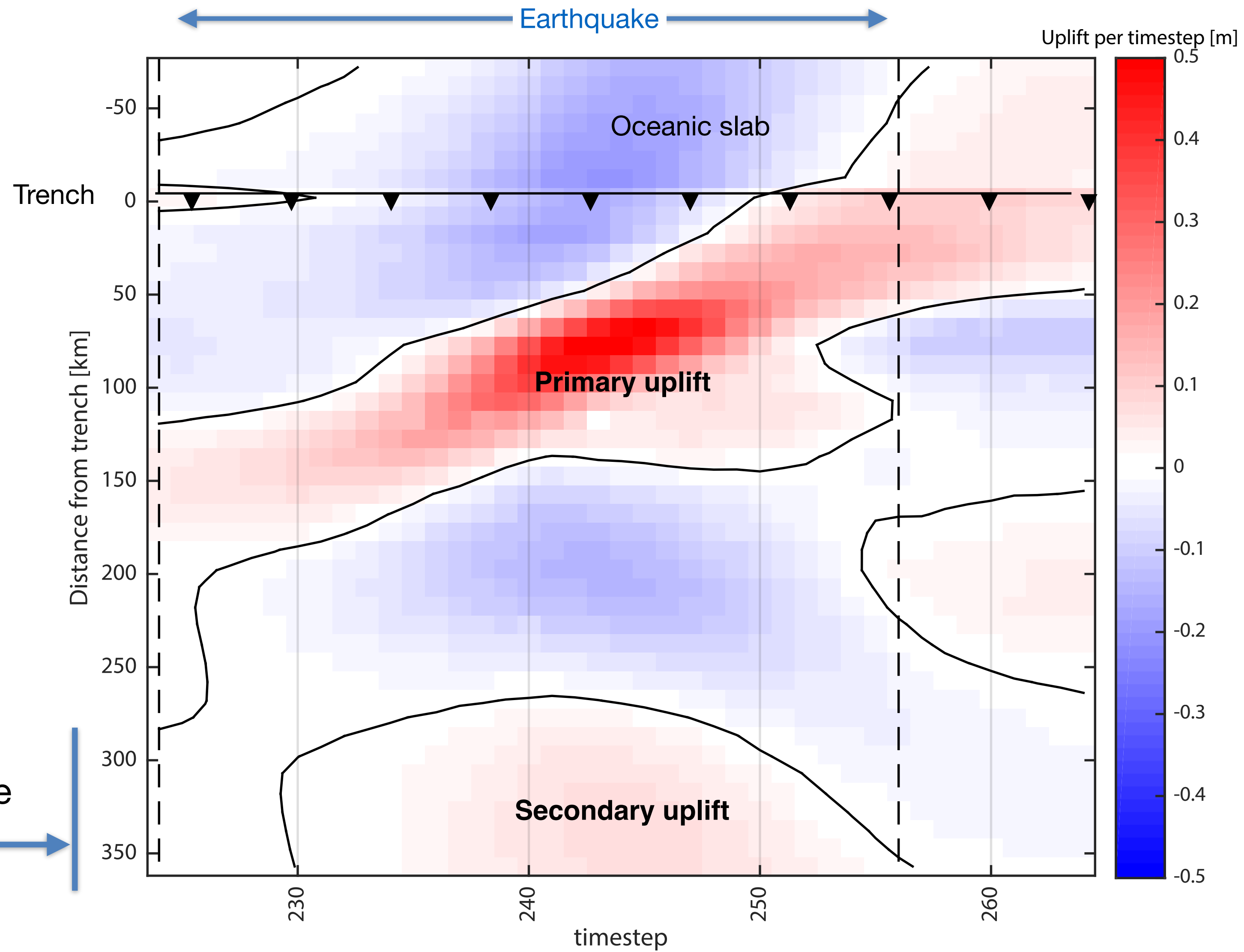


Consistent occurrence of secondary uplift

- In hundreds of realistic experiments we are not able to remove uplift
 - Also not when elastic buckling is inhibited
- » basic mechanism is missing

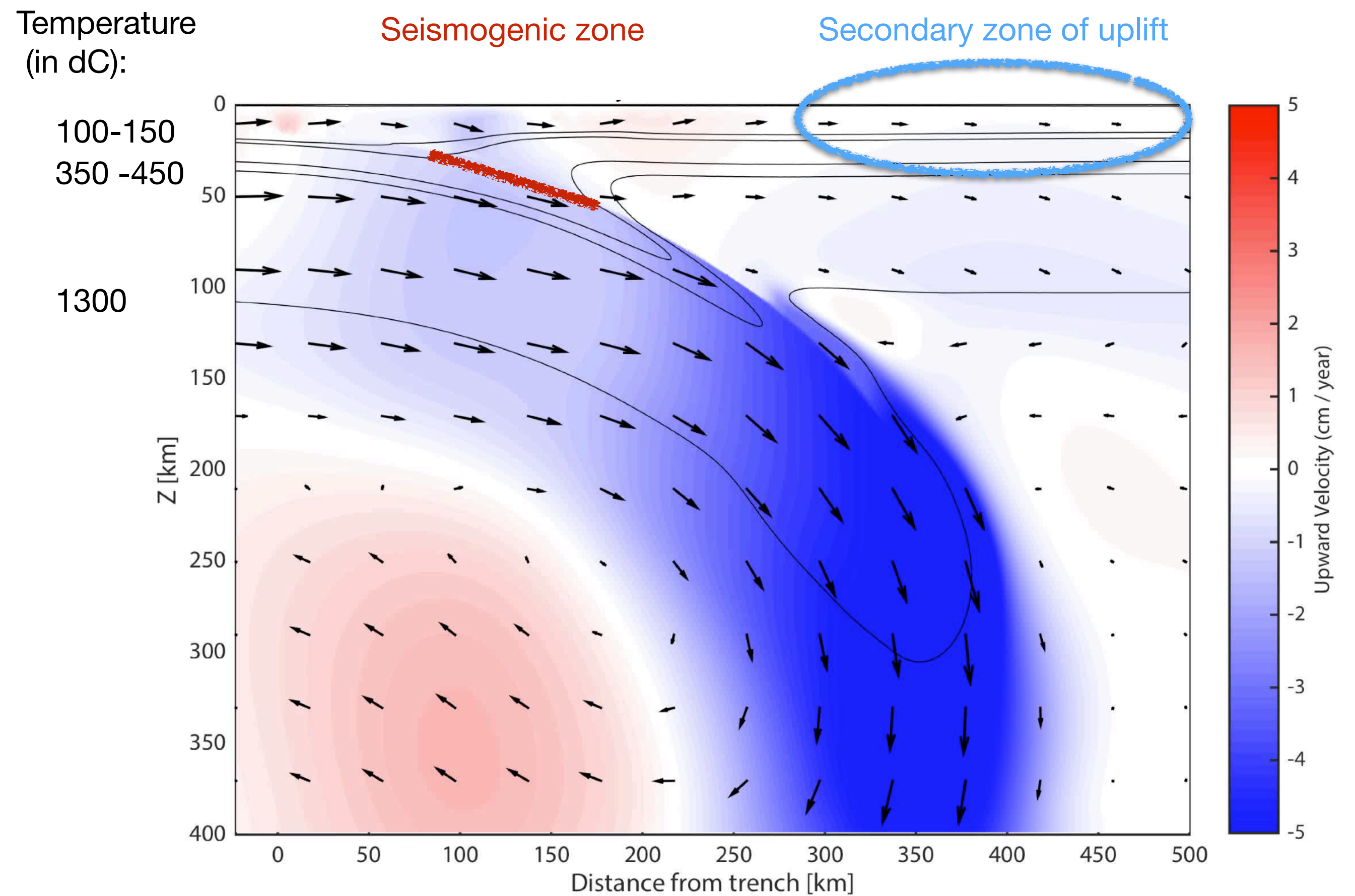
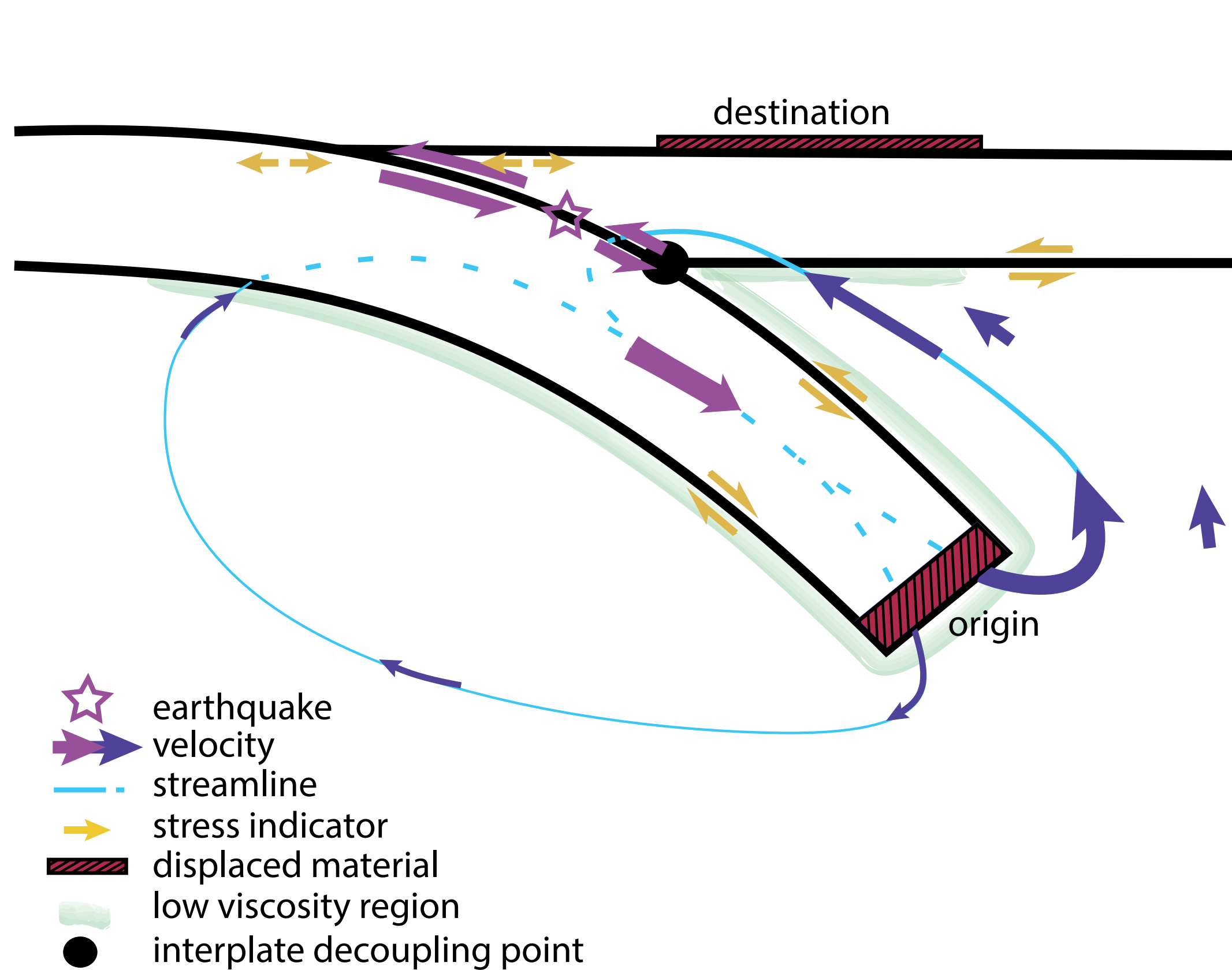


