

Bayesian Parameter Inference in Instantaneous Mantle Flow with Plates

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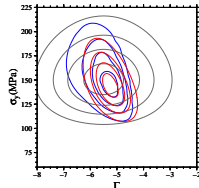
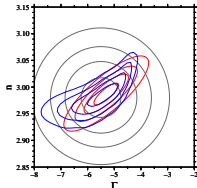
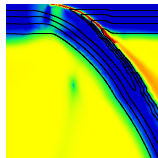
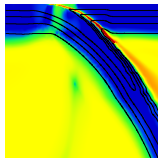
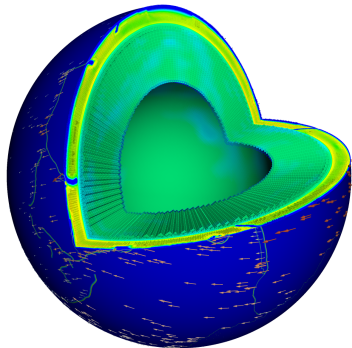
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Goal: Better understanding of fundamental mechanisms

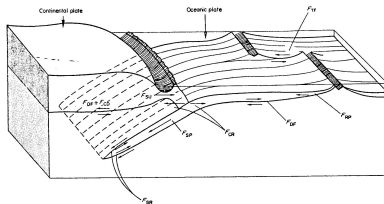
① Instantaneous mantle flow & associated plate tectonics with realistic parameters & resolutions to faulted plate boundaries ("the forward problem").

② Parameter/model inference from observations+mantle flow models.



What drives plate motions?

- ▶ How does slab pull really work?
- ▶ What's the role of back arc-spreading in global mantle motions?
- ▶ Can one infer the degree of mechanical plate coupling between subduction zones based on plate motions?
- ▶ What are the trade-offs of underlying forces [Conrad/Hager '99 using scaling arguments]:
 - ▶ buoyancy from slabs
 - ▶ resistance by inter-plate faults
 - ▶ drag by the underlying mantle
 - ▶ resistance by bending the oceanic lithosphere



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Target is to infer the most uncertain parameters (rheological parameters, plate coupling strength) to global observations (Bayesian inverse problem).

High-resolution global forward models

Equations for momentum/mass conservation

$$-\nabla \cdot \left[\mu(T, \mathbf{u}) (\nabla \mathbf{u} + \nabla \mathbf{u}^\top) \right] + \nabla p = \mathbf{f}(T)$$

$$\nabla \cdot \mathbf{u} = 0$$

\mathbf{u} ... velocity

p ... pressure

T ... temperature

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Rheology is shear-thinning with plastic yielding, and upper/lower viscosity bounds; exponential dependence on temperature:

$$\mu(T, \mathbf{u}) = \mu_{\min} + \min \left(\frac{\tau_{\text{yield}}}{2\dot{\epsilon}(\mathbf{u})}, w \min \left(\mu_{\max}, a(T) \dot{\epsilon}(\mathbf{u})^{\frac{1-n}{n}} \right) \right)$$

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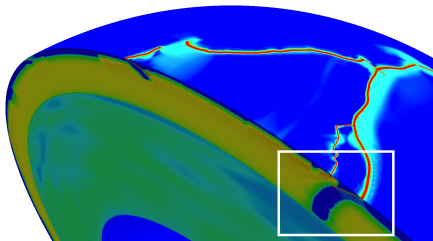
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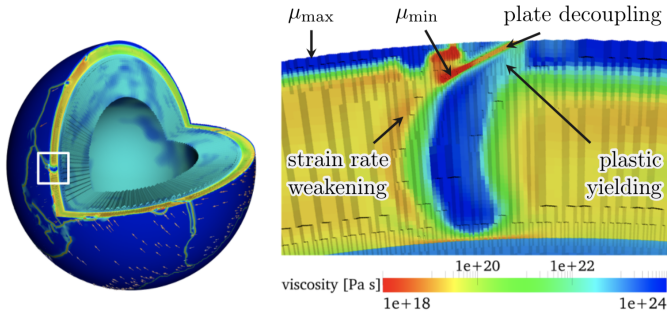
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Plates are modeled as high-viscosity fluid (low T); plate boundaries are narrow zones of weak viscosity (prefactor $\omega(x)$ controls plate coupling).



