Generation and transport of liquid water in the high-pressure ice layers of Ganymede and Titan

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Is there life somewhere else in the Universe?

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How do we find it?

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NASA: Follow the water

life might have started at the bottom of Earth's oceans

#### REVIEWS MICROBIOLOGY

Review Article | Published: 29 September 2008

#### Hydrothermal vents and the origin of life

William Martin , John Baross, Deborah Kelley & Michael J. Russell Noture Reviews Microbiology 6, 805–814 (2008) Download Citation #



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- Enceladus:
- geysers
- libration of outer layer



#### Europa & Enceladus

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#### Ganymede & Titan

- thick HP ice layer between silicates and the ocean
- is water & material exchange possible?



- full differentiation: Mol=0.310
- tenuous O<sub>2</sub> exosphere
- intrinsic magnetic field



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#### Titan

- partial differentiation: Mol=0.3414
- dense N<sub>2</sub>-CH<sub>4</sub>-Ar atmosphere
- $\rightarrow$  possibly ongoing exchange?



- dehydrated mantle
   + thick H<sub>2</sub>O layer
- ► H≥400 km: only ice VI



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#### Titan

- core of hydrated silicates
   + thinner H<sub>2</sub>O layer
- ► *H*~50-300 km: ices VI, V, III?



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- a)  $T = T_{\rm m}$ , heat transfer through HP ice not treated
- b) use of scaling laws based on the hot TBL instability
- effect of ice phases transitions on convection:
- a) linear stability analysis (Bercovici +, 1986; Sotin & Parmentier, 1989)
- b) 2d convection in ice VI and VII mantle (Journaux & Noack, 2015)

Choblet + (2017): solid-state thermal convection in 3d spherical

•  $T_{\max}$  and  $T_{\operatorname{avrg}} > T_{\operatorname{melt}}$  in the upper part



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- $T_{\max}$  and  $T_{\operatorname{avrg}} > T_{\operatorname{melt}}$  in the upper part
- $\rightarrow$  cold thermal boundary layer cannot exist in a solid state
- → two-phase mixture model (Bercovici +, 2001; Šrámek +, 2007)



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- different material densities  $\rightarrow$  2 mass balances
- phase separation  $\rightarrow$  2 linear momentum balances
- $\blacktriangleright$  thermal equilibrium between the two phases  $\rightarrow$  1 energy balance

1. Thermal convection in ice - incompressible extended Boussinesq approximation

mass balance of ice

$$\nabla \cdot \mathbf{v}_i = 0$$

► linear momentum balance of ice  $0 = -\nabla \Pi - \rho_i \alpha_i T g + \nabla \cdot \boldsymbol{\sigma}_i$ 

• energy balance of ice  

$$\rho_i c_i^p \frac{\partial T}{\partial t} + \rho_i c_i^p \mathbf{v}_i \cdot \nabla T - \rho_i T \alpha_i \mathbf{v}_i \cdot \mathbf{g} = \nabla \cdot (k_i \nabla T) + \boldsymbol{\sigma}_i : \nabla \mathbf{v}_i$$

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abla \cdot \left[ \phi(\mathbf{v}_{\mathrm{w}} \! - \! \mathbf{v}_{\mathrm{i}}) 
ight] = rac{\Delta 
ho}{
ho_{\mathrm{i}} 
ho_{\mathrm{w}}} \mathsf{\Gamma}_{\mathrm{m}}$$

linear momentum balance of ice

$$\mathsf{0} = - 
abla \mathsf{\Pi} - (1 - \phi) 
ho_{\mathrm{i}} lpha_{\mathrm{i}} \mathcal{T} oldsymbol{g} + 
abla \cdot (1 - \phi) oldsymbol{\sigma}_{\mathrm{i}} + 
abla \left| rac{(1 - \phi) \mu_{\mathrm{i}}}{\phi} (
abla \cdot \mathsf{v}_{\mathrm{i}}) 
ight|$$

• energy balance of mixture  $\rho_{i}c_{i}^{p}\frac{\partial T}{\partial t} + \overline{\rho c^{p} \mathbf{v}} \cdot \nabla T - \rho_{i}T\overline{\alpha \mathbf{v}} \cdot \mathbf{g} = \nabla \cdot (\overline{k}\nabla T) + (1-\phi)\boldsymbol{\sigma}_{i} : \nabla \mathbf{v}_{i}$ 

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- 3.  $\mu_{\rm w} \ll \mu_{\rm i}: \overline{\sigma} \to (1-\phi)\sigma_{\rm i}$ ,  $T_{\rm w} = T^{\rm m}: \overline{\rho\alpha} \to (1-\phi)\rho_{\rm i}\alpha_{\rm i}$ ,  $\overline{\rho c^{\rm p}} \to (1-\phi)\rho_{\rm i}c_{\rm i}^{\rm p} \to \rho_{\rm i}c_{\rm i}^{\rm p}$
- 4. Melting/freezing  $\rightarrow$  compressible mixture  $\rightarrow \nabla \cdot \mathbf{v}_i \neq 0 \rightarrow \zeta_i \sim \frac{\mu_i}{\phi}$