Spatiotemporal uplift in co- and postseismic period



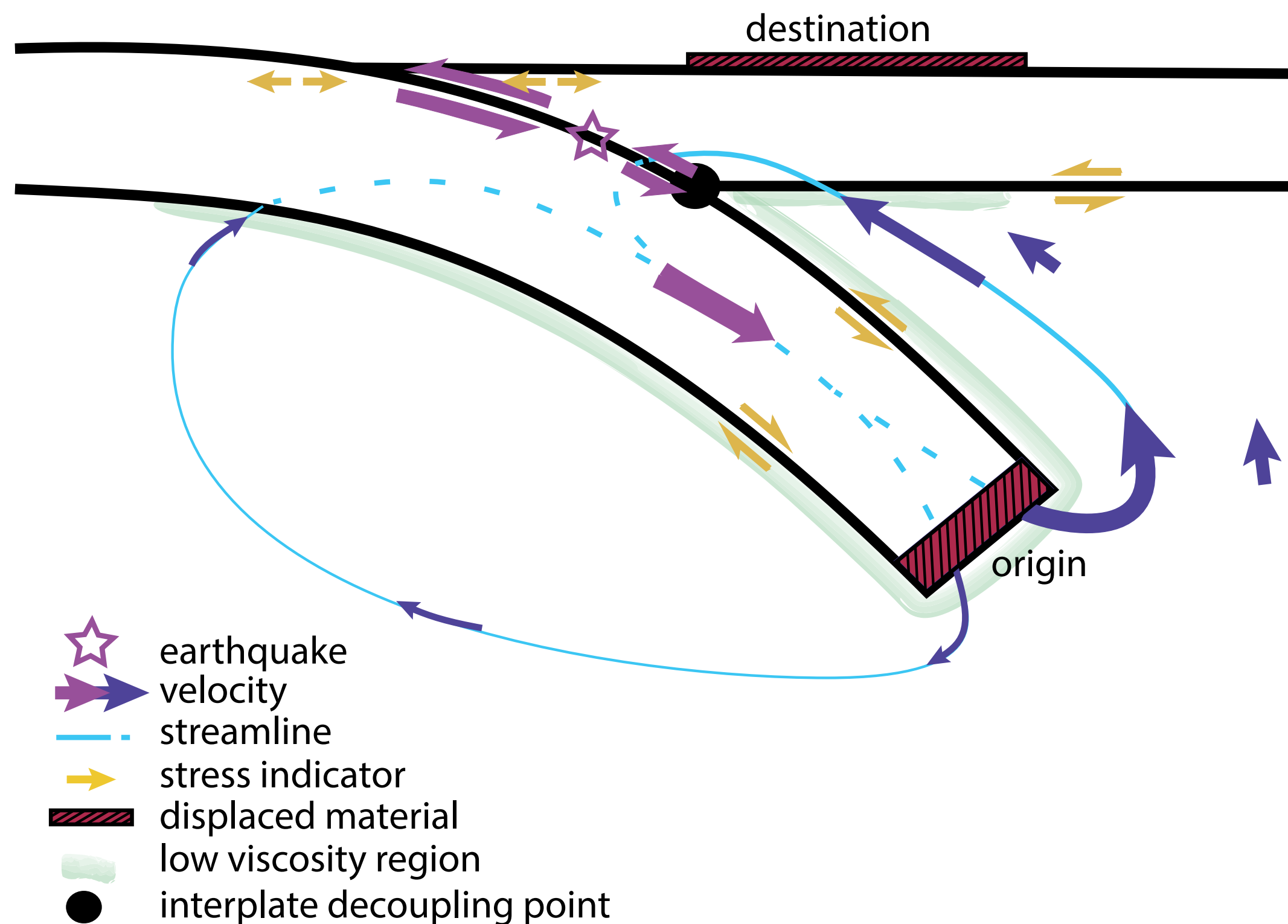
What is the governing physical mechanism?

- (1) Flexural buckling of a thin upper crust facilitated by a visco-elastic lower crust and mantle /backstop
- (2) Penetration of oceanic slab that induces upward flow



Implications

» Deep mantle processes affect the shallow surface also at time scales of minutes and days



» Subduction is NOT a gradual process

» Could see earthquakes as integral driver of mantle flow

• Difficult to apprehend slab-mantle response at such time scales?
(know behaves elastically during seismic wave propagation)

• Several times over last years we tried to disprove it, but we could not...

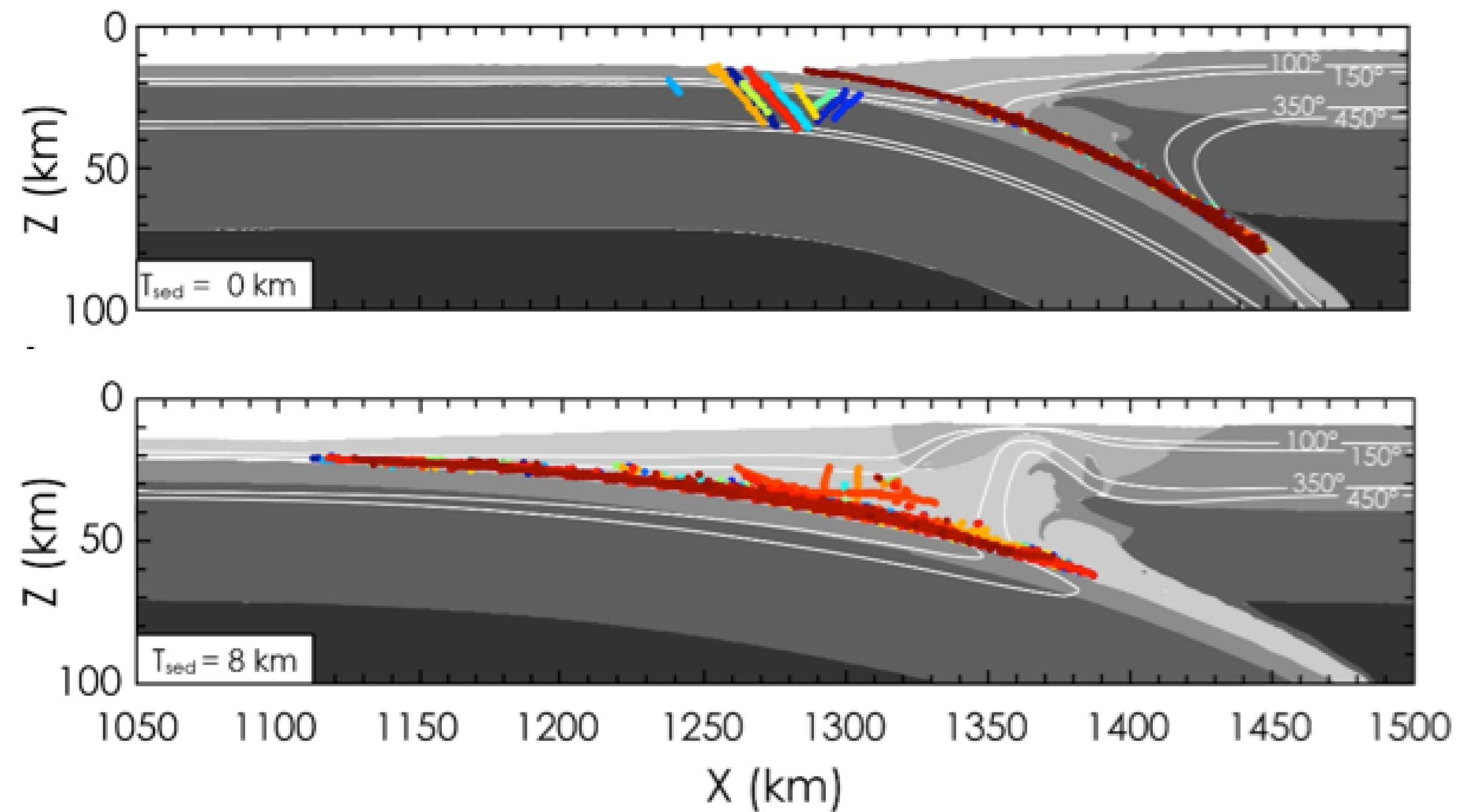
What do new data show?