Solver and discretization

- ▶ Discretization with adaptive finite hexahedral elements
- ▶ Order 2 or 3 for velocity, and corresponding stable discontinuous pressure elements of lower order
- ▶ Mesh refinement used to resolve narrow weak zones, and dynamically weakening areas (hinges)
- ▶ Resolved model has about 500M dofs, smallest mesh elements 1-2km (runs on several K CPUs)

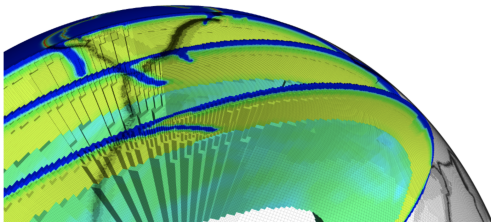


Solver and discretization

- ▶ Nonlinearity is treated with Newton's method; plastic rheology (von Mises) uses Newton modification method to improve convergence

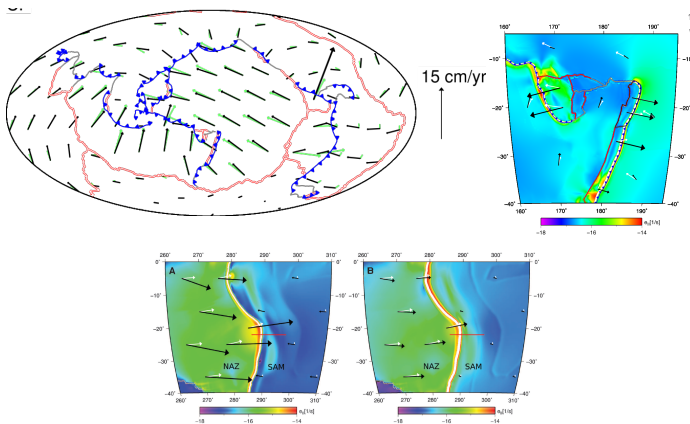
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- ▶ Linearized Stokes solved with BFBT Schur complement, and geometric+algebraic multigrid-preconditioned GMRES



Setup and resolution

- ▶ Free-slip boundary conditions
- ▶ Input is temperature structure and fault weak zones (details later)
- ▶ Shown below: Misfit with plate observations (in green), trench rollback, plate coupling (white are observations)



From forward towards inverse modeling

- ▶ Hand-tuning models to fit observations cumbersome, only possible for a few parameters
- ▶ It is subjective
- ▶ Even if good parameters are found, we have no information about their stability or other possible parameters

Inverse problem theory (deterministic or Bayesian) is a systematic approach to infer parameters and to quantify our confidence in them.

Inverse Problems: Inference from data+model

Bayesian inversion for uncertain parameters \mathbf{m}

- ▶ **Observations \mathbf{d}** : plate motions; average viscosity below lithosphere, . . .
- ▶ **Parameters \mathbf{m}** : strain rate exponent n , yield stress τ_y , activation energy, prefactors in lower/upper mantle, weak zone factors w_i (about ~ 20)

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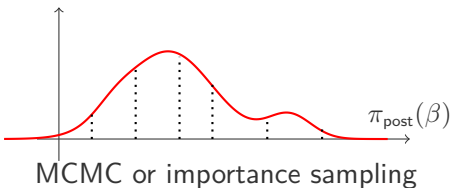
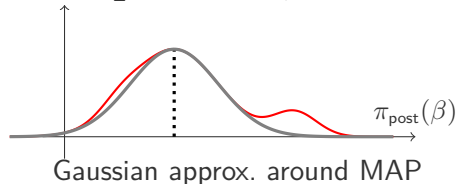
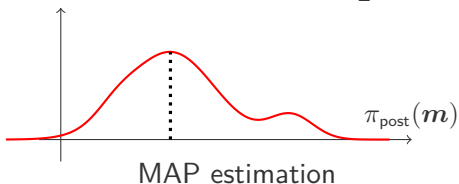
- ▶ **Model $f(\mathbf{m})$** : Notation for nonlinear Stokes solve given \mathbf{m} , "parameter-to-observable map"
- ▶ Additive Gaussian data noise: $\mathbf{d} = f(\mathbf{m}) + \mathbf{e}, \mathbf{e} \sim \mathcal{N}(\mathbf{0}, C_n)$.
- ▶ The **posterior density** is (using Bayes' Theorem):