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 ρ<sub>i</sub>c<sup>p</sup><sub>i</sub> ∂T/∂<sub>4</sub> + ρc<sup>p</sup>v · ∇T - ρ<sub>i</sub>Tαv · g = ∇ · (k∇T) + Ψ̃

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- 4. Melting/freezing  $\rightarrow$  compressible mixture  $\rightarrow \nabla \cdot \mathbf{v}_i \neq 0 \rightarrow \zeta_i \sim \frac{\mu_i}{\phi}$

 $\rightarrow$  latent heat  $L_{\rm m}$  consumption/release

mass balance of mixture

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ightarrow latent heat  $L_{
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#### 5. Melt buoyancy: $\Delta \rho = \rho_i - \rho_w \neq 0$

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ho_{\mathrm{i}}lpha_{\mathrm{i}} T oldsymbol{g} + 
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abla \left\lfloor rac{(1\!-\!\phi)\mu_{\mathrm{i}}}{\phi} (
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ight
angle$$

► energy balance of mixture  $(\tilde{\Psi} = (1-\phi)\sigma_i:\nabla \mathbf{v}_i + \frac{1-\phi}{\phi}\mu_i(\nabla \cdot \mathbf{v}_i)^2)$  $\rho_i c_i^p \frac{\partial T}{\partial t} + \overline{\rho c^p \mathbf{v}} \cdot \nabla T - \rho_i T \overline{\alpha \mathbf{v}} \cdot \mathbf{g} + \Gamma_m L_m + \Gamma_s L_s = \nabla \cdot (\overline{k} \nabla T) + \tilde{\Psi}$ 

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 energy balance of mixture (Ψ̃ = (1-φ)σ<sub>i</sub>:∇v<sub>i</sub> + 1-φ/φμ<sub>i</sub>(∇·v<sub>i</sub>)<sup>2</sup>)

 α∂T
 α

$$\rho_{\rm i} c_{\rm i}^{\rm p} \frac{\partial T}{\partial t} + \overline{\rho c^{\rm p} v} \cdot \nabla T - \rho_{\rm i} T \overline{\alpha v} \cdot g + \Gamma_{\rm m} L_{\rm m} + \Gamma_{\rm s} L_{\rm s} = \nabla \cdot (\overline{k} \nabla T) + \tilde{\Psi}$$

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- ► linear momentum balance of ice  $0 = -\nabla\Pi - (1-\phi)\rho_i\alpha_i T g - \phi\Delta\rho g + \nabla \cdot (1-\phi)\sigma_i + \nabla \left[\frac{(1-\phi)\mu_i}{\phi}(\nabla \cdot \mathbf{v}_i)\right]$
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linear momentum balance of water  $\sim$  Darcy law

$$\mathbf{v}_{\mathrm{w}}-\mathbf{v}_{\mathrm{i}}=-rac{\kappa_{\phi}}{\mu_{\mathrm{w}}\phi}(1{-}\phi)\Delta
ho$$
g

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linear momentum balance of water ~ Darcy law

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 energy balance of mixture (Ψ = (1-φ)σ<sub>i</sub>:∇v<sub>i</sub> + (1-φ)φ/φ μ<sub>i</sub>(∇·v<sub>i</sub>)<sup>2</sup> + μ<sub>w</sub>φ<sup>2</sup>/K<sub>φ</sub>(v<sub>w</sub>-v<sub>i</sub>)<sup>2</sup>)

$$\rho_{\rm i} c_{\rm i}^{\rm p} \frac{\partial T}{\partial t} + \overline{\rho c^{\rm p} \boldsymbol{v}} \cdot \nabla T - \rho_{\rm i} T \overline{\alpha \boldsymbol{v}} \cdot \boldsymbol{g} + \Gamma_{\rm m} L_{\rm m} + \Gamma_{\rm s} L_{\rm s} = \nabla \cdot (\overline{k} \nabla T) + \Psi$$

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$$oldsymbol{v}_{
m w}-oldsymbol{v}_{
m i}=-rac{{\cal K}_{\phi}}{\mu_{
m w}\phi}(1{-}\phi)\Delta
hooldsymbol{g}$$

► mass balance of water  $\rightarrow$  advection of porosity  $\frac{\partial \phi}{\partial t} + \nabla \cdot (\phi \mathbf{v}_w) = \frac{\Gamma_m}{\rho_w}$ 