- Model predictions on secondary zone of uplift keep being confirmed by new data
 - Predicted Secondary Zone of Interseismic Subsidence is observed
 - Revealed across Northwest Pacific (*Bill Hammond et al.*)
 - Postseismic data 2010 and 2011 events show majority is coseismic
 - Korean peninsula uplifts in days after 2011 M9 Tohoku earthquake
(*Kim and Bae, 2012*)

Presentation objectives

» Illustrate how lithosphere- and mantle dynamics, structure and rheology can influence shallow tectonics and seismicity

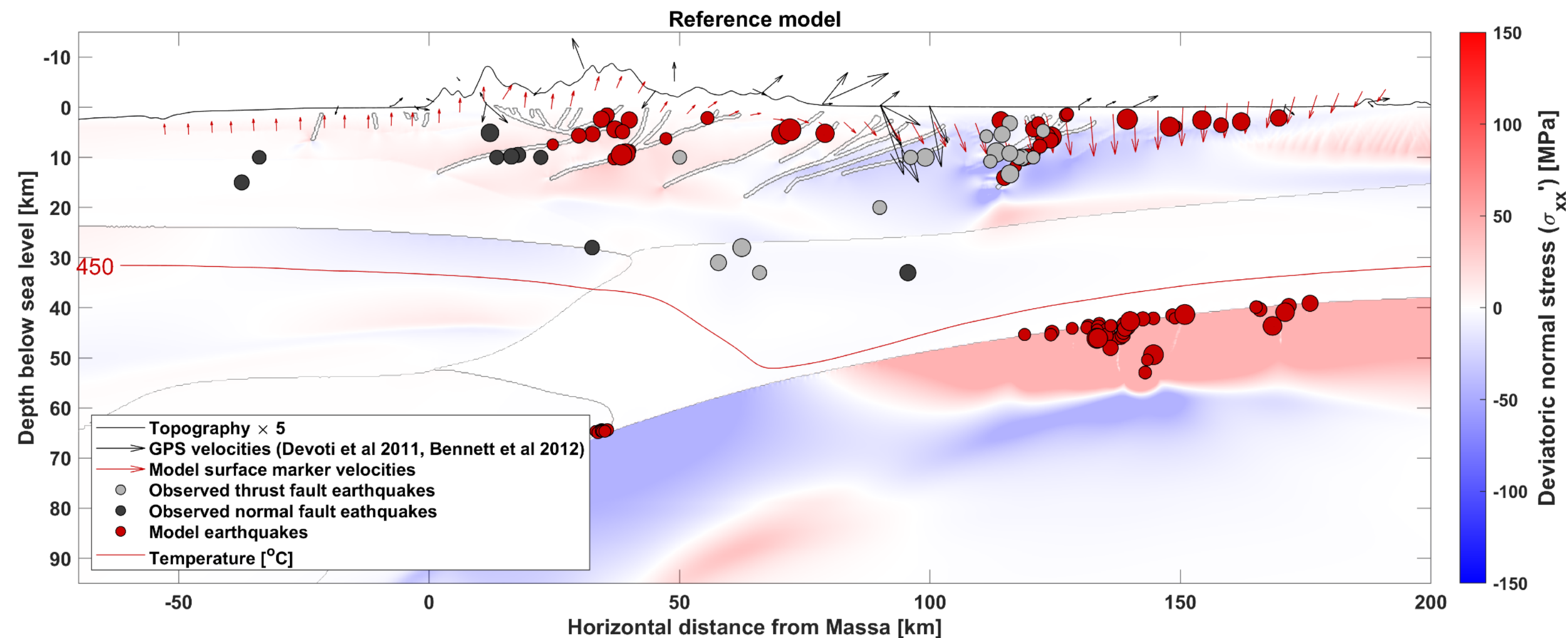
- Does incoming sediment thickness increase maximum earthquake magnitude?
 - Yes; larger sedimentary wedge → trench moves seaward and slab unbends → slab dip ↓ → seismogenic width ↑ → M_{max} ↑
 - Modeling long-term dynamics and sediment presence increases M_{max} by an order of magnitude !