$$\pi_{\text{post}}(\mathbf{m}) \propto \exp \left(-\frac{1}{2} \|f(\mathbf{m}) - \mathbf{d}\|_{C_n^{-1}}^2 - \frac{1}{2} \|\mathbf{m} - \mathbf{m}_0\|_{C_{\text{pr}}^{-1}}^2 \right).$$

Approximations of the posterior pdf $\pi_{\text{post}}(\cdot)$

With these assumptions (Gaussian prior and noise), the posterior is is:

$$\pi_{\text{post}}(\mathbf{m}) \propto \exp\left(-\frac{1}{2}\|f(\mathbf{m}) - \mathbf{d}\|_{C_n^{-1}}^2 - \frac{1}{2}\|\mathbf{m} - \mathbf{m}_0\|_{C_{\text{pr}}^{-1}}^2\right)$$



In high dimension, we use...

- ▶ ... MAP estimation (i.e., Stokes-constrained optimization)
- ▶ ... Gaussian approximation of posterior

Inversion Algorithms

MAP optimization: Maximizing the posterior equals minimizing its negative log, i.e.:

$$\min_{\mathbf{m}} F(\mathbf{m}) := \frac{1}{2} \|f(\mathbf{m}) - \mathbf{d}\|_{C_n^{-1}}^2 + \frac{1}{2} \|\mathbf{m} - \mathbf{m}_0\|_{C_{pr}^{-1}}^2$$

Compute gradient of $F(\cdot)$ using adjoint method!

Adjoint method to compute $F'(\mathbf{m})$:

1. For given \mathbf{m} , solve nonlinear Stokes equations:

$$\begin{aligned} -\nabla \cdot \left[\mu(T, \mathbf{u}, \mathbf{m}) (\nabla \mathbf{u} + \nabla \mathbf{u}^\top) \right] + \nabla p &= \mathbf{f}(T) \\ \nabla \cdot \mathbf{u} &= 0 \end{aligned}$$

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2. Solve an adjoint problem for the linear (!) adjoint velocity \mathbf{v}

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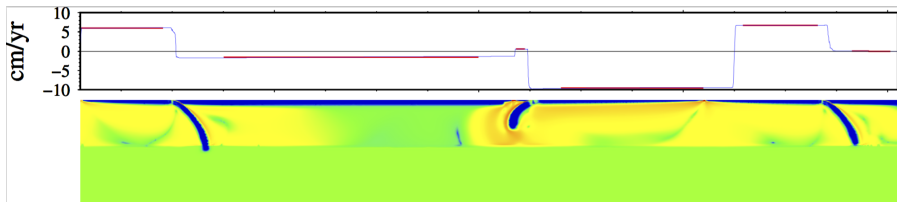
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Requires one additional linear solve only!

Inversion Algorithm

- ▶ We also compute directional second derivatives of $F(\cdot)$ using similar approach
- ▶ Solve optimization (MAP point) using iterative descent method
- ▶ At MAP estimate, compute Gaussian approximation of posterior using that inverse Hessian approximates covariance matrix

Example I: Synthetic models



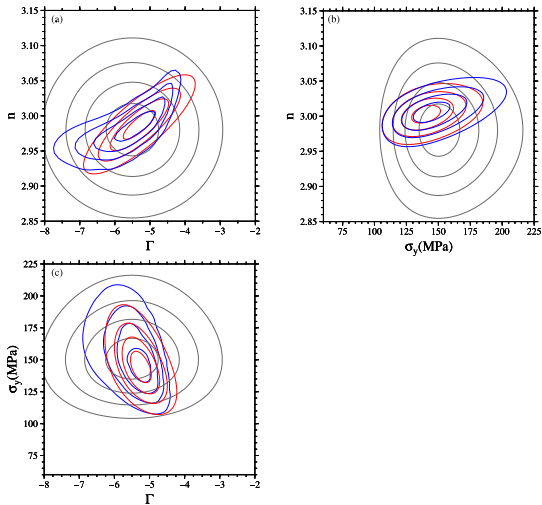
Parameters

- ▶ Parameters: plate coupling strength, global rheology parameters (≤ 5 params)