#### Material parameters

• ice permeable to water transport above a threshold porosity  $\phi_{\rm c}$ 

$$K_{\phi} = \begin{cases} k_{0}(\phi - \phi_{c})^{2} & \phi \geq \phi_{c} \\ 0 & \phi < \phi_{c} \end{cases}$$

temperature and depth dependent ice viscosity

$$\mu_{\rm i} = \mu_{\rm 0} \exp\left[\frac{Q}{R}\left(\frac{1}{T} - \frac{1}{T^{\rm m}(z)}\right)\right]$$



## Implementation

FEniCS (fenicsproject.org, Logg +, 2012; Alnaes +, 2015), MPI

ocean: free slip, free water outflow,  $T=T_m^o$ sides: free slip, water impermeable, thermally insulating silicates: no slip, water impermeable, heat flux  $q_s$ 

Η

H=200 km,  $\mu_{\rm i}$ =10<sup>15</sup> Pa s,  $q_{\rm s}$ =20 mW m<sup>-2</sup>



 $H{=}200$  km,  $\mu_{\rm i}{=}10^{15}$  Pa s,  $q_{\rm s}{=}20$  mW m  $^{-2}$ 

H=200 km,  $\mu_{\rm i}{=}10^{15}$  Pa s,  $q_{\rm s}{=}20~{\rm mW}~{\rm m}^{-2}$ 



▶ silicates:  $T = T^{m}$ ,  $\phi \sim \text{few } \%$ : volatiles leaching

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- ▶ interior: *T*<sup>av</sup> < *T*<sup>m</sup>: melt can freeze; upwellings: volatiles transport

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- ▶ silicates:  $T = T^{m}$ ,  $\phi \sim$  few %: volatiles leaching
- ▶ interior: *T*<sup>av</sup> < *T*<sup>m</sup>: melt can freeze; upwellings: volatiles transport
- ▶ ocean:  $T = T^{m}$ ,  $\phi \sim \phi_{c}$ : volatiles dissolution into the ocean

## Increasing HP ice layer thickness H ( $\sim$ increasing Ra)



## Increasing HP ice layer thickness H ( $\sim$ increasing Ra)



## Increasing HP ice layer thickness H ( $\sim$ increasing Ra)



## Increasing HP ice layer thickness H (~ increasing Ra)



## Increasing HP ice layer thickness H (~ increasing Ra)



# Results of parametric study



- degree of exchange decreases with
  - ► increasing ice layer thickness H (~ time)
# Results of parametric study



degree of exchange decreases with

- increasing ice layer thickness H ( $\sim$  time)
- decreasing ice viscosity µ<sub>i</sub>

# Results of parametric study



degree of exchange decreases with

- increasing ice layer thickness H ( $\sim$  time)
- decreasing ice viscosity µ<sub>i</sub>
- decreasing heat flux from silicates q<sub>s</sub>

# Long-term evolution modeling

- 2d model: strong dependence on input parameters:
- HP ice layer thickness H
- ice viscosity  $\mu_{
  m i}$
- silicates heat flux  $q_{
  m s}$

#### Long-term evolution modeling

- 2d model: strong dependence on input parameters:
- HP ice layer thickness H
- ice viscosity  $\mu_{\mathrm{i}}$
- silicates heat flux  $q_{
  m s}$
- Id thermo(-chemical) evolution model:
- scaling laws for rocky interior, HP ice layer, ocean, ice I crust
- $q_{
  m s}$  and H are part of solution
- $\mu_{
  m i}$  remains an input parameter

• fixed viscosity  $\mu_i$ , range of  $q_s$  and H



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 $\rightarrow$  critical  $q_{\rm s}^{\rm c}$  for bottom melting:  $q_{\rm s}^{\rm c} = \alpha H - \beta$ 



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```
\blacktriangleright Titan, \mu_{
m i}=10^{15} Pa s: q_{
m s}^{
m c}\propto H
```



- $\blacktriangleright$  Titan,  $\mu_{
  m i}=10^{15}$  Pa s:  $q_{
  m s}^{
  m c}\propto H$
- ▶  $\mu_{\rm i} = 10^{14} 10^{16}$  Pa s



- $\blacktriangleright$  Titan,  $\mu_{
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 $\rightarrow$  Titan: melting at silicates interface and leaching of <sup>40</sup>Ar?

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#### juice



# → JUPITER ICY MOONS EXPLORER

Exploring the emergence of habitable worlds around gas giants

Two mission phases: - Jupiter Tour (-2.5 yr): Jovian atmophere & magnetosphere Europa & Callisto flybys - Ganymede Tour (-1 yr): in Ganymede orbit

# Dragonfly

- rotocraft lander
- dozens of locations
- launch 2026, arrival 2034







Search for prebiotic chemical processes common on both Titan and Earth ...

#### Conclusions & perspectives

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# Thank you for your attention!

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