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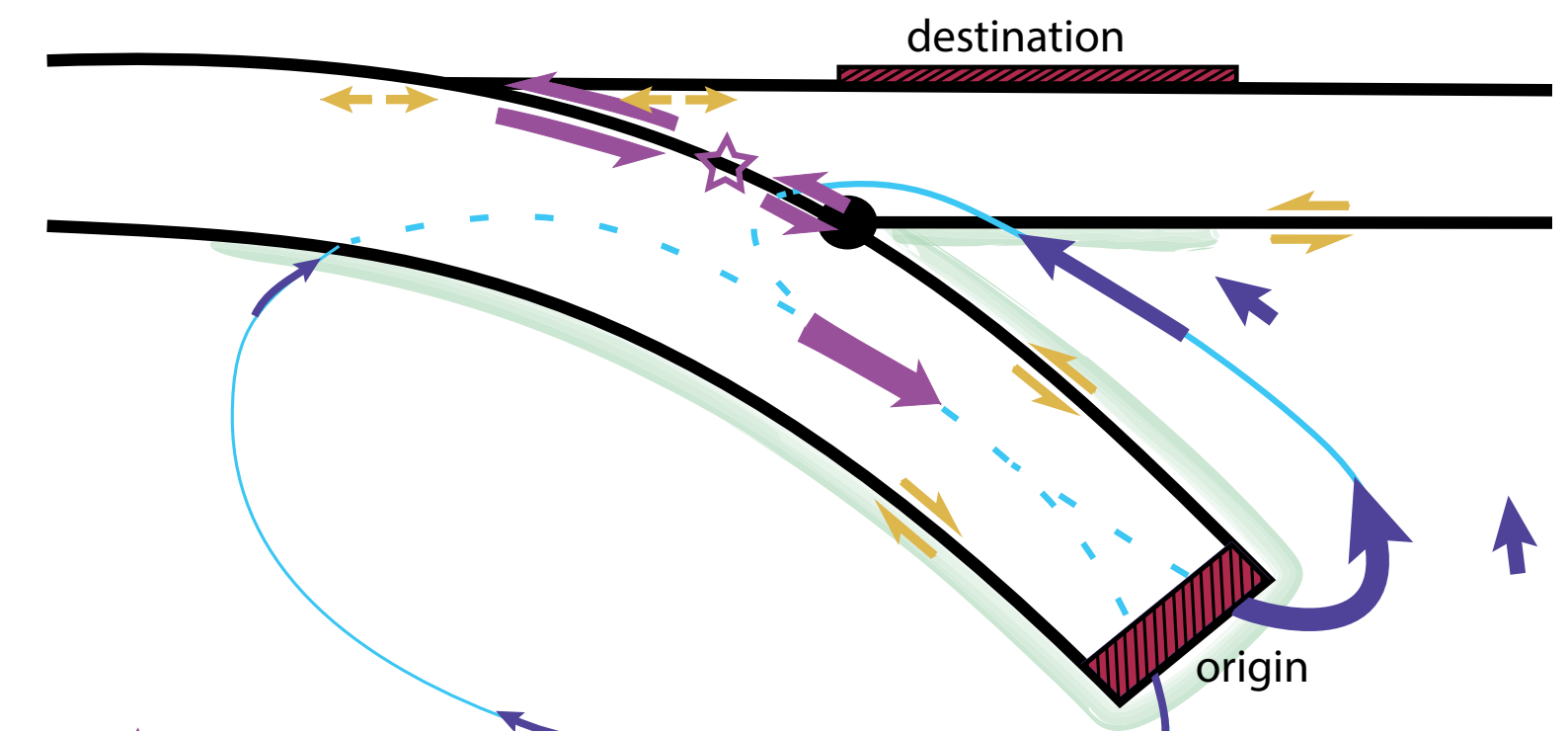
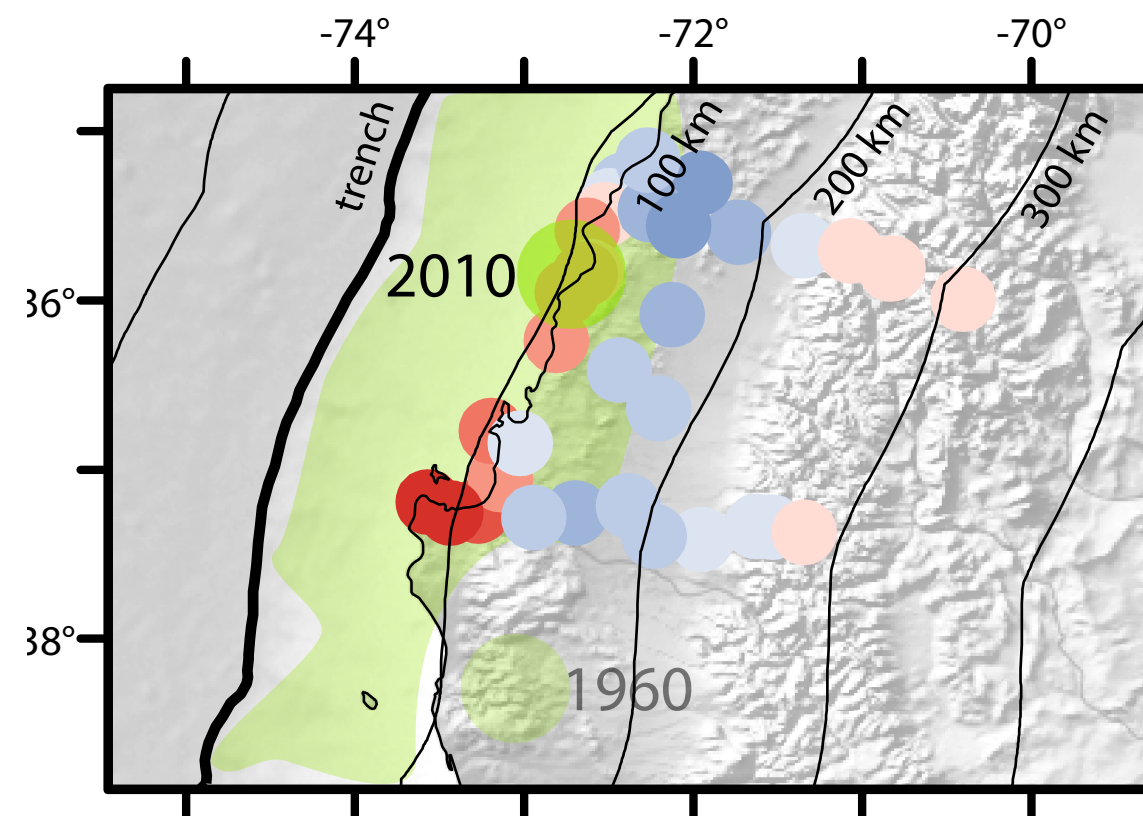
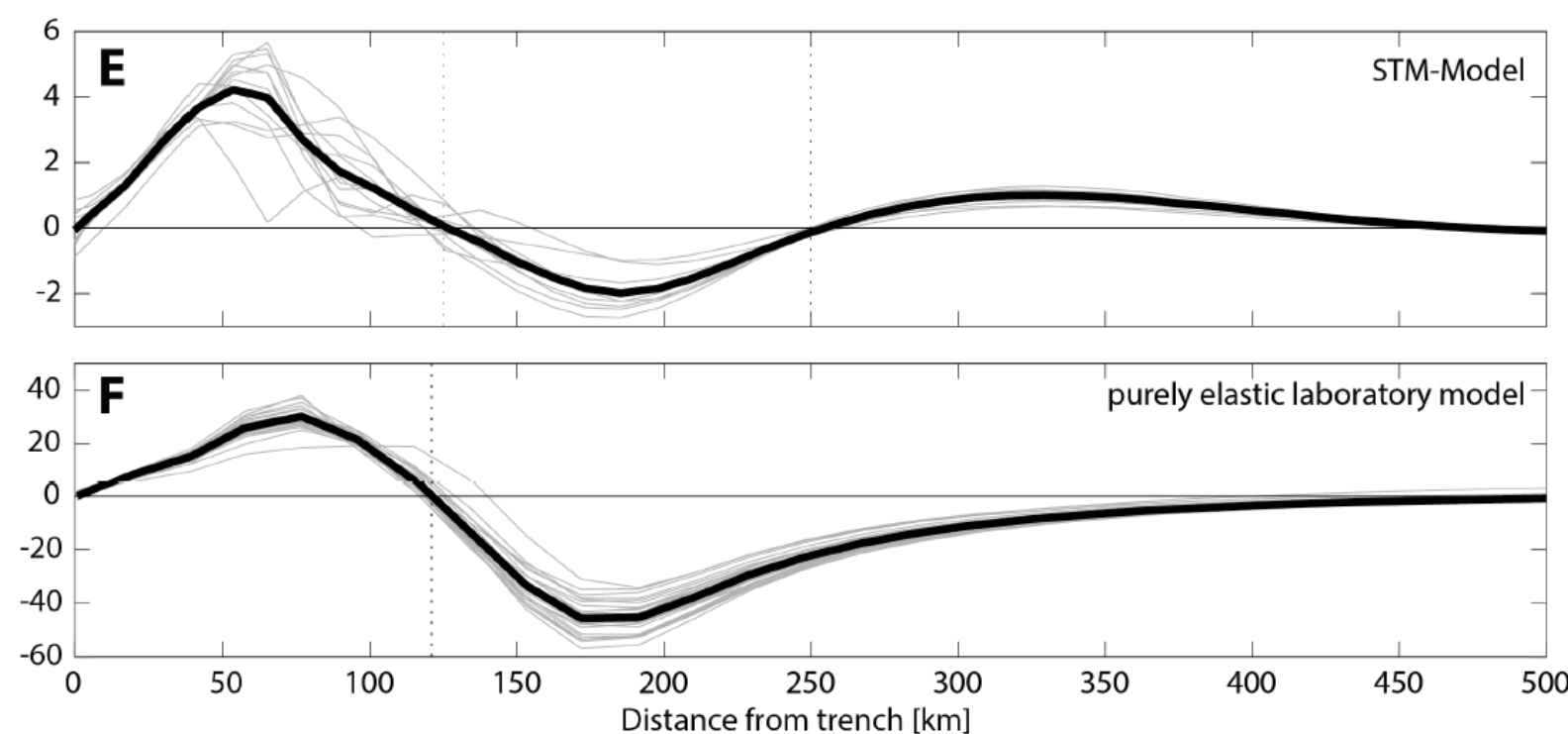
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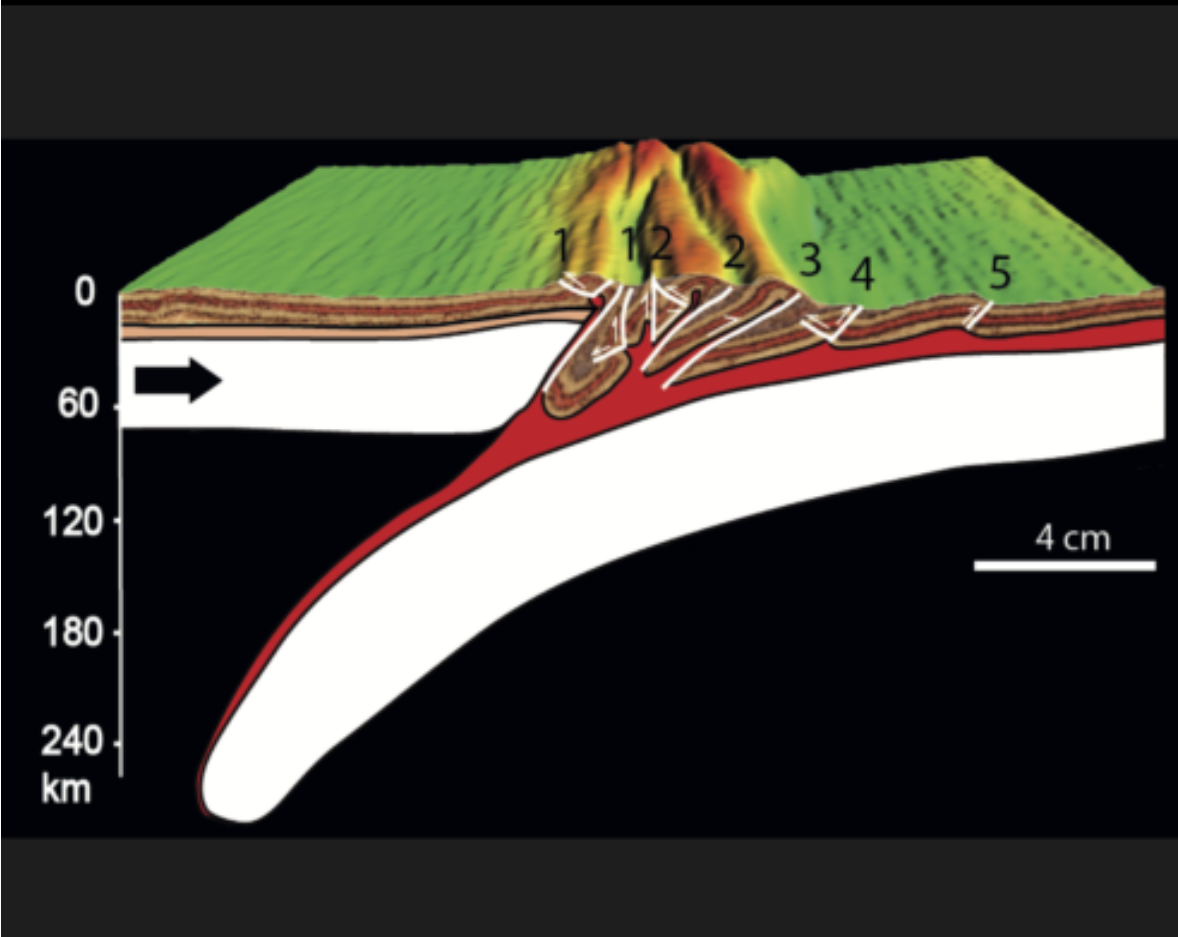
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- Do lower crustal and mantle depth temperature and rheology affect seismicity and tectonics?
 - Yes, in complex tectonic settings driven by deep loading, such as Northern Apennines
- Does the mantle affect surface displacements at time scales of minutes to days?
 - Yes, STM models predicted a secondary zone of “coseismic” uplift, which was confirmed by observations of 4 out of 4 great megathrust earthquakes
 - Accelerated slab penetration causes upward return “flow”