Observation data

- ▶ Data: plate velocities
- ▶ Computed from forward simulation (i.e., synthetic data)

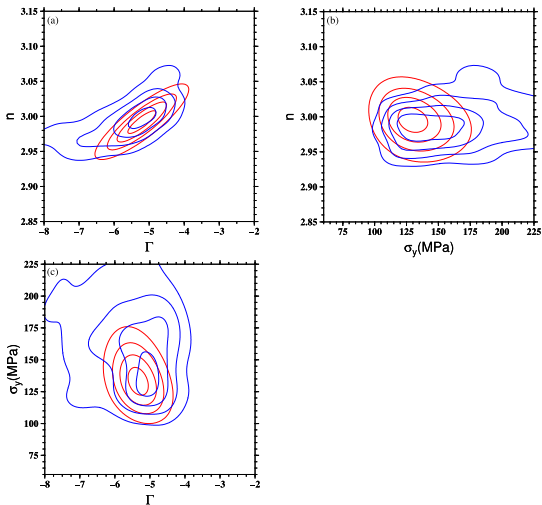
Example I: Inversion of plate coupling & rheology params



2D pairwise conditionals for Γ , n , σ_y . Prior, Gaussian approximation and true posterior.

- ▶ Inference for plate coupling Γ (w) and rheology parameter n from surface velocity data.
- ▶ Gaussian prior distribution for parameters.
- ▶ Posterior approximation uses inverse Hessian as covariance matrix.
- ▶ Full posterior computed using Delayed Rejection Adaptive Metropolis (DRAM) sampling.

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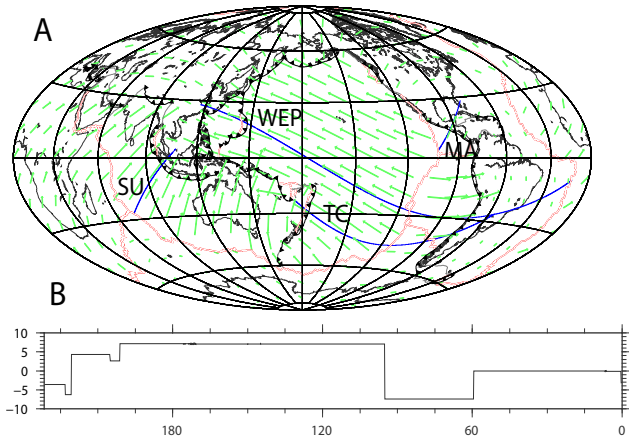
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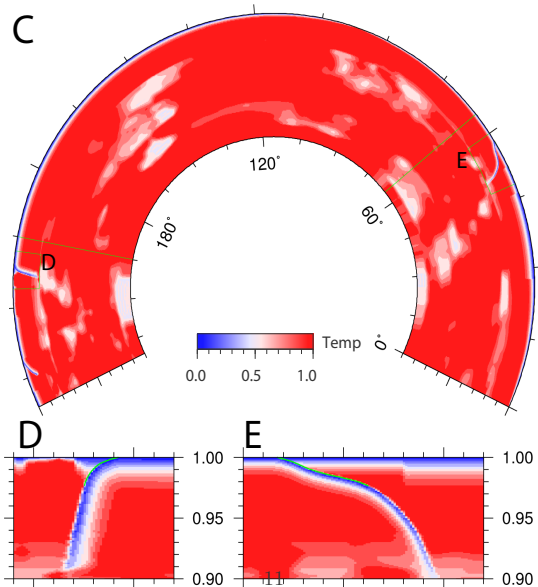
Updated input data and 2D inversions