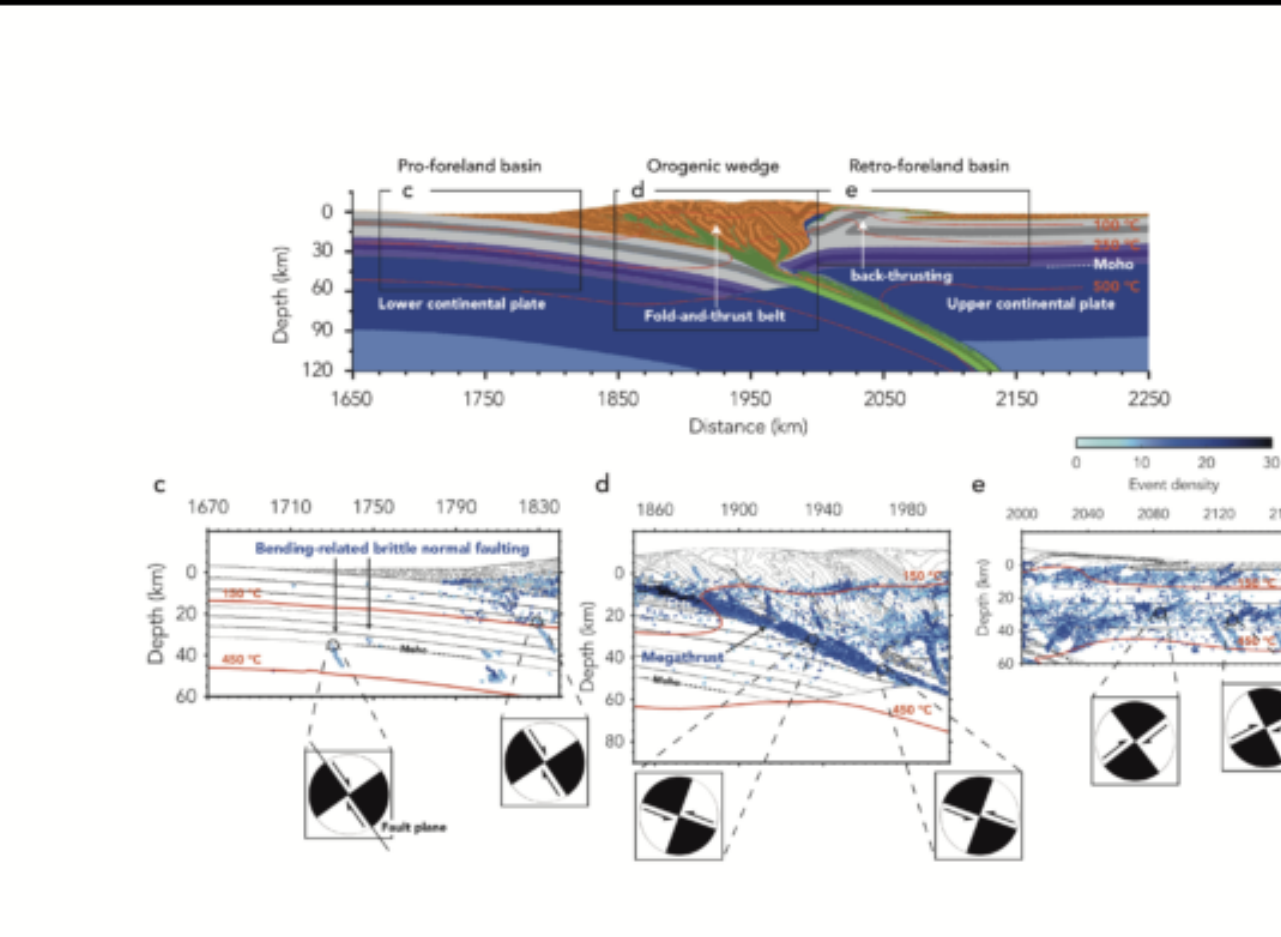
GeoMod 2020

[Welcome](#) [Venue](#) [Programme](#) [Organization](#) [Deadlines](#) [Contact](#)

<https://geomod2020.uu.nl>







SAVE THE DATE: June 28 - July 2, 2020

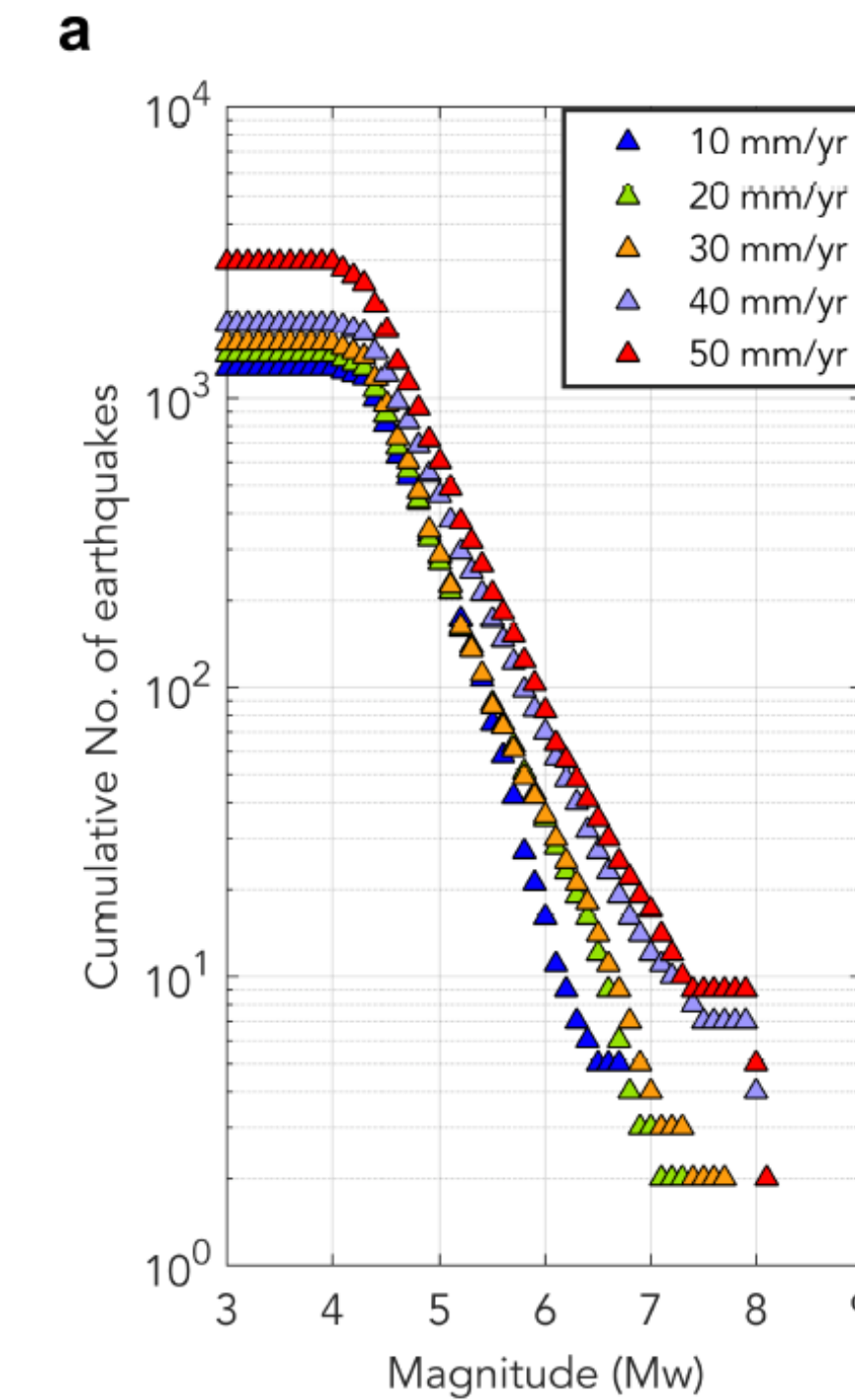
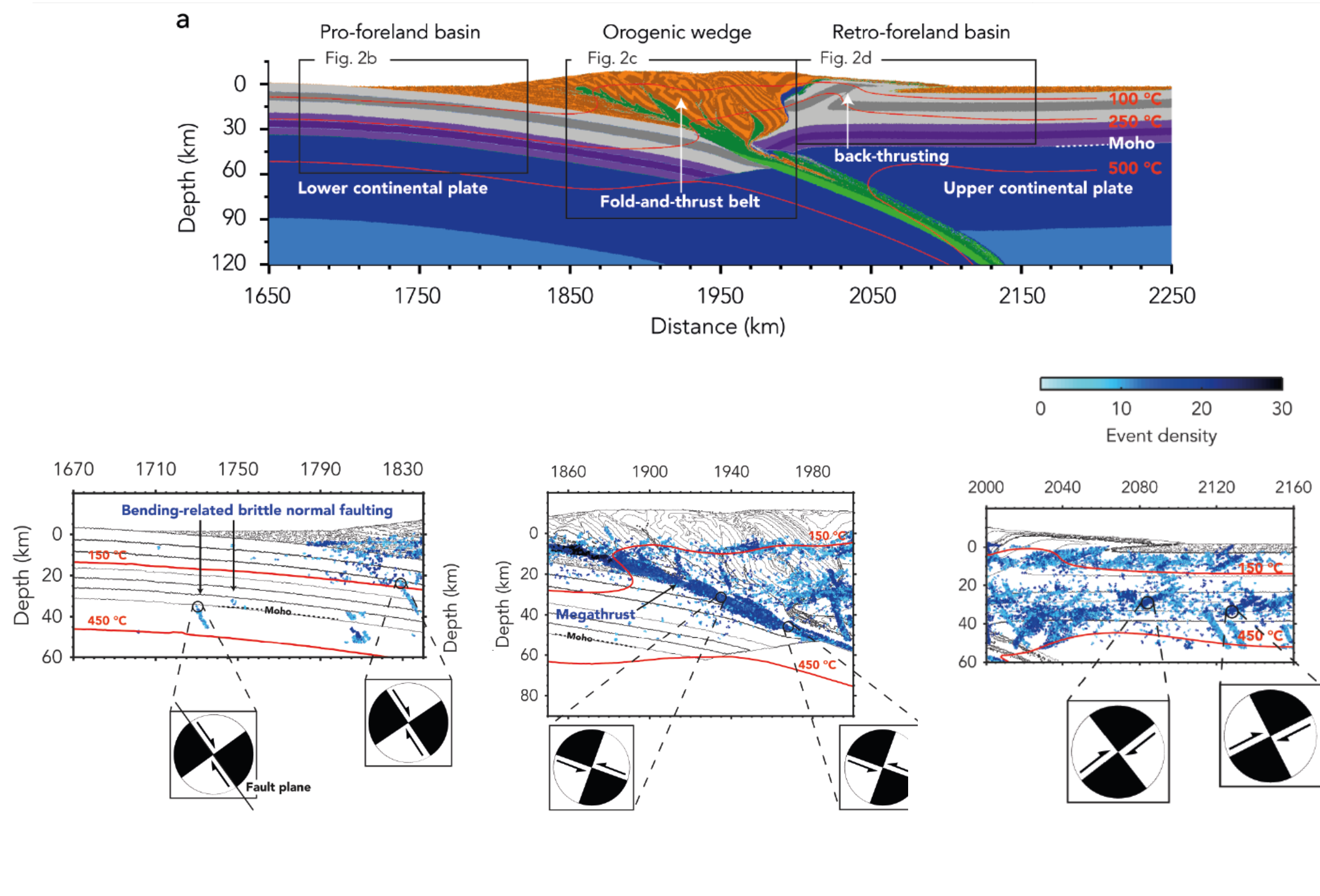
- Venue: Estate “Sunny Hill” near Utrecht, The Netherlands
- Host: Utrecht University
- Programme:
- 7 topical sessions
 - 18 key note presentations
 - vivid poster presentations
 - 2 evening lectures
 - visit to the new Earth Simulation Laboratory at Utrecht University



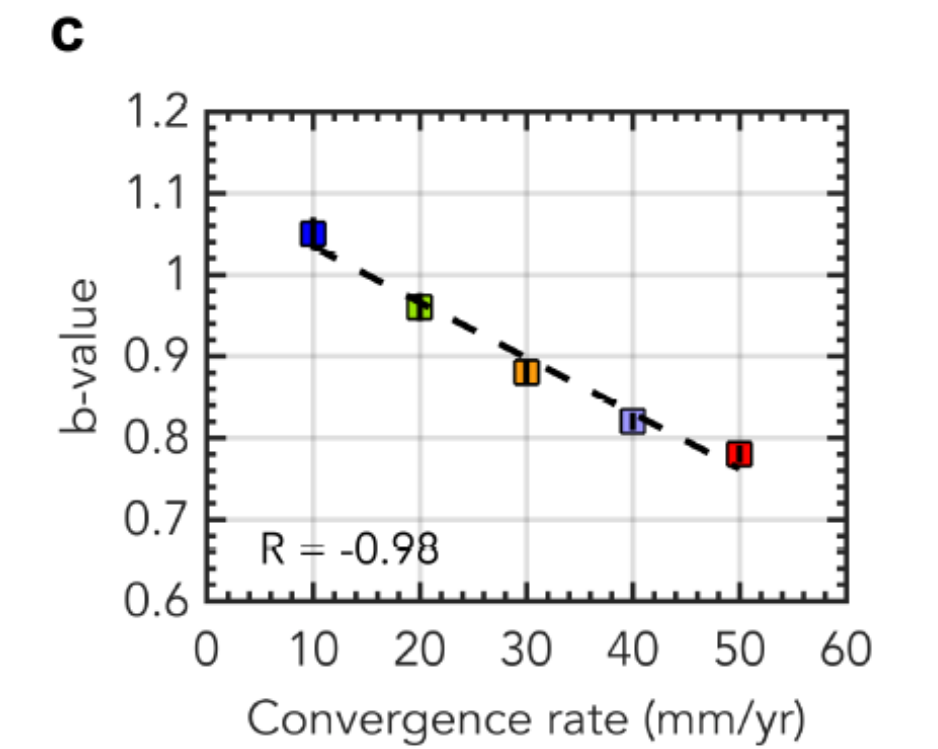
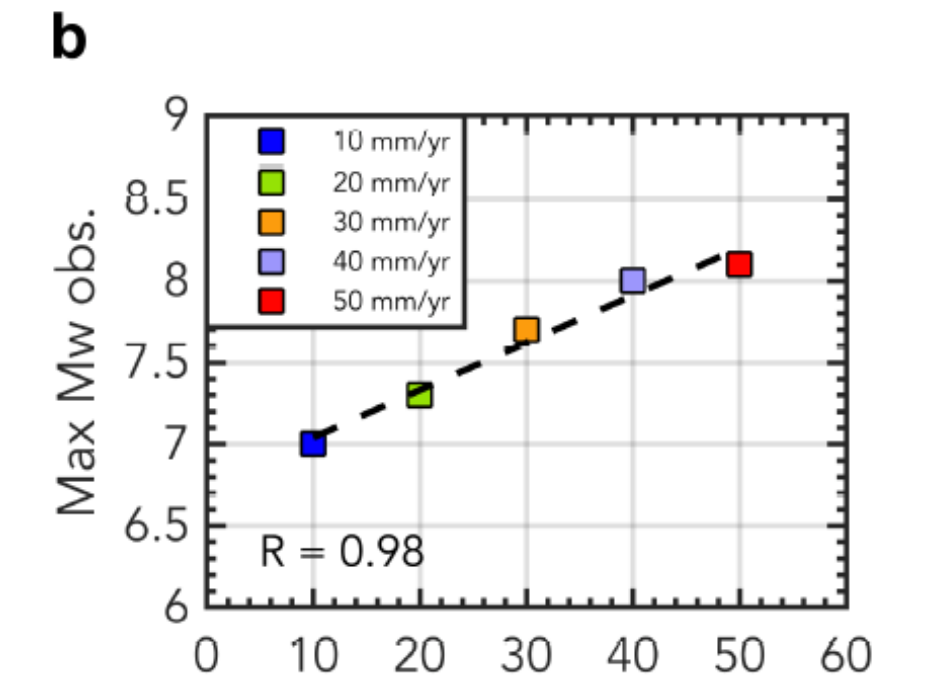
Earth Simulation Laboratory at UU

Role of convergence velocity V_c

- Mechanical experiments show V_c only increases seismic rate, not maximum magnitude M_{max} (e.g., Corbi et al., GRL, 2017).
- Adding temperature and long-term dynamics changes the story...