- ▶ Mega-thrust interface using Slab1.0
- ▶ Thermal structure from seismic models and convergence rate of slabs
- ▶ Lithosphere model
- ▶ Slab structure blend with tomography-based models
- ▶ Morvel56



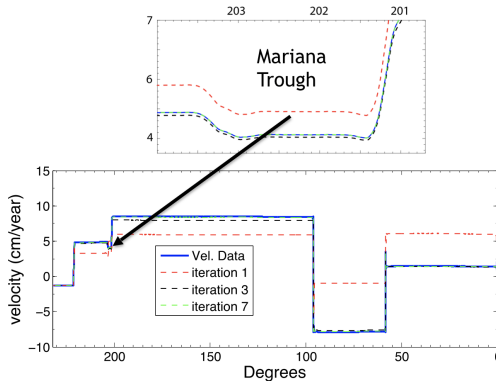
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Example II: Inversion results for WEP slice

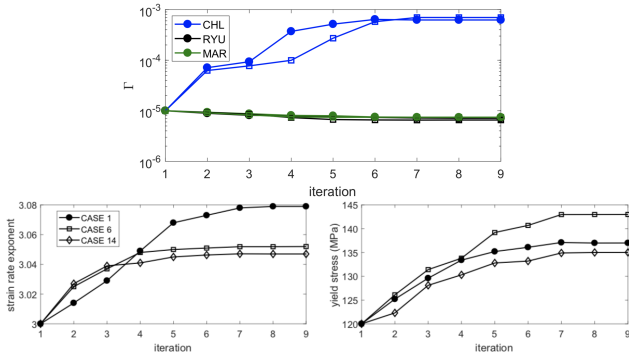
This is a 3D thin slice



- Fitting the observations after about 7 inverse problem iterations.

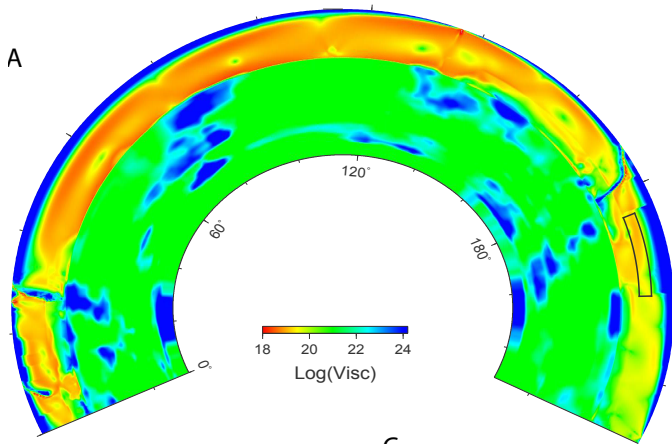
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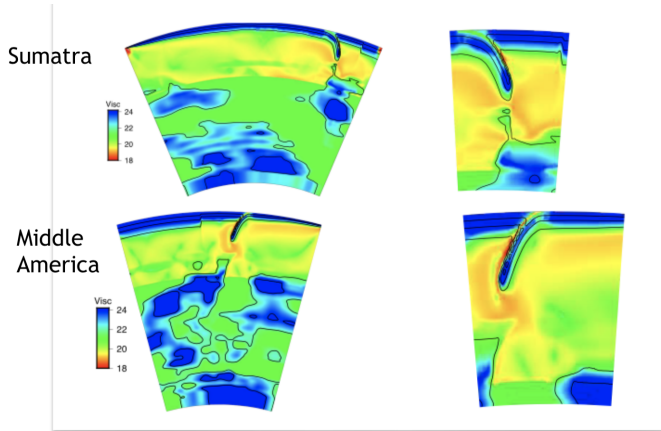
- ▶ Fitting the observations after about 7 inverse problem iterations.
- ▶ Convergence history for weak zone parameters (Chile, Ryukyu, Mariana), for strate rate exponent and yield stress.