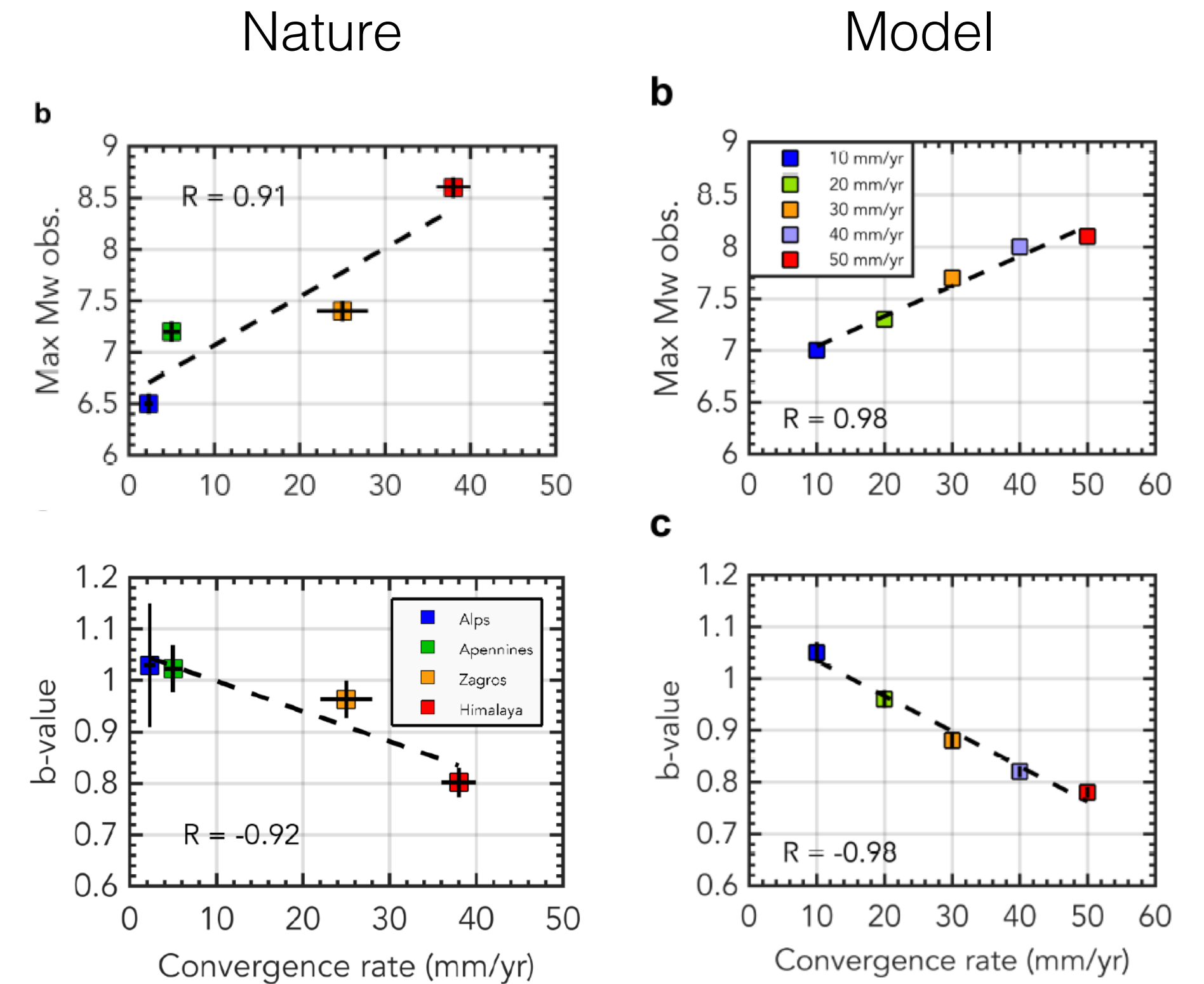
- M_{max} does increase!



- Gutenberg-Richter distribution is obtained!