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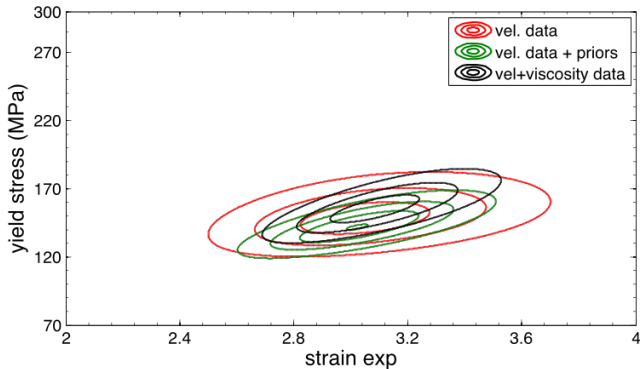
Effective viscosity. Here, besides plate motion data, we have also used the average viscosity in the boxed region.

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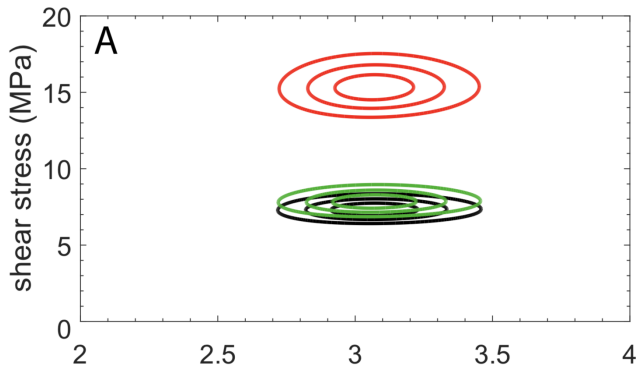
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Two-dimensional distribution between shear stress and n with and without using average viscosity data.

Example II: Inversion results for WEP slice



Shear stress estimates (with uncertainty) for Chile (red), Ryukyu (black) and Marianas (green), versus strain rate coefficient n .

Summary & Perspectives

Solvers

- ▶ High-resolution adaptive discretization
- ▶ Full Newton solver for instantaneous Stokes with strain-rate weakening and yielding
- ▶ Consistent temperature field (slabs) and faults;
Self-consistent flow field

Inversion

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- ▶ Gives confidence/uncertainty regions, and trade-offs
- ▶ Gradients using adjoints are efficient and often do not require much new implementation

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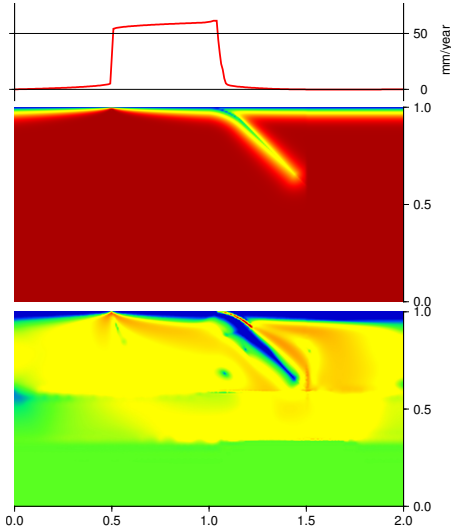
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Thanks!

Extra Slides

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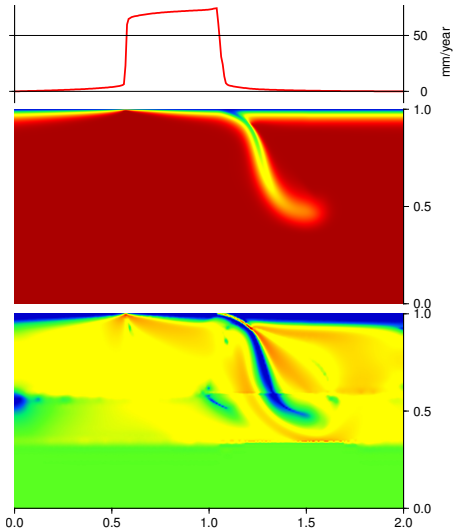
Forward simulation