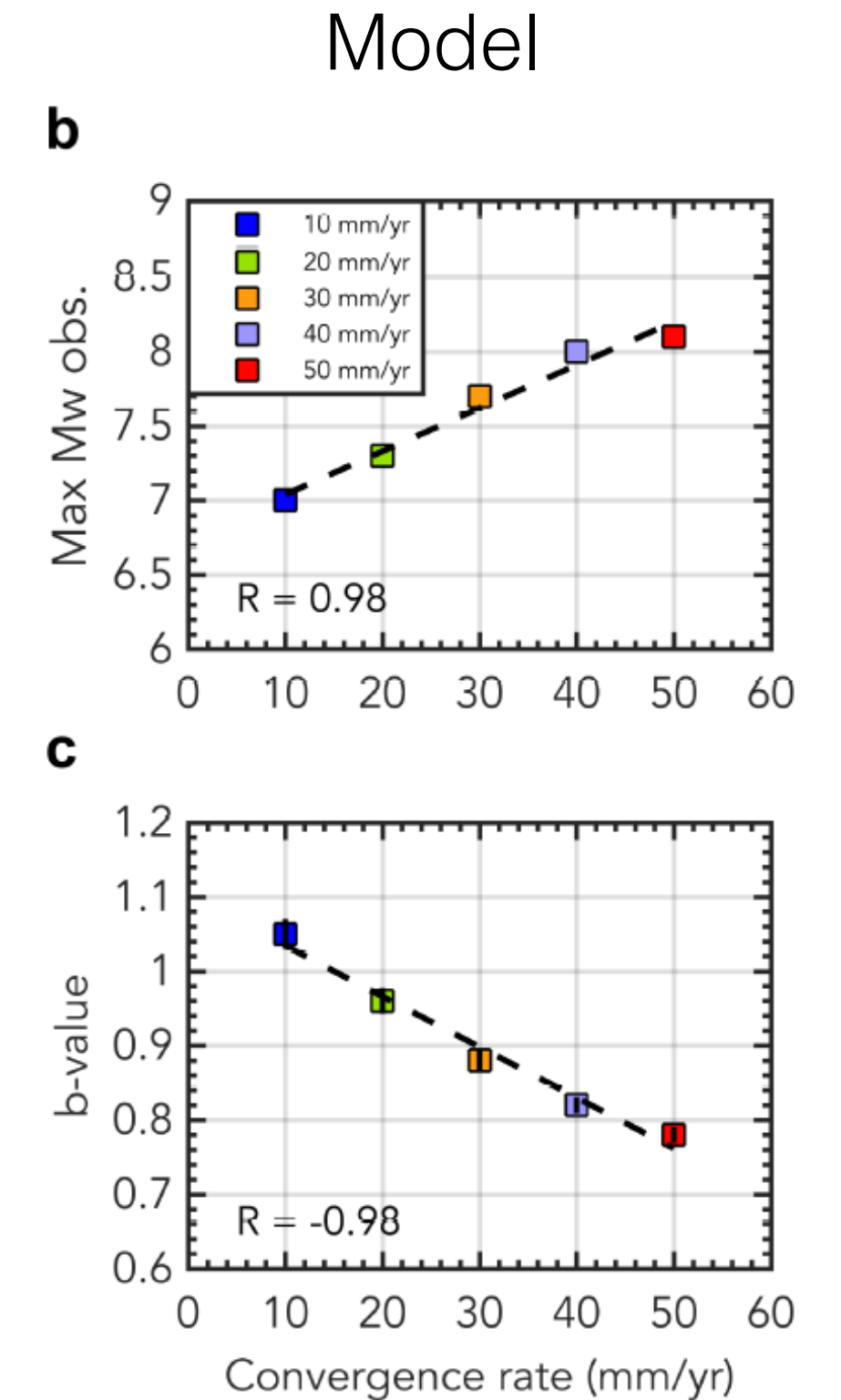
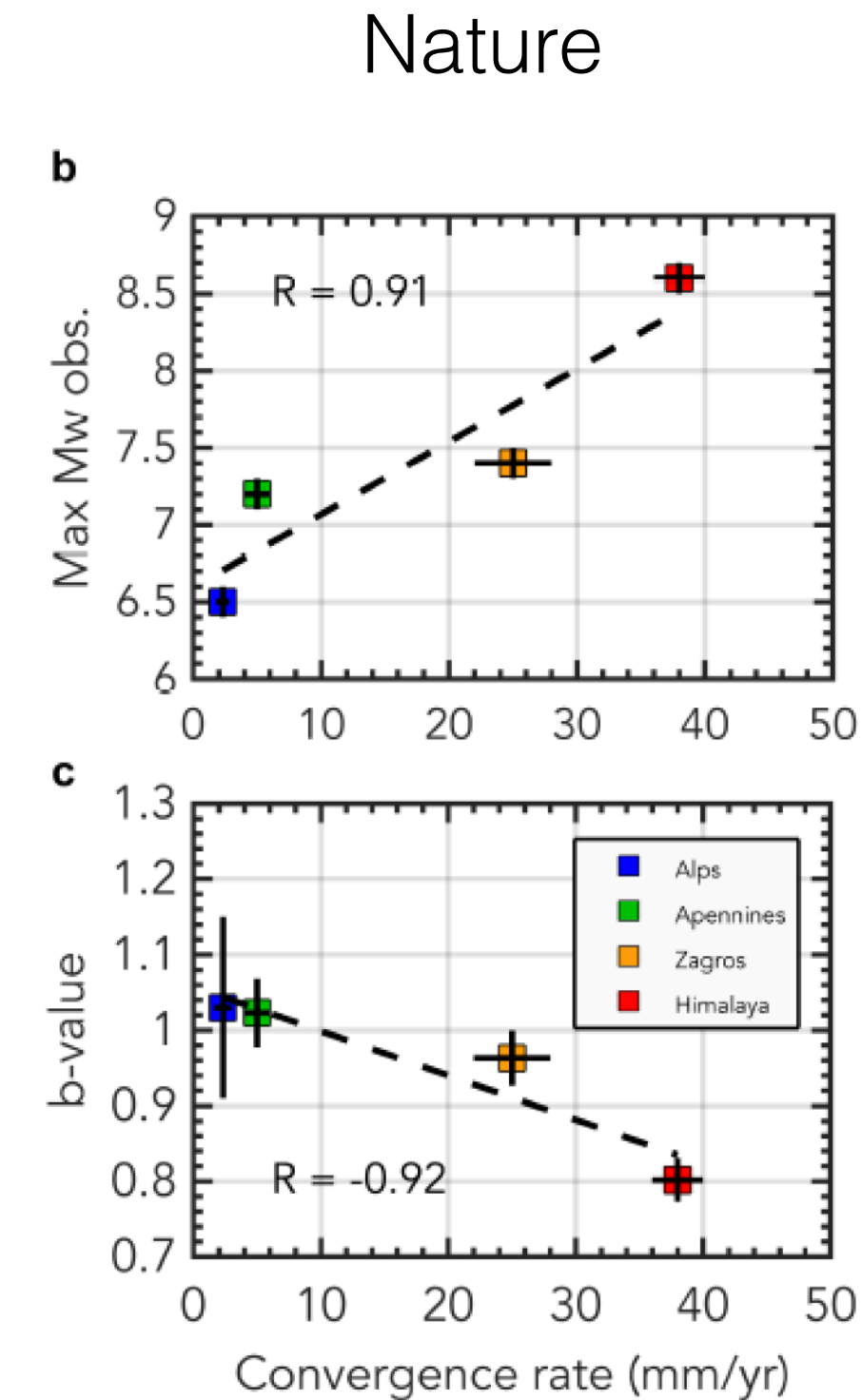
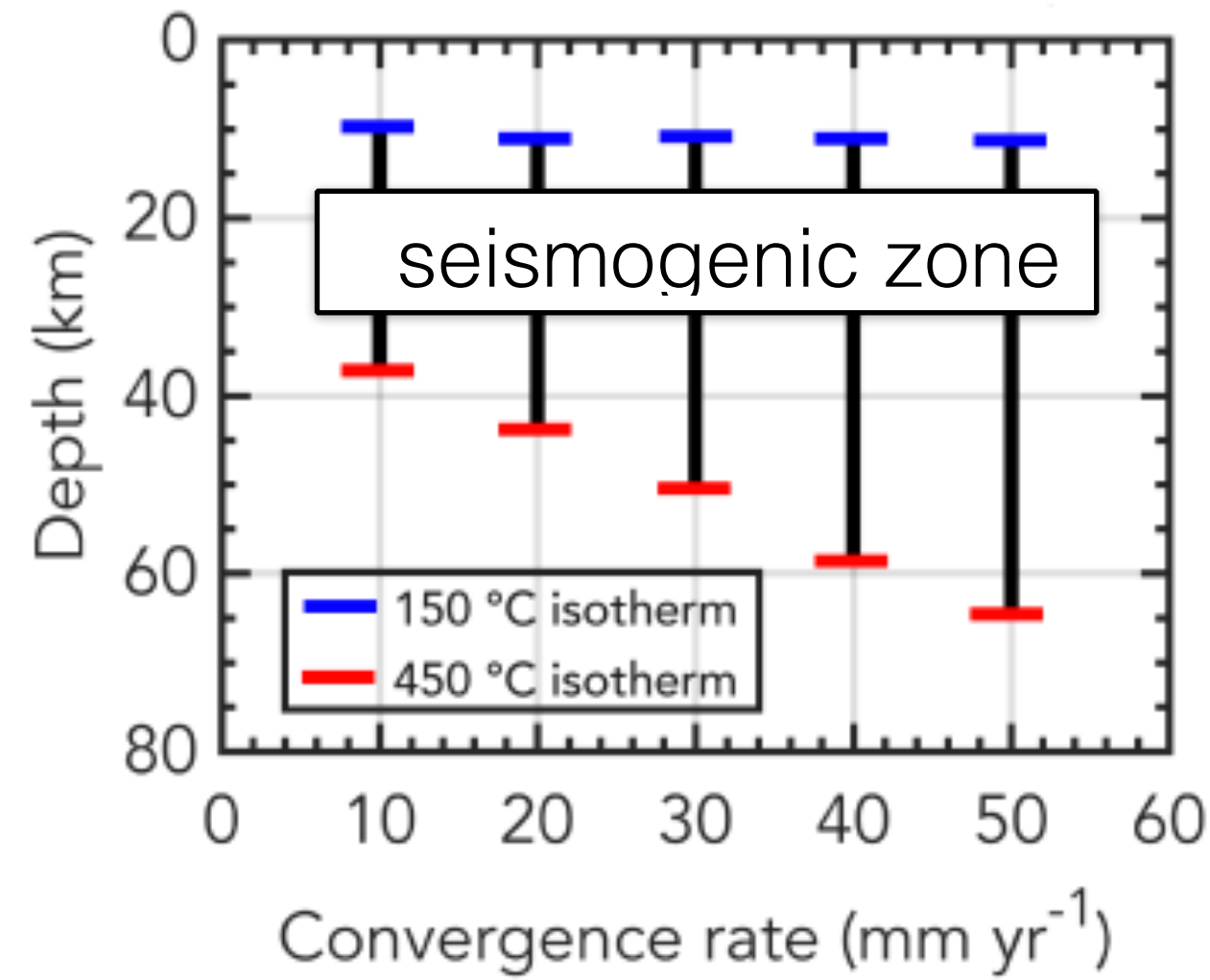
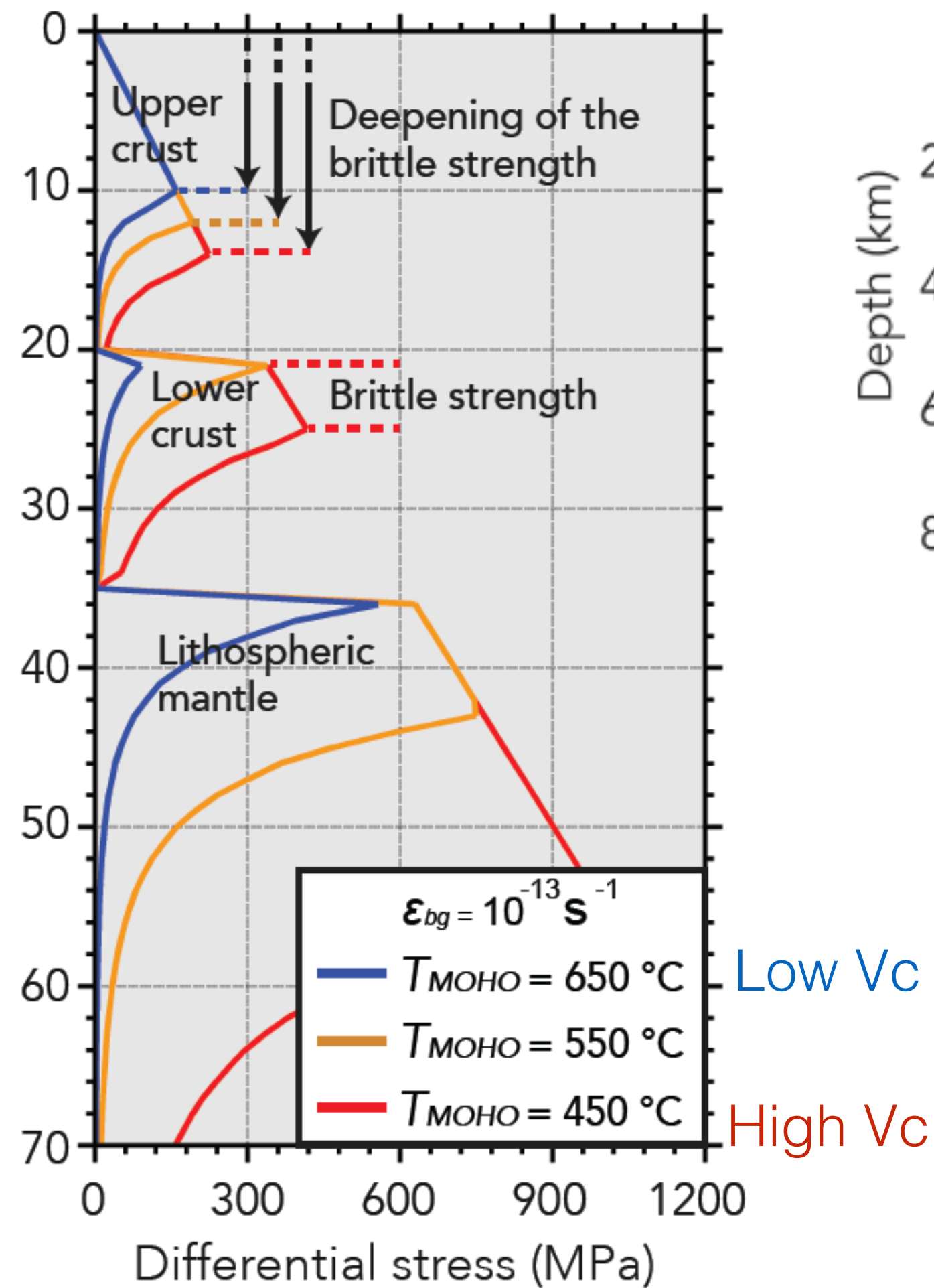
Role of convergence velocity highlights importance temperature

- Faster penetration of cooler temperatures to larger depths
- Larger brittle portion → larger and relative more larger events
- Relation and magnitude agree with regional observations



Role of convergence velocity highlights importance temperature

- Faster penetration of cooler temperatures to larger depths
- Larger brittle portion → larger and relative more larger events



» More ductile deformation in Alps, so lower seismic hazard than in Himalaya