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- ▶ Gradient computation requires solving state equation, and adjoint equation backwards in time.
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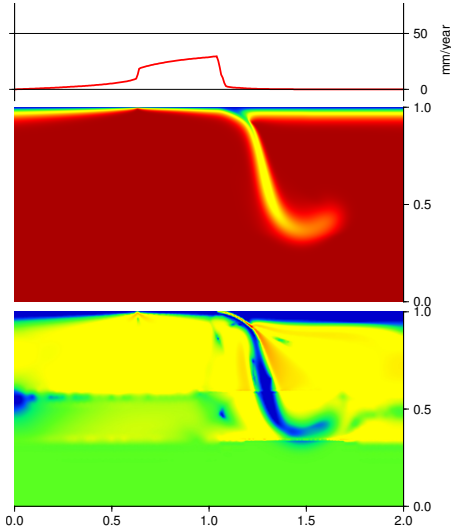
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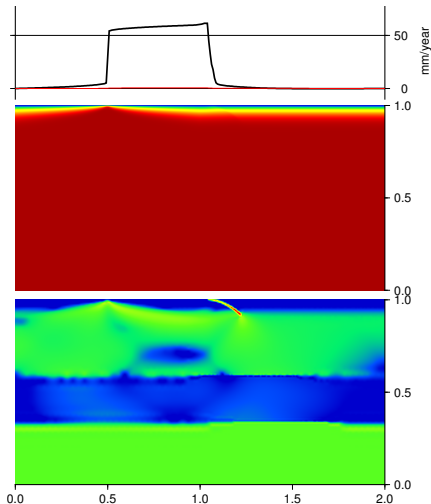
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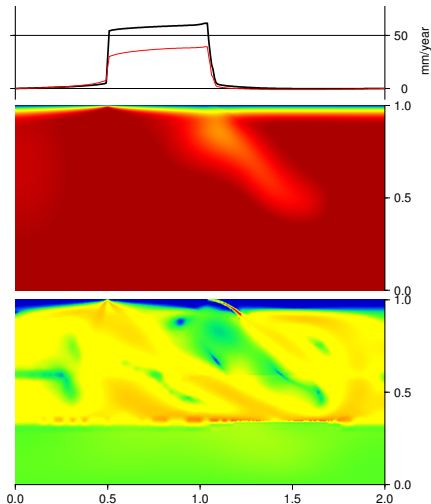
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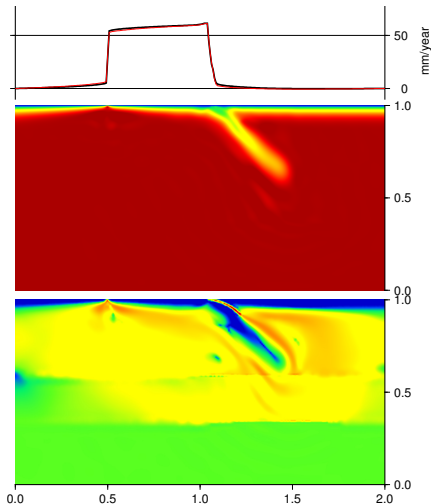
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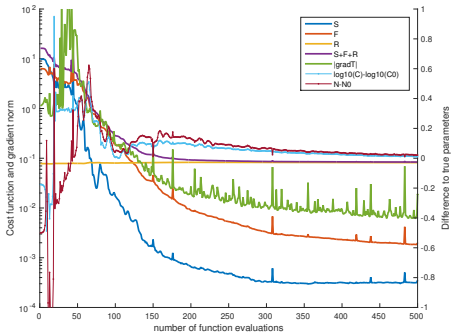
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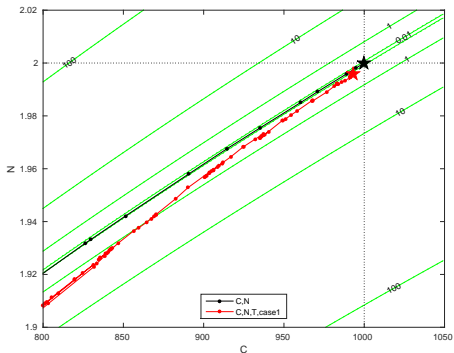
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- ▶ I'd like convergence to be faster
- ▶ Part of reason for slow convergence is flat valley in cost landscape indicating trade-off between parameters